

source of the magnetic anomaly at Strahan is at considerable depth and of no interest. The anomalies recognised by Corbett et al (1982) include Zones 8, 9, 13 and 14. Zone 8 Asbestos Point to Nobby Creek area; is due to ultramafic, intermediate to acid volcanic rocks and gabbros. Zone 9 Birch Inlet; is assigned to basalt however the size of the anomaly is typical of serpentinite. Zone 13 Mouth of King River; is most probably due to a buried extension of Zone 8, with the source of the anomaly due to ultramafic or basic volcanics. Zone 14 Ocean Beach; could be similar to Zone 13 but displaced to the north by a large regional fault, e.g. a north-easterly extension of the Henty Fault.

Rampore & King 1968, Geol Soc Aust Trans 1962 p 129, 327

The other main area of ultrabasic rocks is located either side of the Huskisson River Syncline (new name - Figure 11 & Plates 3 & 5), and has an intense coincident magnetic anomaly. The mapping of the area by Brown (1982) showed that wherever Eocambrian or Cambrian serpentinites or amphibolites were mapped there was a coincident, large amplitude magnetic anomaly. The Wilson River serpentinite, Zone 45 Pieman River to Parsons Hood; is actually two magnetic linears due to layered pyroxenite-peridotite-serpentinite. The anomaly on the eastern side of the syncline, Zone 44 Pieman River to Webb Creek; has a similar source. The qualitative interpretation of the western serpentinite belt is consistent with an intrusive sill dipping to the north-east. The inclination varies from steeply dipping to the south-east of the Alfred Fault, to a shallow dip north-west of the Fault, (Figures 12 A, 12 B & 12 C). The interpretation of the eastern serpentinite is not as simple. The source of the magnetic anomaly could be interpreted as a magnetic sill intruded into an anticlinal structure, which outcrops as far north as the Huskisson River and then plunges further to the north. South of 380000N the eastern limb of the anticline has been removed. To the north, the Mt. Ramsay skarn-type magnetic anomalies are relatively shallow compared to the serpentinites at depth (Figure 12 b). This model is far from perfect and is open to major variation, however it matches the ideas presented by Zarzavatjian (1965). The main difference between the two models is that the recent mapping of the area has identified Precambrian basement either side of the Huskisson Syncline.

PLUTONIC INTRUSIVES

During the mid-Devonian Tabberabberan Orogeny the region was intruded by a number of post tectonic granites. These are mainly contact aureole biotite granites and hornblend-biotite granodiorites including S and I types, which appear to have been emplaced at relatively low pressures largely by diapiric intrusion and roof lifting, (Solomon 1981). The granites vary greatly in size from the Meredith Granite in the centre of the area to the smaller granitic bodies intruded in the Mt. Read Volcanics. The magnetic expression of each of the plutons will be discussed separately.