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INTERIM GEOLOGICAL REPORT

ON THE SOUTH WEST PORTION OF EXPLORATION

LICENCE 13/65. SOUTH WEST TASMANIA

NOVEMBER 1965 - MAY 1966

W.D.M. HALL

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INTERIM GEOLOGICAL REPORT  
ON THE SOUTHWEST PORTION OF EXPLORATION LICENCE 13/65  
SOUTHWEST TASMANIA  
NOVEMBER 1965 - MAY 1966  
by W.D.M. Hall

SUMMARY

Strongly deformed older Precambrian rocks occupy an anticlinal core, extending from the south coast to the Gordon River. They include quartzite, quartz schist, pelitic and sub-pelitic schist, coarse schist, gneiss, and amphibolite, which have undergone two metamorphisms and at least three periods of folding. The earlier metamorphism gave rise to garnet-bearing rocks.

Younger Precambrian rocks rest unconformably on the older Precambrian, and consist of 10,000 feet of moderately deformed turbidites and conglomerate.

Cambrian rocks include acid volcanic and sedimentary rocks, and occur on the west and south-east flanks of the Precambrian core. Ordovician rocks include Owen conglomerate, Caroline Creek sandstone, and Gordon limestone and occur in the Mt. Osmund and Olga Synclines, respectively west and east of the Precambrian, and also in the south east. They are unconformable on both Cambrian and Precambrian rocks. Silurian

003

quartzite and slate occur in the Olga Syncline, the eastern margin of which is formed by a major north-west striking fault.

The Permian, Triassic, and Jurassic consist of flat-lying tillite, siltstone, sandstone, and dolerite, which rest unconformably on the older rocks.

Copper and tin surface mineralisation has been located, and is associated with granite, pelitic schist, gneiss, and Cambrian Volcanics.

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STRATIGRAPHYOLDER PRECAMBRIAN

Older Precambrian rocks are the most widely exposed, and consist of strongly deformed, regionally metamorphosed quartzites and schists. A generalised stratigraphic sequence includes:

Massive quartzite at top

Quartz schist

Pelitic and sub-pelitic schist and phyllite

Coarse grained schist and gneiss with amphibolite intrusions.

This sequence, although sometimes inverted, is fairly well established from the east side of Port Davey, the west side of Bathurst Harbour, and at Nye Bay. Individual beds of the various rock types are, however, often interbedded.

The quartzite and quartz schist fall in the chlorite zone of the greenschist metamorphic facies, while the coarse schist, gneiss and amphibolite generally fall in the almandine zone. The pelitic schist occurs in both the chlorite and almandine zones. A biotite zone appears to be absent.

The coarse grained schist and gneiss is exposed at Wilson Bight and Ketchem Bay on the south coast, in the Kelly Basin area, and along the west coast to Nye Bay, and is up to 1,500 feet thick.

The rocks on the south coast are massive and pale green with a lustrous sheen on fresh surfaces. They are gneissic, have an indistinct lineation, and contain abundant albite porphyroblasts in a finely-foliated quartz-muscovite-chlorite (-epidote?) matrix.

In the Kelly Basin area the most common coarse schist is a finely-layered, white and grey mottled rock with a gneissic texture. It contains quartz-albite augen in irregular muscovite and rare biotite and chlorite layers, and chloritised garnet porphyroblasts up to one inch across.

Varieties include fine-grained, equigranular quartz-albite-muscovite-biotite rocks with a poorly developed foliation, and very finely-layered biotite-quartz rocks with gneissic texture and garnet prophyroblasts up to half an inch across.

An unusual rock, probably an unfoliated granulite, is interbedded with coarse schists at Davey Head, south of Kelly

Basin. It is medium-grained, equigranular, mottled white and dark green, and consists of quartz, plagioclase, rare actinolite, pale almandine, and crenulated chlorite.

Lineations and mesoscopic folds are usually indistinct or absent in the coarse schists.

At a number of localities along the west coast the coarse schists are intruded by small, concordant, but discontinuous bodies of amphibolite. These are dense, black, and usually structureless rocks. They are fine grained, and consist of almandine garnet and actinolite, with rare biotite, chlorite and plagioclase. The amphibolites are probably highly metamorphosed basalts and dolerites.

Pelitic schist is exposed in the Kelly Basin and Port Davey areas, along the west coast to north of Nye Bay, at Joe Page Bay, the north west corner of Bathurst Harbour, and on the south coast between Ketchem Bay and New Harbour Point, on the east side of Cox Bight, and from Red Point to the Iron bound Range, and is up to 8,500 feet thick.

It is dark grey to black and very finely foliated but in the almandine zone becomes more massive and gneissic.

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It is locally phyllitic in the chlorite zone. Lamination is generally poorly developed, but two mica-crinkle lineations have been observed.

The mesoscopic folds are dominantly  $F_2$ . They are isoclinal, usually with attenuated limbs, and have thickened, moderately sharp crests.  $F_3$  folds are open crinkle and kink folds.

The layering is  $S_1$  (=So?) with  $S_2$  slightly oblique to it, and forming the axial planes at the  $F_2$  folds. Locally  $S_3$  has developed at right angles to both  $S_1$  and  $S_2$  as kink planes associated with the  $F_3$  folds.

In thin section the pelitic schist is fine-grained, and has a foliated matrix of quartz, muscovite, biotite, and albite, and minor chlorite and graphite. The snowball almandine garnets are usually chloritised, and indicate later retrograde metamorphism.

Sub-pelitic schist is distinguished from the pelitic schist by its well-developed foliation. It is up to 5,000 feet thick, and occurs from Muleahy Bay, to north of Nye Bay, on the east side of Port Davey and in the Kelly Basin area, and as thin bands interbedded with quartz schist at Stephens

008

Bay and New Harbour.

The rocks are fine-grained with a black and white segregation layering (i.e. foliation) usually 1 inch thick, but up to one inch thick. They are quartz-albite-muscovite-chlorite rocks with minor biotite flakes and minute chloritised garnet porphyroblasts in the almandine zone.

Up to three lineations occur, the most common being a minute crinkle lineation, emphasised in the almandine zone by small, randomly distributed biotite flakes aligned parallel to it.

Mesoscopic folds developed in the sub-pelitic schists usually belong to  $F_2$ , and are generally similar to those in the pelitic schist, but with slightly more rounded hinges and less attenuated limbs.

The layering is  $S_1$ , and is often cut by a strongly developed  $S_2$  which forms the axial planes of the  $F_2$  folds.

Quartz schists are widely distributed, and are probably best illustrated in outcrops along the north-east coast of Bathurst Harbour and north-west coast of Joe Page Bay, and are in excess of 4,000 feet thick.

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They are white to light green, and are finely layered in bands one inch to one foot thick, with the original cross-bedding sometimes preserved. They include quartz-muscovite and quartz-albite muscovite schists with minor to rare chlorite and epidote. The muscovite is commonly altered to sericite.

Lineations are usually poorly developed, but two may occur, particularly where  $S_1$  and  $S_2$  coincide. The earlier lineation is formed by streaked mineral elongation, and has been refolded by a later mica-crinkle lineation.

A variety of fold styles occur,  $F_1$ ,  $F_2$ , and  $F_3$ , all being preserved. Possible  $F_1$  folds are flattened recumbent folds with partly thickened hinges.  $F_2$  folds are most common, and are open to isoclinal with rounded hinges. Rare examples have thickened hinges and attenuated limbs.  $F_3$  folds are kink folds, sometimes slightly overturned, and have a rare axial plane fracture cleavage.

At one locality on the Crossing River  $F_2$  folds were observed to be refolding  $F_1$  folds.

The layering is  $S_1$ , which is plainly parallel to  $S_0$ , the original bedding.  $S_2$  has developed as an axial plane

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cleavage, particularly where the rocks are strongly folded.

Massive quartzite is also widely distributed, and is well illustrated by outcrops on the east coast of Bathurst Harbour and in the hills immediately to the east, and at New Harbour Range. It is over 10,000 feet thick.

The quartzite is white to pale green and occurs in beds six inches to three feet thick, which are frequently cross-bedded, finely-laminated and ripple-marked, and often contain thin bands of finely-bedded quartz schist.

A distinct pink quartzite occurs in the Frankland and Giblin Ranges immediately south of Lake Peddar, and along the Cracroft River. It may represent a stratigraphic horizon, but its precise position is uncertain.

The quartzites are composed of a fine-grained quartz mosaic with rare muscovite and chlorite. Lineation is indistinct or absent, and the layering is  $S_1 = S_0$ .  $F_1$  folds are rare partly compressed recumbent folds with slightly thickened hinges.  $F_2$  folds are most common, and have a pronounced asymmetry and slightly thickened hinges.

In the Herold Mountains north-west of Bathurst Harbour,

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recumbent  $F_1$  folds with an amplitude of about 100 feet have been slightly refolded by  $F_2$  folds which have a pronounced axial plane cleavage.

Possible Older Precambrian rocks are exposed in the area between the head of the Cracroft and New Rivers and along Federation Creek, and consist of a north-east striking belt of chlorite grade rocks unlike any previously seen in the Exploration Licence area.

They dip steeply west, and consist of about 5,000 feet of massive, finely-laminated pink quartzite, overlain by 100 feet of finely-foliated green schist, then at least 8,000 feet of green, brown, and purple phyllite with conglomerate bands up to 100 feet thick. The conglomerates contain rounded, slightly stretched quartzite boulders up to one foot across.

The phyllite is locally mesoscopically folded and has a steeply plunging mica-erinkle lineation. The layering is  $S_1 = S_0$ .

To the west these rocks are probably faulted against south-east trending quartzite and quartz schist.

Younger Precambrian.

Younger Precambrian rocks occur in a complex outlier faulted into older Precambrian rocks at Bathurst Harbour and along Bathurst Channel, and at the southern end of the Olga Syncline near the junction of the Davey and Crossing Rivers.

In the Bathurst Harbour area the Young Precambrian rocks consist of about 10,000 feet of turbidites and conglomerate. The conglomerate rests unconformably on Older Precambrian schists and quartzite, and is conformably overlain by the turbidites. They are generally steeply dipping and have a well-developed slaty cleavage. The conglomerate pebbles are sometimes crumpled and distorted, while the turbidites are commonly altered to quartz-muscovite-chlorite schist and phyllite, and are often mesoscopically folded.

The most continuous section of Younger Precambrian rocks is exposed along the south coast of Bathurst Harbour between Moulters Inlet and Celery Top Islands, and consists of the following rocks:

Turbidites - at top.	
Buddy conglomerate	+ 1,500 feet
Graded bedded conglomerate, sandstone and siltstone	} 1,000 feet
Fine conglomerate	600 feet
Graded beds as above	1,500 feet
Medium conglomerate	200 feet
Graded beds as above	2,000 feet
Massive conglomerate at base	1,600 feet

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The turbidites are excellently exposed along Bathurst Channel and the south coast of Bathurst Harbour. The graded bedded members consist of fine, quartz pebble conglomerate, quartzose greywacke sandstone, and dark siltstone in beds up to eight inches thick. They are often finely laminated and have well preserved sedimentary structures. The muddy conglomerate has a dark schistose matrix which contains angular to rounded pebbles of quartzite and quartz schist up to six inches across.

Rare bands of white quartzite and dark grey slate are interbedded with the turbidites.

The conglomerate is best exposed on the east side of Horseshoe Inlet, and at Mounts Rugby, Berry and Misery. At the base it contains angular fragments of quartz, pink and white quartzite, quartz schist and rare graphitic schist up to two feet across, in a matrix of angular quartz grit. The matrix becomes finer-grained higher in the formation, and the pebbles more rounded. Current-bedding and cut and fill structures sometimes occur, and the conglomerate is locally hematite-stained.

Younger Precambrian rocks are exposed on either side of the Davey River at the southern end of the Olga syncline, where they rest unconformably on Older Precambrian schists, and are unconformably overlain by Ordovician rocks.

The sequence is similar to that at Bathurst Harbour, and includes a basal breccia overlain by a conglomerate of unsorted, rounded quartzite pebbles up to four inches across in a fine quartz sandstone matrix, and blue grey to white

fine sandstone which is strongly folded.

Faulted outliers of the basal breccia occur along the Crossing River and at the head of Spring River.

Cambrian

Cambrian rocks have been mapped in two areas, on the west coast north of Elliott Bay, and along the south coast adjacent to New River.

The Cambrian rocks exposed in the area from Elliott Bay to the Wanderer River can be divided into two lithological groups; a volcanic sequence to the east, and a sedimentary sequence exposed along the coast. The relation between these two groups is not precisely known, but they are locally interbedded and may be lateral equivalents.

The Volcanic rocks are exposed in scattered outcrops on the open country south and east of the Mt. Osmond Syncline, and along the coast between the Lewis River and Copper Creek. They include pyroclastics, acid lavas, and tuffs, and are generally schistose and have a well developed slaty cleavage.

The pyroclastics and tuffs are generally similar, differing only in grain-size and phenocryst development.

They are commonly fine-grained, green, very finely bedded, and often show cross bedding and graded bedding.

The acid lavas include porphyritic and fine-grained varieties. They are green to pale grey, have a fine-grained quartz-feldspar groundmass, and glassy quartz phenocrysts.

The sedimentary sequence is composed dominantly of grey to dark brown, thinly-bedded siltstone and greywacke with minor slate, chert, and tuffaceous sandstone. The finer beds have a strongly developed north-striking slaty cleavage and are locally strongly mesoscopically folded.

At the mouth of the Wanderer River and at Sandy Point, the rocks are graded bedded, and probably part of a turbidite sequence.

Rocks inferred to be Cambrian are exposed at Telopea Point at the southern end of the Amy range. They are about 500 feet thick, and consist of a basal breccia of quartzite and quartz schist blocks resting unconformably on Precambrian quartzite, overlain by pebble conglomerate and massive argillite with thin pebble and sandstone bands. The argillite is slightly schistose.

West of New River Cambrian rocks rest unconformably on Precambrian schists, and occupy a south-east plunging syncline which is exposed on Mt. Louisa and Ironbound Ranges, and along the south coast between Deadman's and Prion Bays.

At Deadman's Bay, the basal beds consist of 50 feet of massive breccia containing angular blocks and pebbles of quartzite, quartz schist, and graphitic schist up to six feet across, and thin lenses of grit and fine conglomerate. Near Mt. Louisa and on the Ironbound Range the breccia consists of blocks of quartzite and quartz schist up to two feet across and rare blocks of graphitic schist 50 feet across, with beds of grey, gritty quartz sandstone.

Overlying the breccia is about 2,000 feet of massive, blue-grey siltstone in beds two to four feet thick, with thinner beds of gritty sandstone containing scattered quartz pebbles. The siltstone has a pronounced vertical slaty cleavage, and is commonly phyllitic.

At the west end of Prion Bay, the siltstone is unconformably overlain by Ordovician sediments.

East of New River Cambrian rocks are exposed along the coast from the mouth of New River to just west of Pretty's

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Point. They dip steeply east, and are faulted against Precambrian dolomite to the west and are unconformably overlain by Ordovician rocks to the east.

The following section was measured:

- At top                    Mottled maroon-green grit and fine conglomerate with angular pebbles                    100 feet
- blue-grey, calcite-veined siltstone with thin bands of graded bedded sandstone, grit and muddy conglomerate                    2,500 feet
- Conglomerate with unsorted, rounded inclusions of granite, dunite, dolomite, amphibolite, and quartz up to two feet across in a brown-grey sandstone and locally maroon mudstone matrix. Thin bands of graded bedded sandstone, siltstone and chert                    800 feet
- finely laminated, blue-grey mudstone with thin bands of brown sandstone, and rare beds of conglomerates and grit                    1,200 feet
- Interbedded cycles of boulder and red-green conglomerate beds, gritty sandstone, and finely laminated, light grey chert and breccia                    1,200 feet

Massive, indurated conglomerate and breccia of quartz, quartzite, and quartz schist pebbles in a matrix of quartz grit and sandstone in beds up to six feet thick, with beds of fine conglomerate, grit, and finely laminated, cross bedded sandstone up to three feet across. A conspicuous horizon of dolomite boulders occurs at the top of the formation + 300 feet

The conglomerate and breccia at the base of this section strongly resembles that at the base of the Cambrian west of New River.

### Ordovician

The Ordovician rocks in the area mapped are divided into three formations.

Gordon Limestone, at top  
 Caroline Creek Sandstone  
 Owen conglomerate

Owen Conglomerate was mapped only in the area close to Moores Valley. It is exposed in the Mt. Osmund Syncline, at Thirkell Hill and as a thin strip just west of Moores Lookout.

In the Mt. Osmund Syncline the formation is over 5,000 feet thick, and consists of rounded siliceous pebbles in beds

up to 30 feet thick, interbedded with bands of grit, sandstone and siltstone, some showing current bedding. These rocks overlie Cambrian volcanics, probably unconformably, in the east, and are faulted against them on the west limb of the syncline.

West of Moores Lookout the formation is 200 feet thick. It rests unconformably on Precambrian schist and quartzite, dips 45 degrees west, and is faulted at its western margin. The basal rocks are hematite - stained quartzites and tubicolar quartzites, and are overlain by conglomerate and grit beds two feet thick. The conglomerate contains hematite pebbles and irregular patches of fine-grained hematite.

Caroline Creek Sandstone is exposed in the Olga Syncline, along the south coast and in the Upper Cracroft Valley.

It occurs as a thin strip along the west limb of the Olga Syncline from near Frederick Hill to the junction of the Davey and Crossing Rivers, and for a short distance on the east limb parallel to the Davey River. At View Hill it consists of east dipping tubicolar quartzite and hematite - stained quartzite which rest unconformably on the underlying schist, and mark the base of the Palaeozoic succession in the area.

Around the southern end of the syncline the formation consists of grey-weathering friable quartz sandstone with thin pebble bands and scattered quartz pebbles, and forms a prominent strike ridge which has been strongly drag folded.

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South of the Crossing River and along Spring River similar rocks occur as a series of isolated, faulted outliers. The basal unconformity is well-exposed at a number of localities.

On the south coast Caroline Creek sandstone is exposed at the west end of Prion Bay, and at Pretty's Point at the west end of Surprise Bay.

At Prion Bay it consists of well-bedded, medium-grained quartz sandstone with rounded quartz pebbles. It unconformably overlies Cambrian siltstone and dips south-east at 30 degrees. At Pretty's Point it consists of about 500 feet of hard, brown-weathering, grey, fine sandstone with a slight sulphur efflorescence, in beds two to eighteen inches thick, and unconformably overlies Cambrian turbidites and dips 40 degrees east.

The formation also occurs as a faulted outlier at the junction of the East and West Crocroft Rivers, where it dips 35 degrees south-west, and rests unconformably on Precambrian schist.

Gordon Limestone was observed on the south coast at the east end of Surprise Bay, and is reported by H.E.C. Geologists to occur in rare outcrops on the floor of the Olga-Hardwood Valley.

At Surprise Bay it is about 2,000 feet thick, and consists of alternating fine-grained, dark-grey limestone and black, micaceous siltstone in beds two to twelve inches thick.

The limestone was also observed during helicopter

reconnaissance on the east side of the New River Lagoon.

Silurian

Rocks inferred to be Silurian are shown as occurring along the east side of the Olga-Hardwood Valley. They are well-bedded quartzites and slates which rest conformably on Gordon Limestone and are faulted against Precambrian quartzite.

Permian, Triassic, Jurassic

Permian, Triassic and Jurassic rocks were mapped only in the south east corner of the Exploration Licence area. An excellent section, dipping gently east, was examined along the south coast between Shoemakers Point and South Cape.

Tillite at the base (i.e. western end) of the section rests with marked angular unconformity on Gordon Limestone. It is approximately 1,300 feet thick, and consists of a matrix of indurated dark siltstone with abundant, unsorted debris up to boulder size. The formation is probably of marine origin.

Between the south coast and the Cracroft River the base of the tillite has been interpreted from aerial photos and brief helicopter reconnaissance.

The tillite is conformably overlain by 100 to 200 feet of dark siltstone, probably equivalent to the Woody Island siltstone.

Above the siltstone is a dolerite sill about 800 feet

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thick, which forms a cap on most of the peaks north to Huon River. It is well-exposed at South Cape and on Precipitous Bluff.

At South Cape the dolerite has a thin, impersistent cover of yellow-grey quartz sandstone with scattered quartz pebbles, and thin bands of fine conglomerate.

### Igneous Rocks

Acid igneous rocks are exposed at Elliott Bay, South West Cape, and Cox Bight.

At Cox Bight, a small, fine-grained biotite granite body intrudes Precambrian quartzite, while at South West Cape a strongly-sheared, coarse-grained granite with thin aplitic dykes intrudes quartz schist. Both bodies have small contact aureoles of quartz-plagioclase - biotite hornfels associated with them.

At Elliott Bay granite intrudes pyroclastics, lavas, and quartz-porphyry. It is locally very <sup>coarse</sup> fine grained, with individual crystals up to two inches long, but also includes aplitic and porphyritic varieties. Near its outer margins the granite becomes finer grained and schistose.

Schistose quartz porphyry is exposed at the east end of Elliott Bay. It intrudes Cambrian volcanics and is probably intruded by the granite. It extends inland in a

northerly direction for about 12 miles, and is faulted against Precambrian rocks along its eastern margin.

### STRUCTURE

The main structural feature is an anticlinal core of strongly deformed Younger and Older Precambrian rocks extending from the south coast, through Port Davey along the west coast and north almost to the Gordon River. It is almost 20 miles wide at its southern end, but becomes narrower northwards and is concealed beneath Ordovician rocks.

Within this core are large  $F_1$  recumbent folds which have been strongly isoclinally folded during  $F_2$  deformation. Metamorphism accompanied both deformations. The oldest, or tectonically oldest rocks are exposed in the Kelly Basin area, and along the west coast to Nye Bay.

West of the anticlinal core are the Ordovician and Cambrian rocks of the north-pitching Mt. Osmond syncline, and to the east is the Olga syncline containing Ordovician and Silurian rocks.

The eastern boundary of the Olga syncline is formed by a major fault extending from north of the Gordon River south east to the head of the Cracroft River.

In the New River area there is another syncline with Cambrian and Ordovician rocks, cut by a major fault extending

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along the valley of New River. Cambrian and Ordovician rocks east of the fault are apparently up faulted, while the overlying and undeformed Permian, Triassic, and Jurassic rocks have been apparently down faulted about 2,000 feet. The fault probably has a large transcurrent displacement.

ECONOMIC GEOLOGY

Surface mineralisation containing copper and tin has been observed.

The tin occurs in quartz schist and quartzite and in small alluvial deposits at Cox Bight and Melaleuca Inlet. The Cox Bight field is associated with granite, which may underlie a considerable area in the extreme south-west of the Exploration Licence area, and also have some bearing on the significant geochemical tin anomaly in the Hannant Creek area west of Melaleuca Inlet.

Copper occurs in graphitic schist and gneiss at Kelly Basin and on the west coast immediately north west of Kelly Basin. The geochemical results for this area are not yet available, but one rock sample spectrographically assayed indicated Cu greater than 10,000 p.p.m.

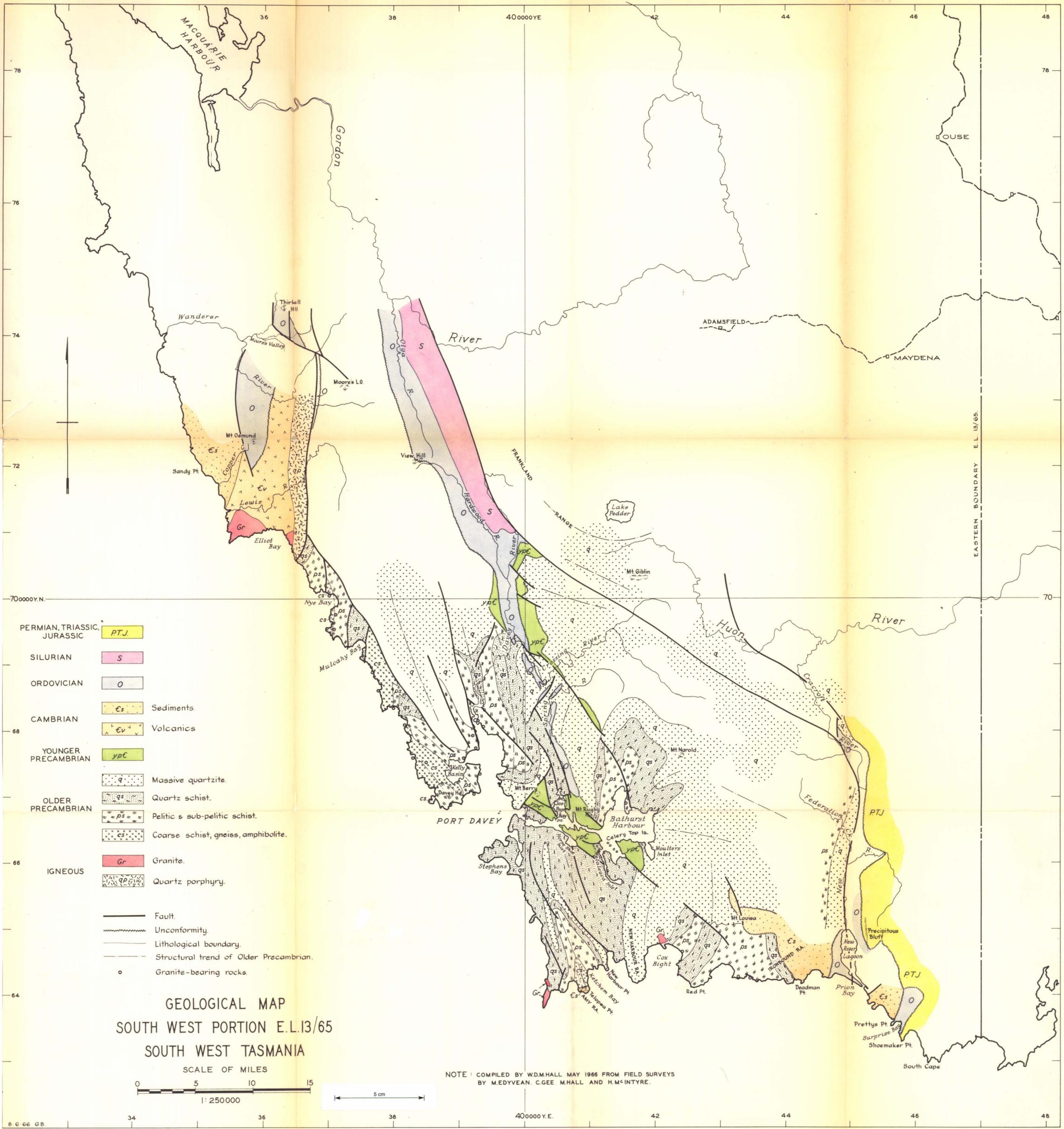
Copper also occurs in Cambrian volcanics at Elliott Bay and on the west coast just north of Low Rocky Point. Large pyrite veins are associated with these occurrences.

Significant geochemical zinc anomalies occur in the pelitic schist east of Cox Bight.

The most interesting economic considerations are the possibility of extensive granite at shallow depth in the south west giving rise to tin areas, the possibility of hydrothermal mineralisation of the pelitic and coarse schists associated with metamorphism, and the possibility of hydrothermal mineralisation associated with granitic intrusion in the Cambrian volcanics.

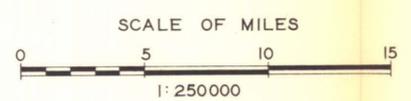
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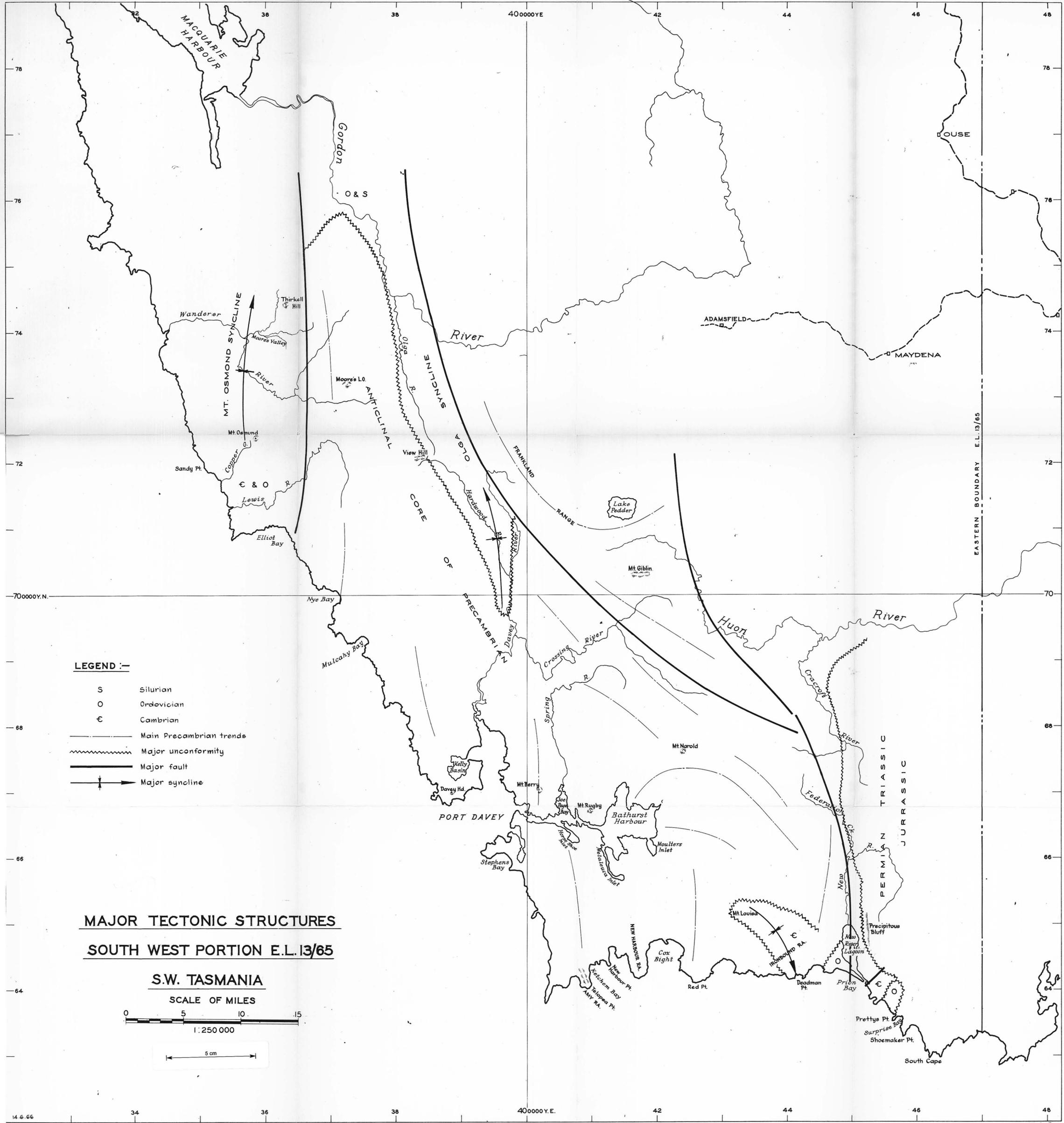


- PERMIAN, TRIASSIC, JURASSIC PT.J
- SILURIAN S
- ORDOVICIAN O
- CAMBRIAN
  - Cs Sediments.
  - Ev Volcanics
- YOUNGER PRECAMBRIAN ypC
- OLDER PRECAMBRIAN
  - q Massive quartzite.
  - qs Quartz schist.
  - ps Pelitic & sub-pelitic schist.
  - cs Coarse schist, gneiss, amphibolite.
- IGNEOUS
  - Gr Granite.
  - qp Quartz porphyry.
- Fault.
- ~ Unconformity.
- Lithological boundary.
- Structural trend of Older Precambrian.
- o Granite-bearing rocks.

**GEOLOGICAL MAP**  
 SOUTH WEST PORTION E.L.13/65  
 SOUTH WEST TASMANIA



NOTE: COMPILED BY W.D.M.HALL MAY 1966 FROM FIELD SURVEYS  
 BY M.E.DYVEAN, C.GEE, M.HALL AND H.M. INTYRE.



- LEGEND :-**
- S Silurian
  - O Ordovician
  - € Cambrian
  - Main Precambrian trends
  - ~ Major unconformity
  - Major fault
  - ↔ Major syncline

**MAJOR TECTONIC STRUCTURES  
SOUTH WEST PORTION E.L.13/65**

**S.W. TASMANIA**

