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COMPAGNIE GÉNÉRALE DE GÉOPHYSIQUE

MICROFILMED

THE CONSOLIDATED SYNDICATE

OPEN FILE

GEOPHYSICAL SURVEYS

AT MOUNT TYNDALL (TASMANIA)

- LAKE ROLLESTON
- LAKE SELINA
- GOOSENECK

January - February 1971

Mission 501-18-80

C.G.G.

26-28 Manning Street, SOUTH BRISBANE.

ABSTRACT

In January and February 1971, C.G.G. carried out a geophysical survey in the Mount Tyndall (Tasmania) area for The Consolidated Syndicate. Apparent Resistivity, Apparent Chargeability and Self Potential measurements were carried out with IPR7 time domain receivers.

Three different areas were surveyed - the Lake Rolleston Grid, the Lake Selina Grid and the Gooseneck Grid.

At Lake Rolleston, measurements were carried out with a gradient array, a 50 foot receiving dipole and a 50 foot spacing between stations, in order to obtain an accurate structural map of the bedrock beneath the moraine. Structural features were successfully brought out. A north/south conductive strip runs across the area between two resistant areas. A transverse fault intersects the structure - its throw is about 400 feet. The I.P. measurements brought out the southern extension of anomaly A1 which is already drilled further north. High Apparent Chargeability readings coincide with the western resistant area. Weak anomalies exist within the conductive area. A drillhole has been recommended on A20 which coincides with a conductive axis. Further geological reconnaissance has been recommended with regard to the western anomalous resistant area.

On the Lake Selina Grid, the northern extension of anomaly A1 has been followed over a distance of 5,000 feet. The I.P. anomaly coincides with a conductive axis and a magnetic anomaly. It extends further north, beyond a contact between the volcanics

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and a Jukes Breccia body. The anomaly is stronger near the contact on the southern side. A -400mv S.P. anomaly indicates that the mineralization is strongly pyritic and four drillholes have been recommended.

On the Gooseneck Grid, the southern extensions of anomalies A1 and A5 have been followed. The total length of A1 is 4,500 feet. It extends further southward beneath the Owen Conglomerate A1 is due to disseminated magnetite, pyrrhotite, pyrite and chalcopyrite in acid volcanics.

A5 is due to a layer of black shale.

One drillhole has been recommended.

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INTRODUCTION

From the 11th January to the 14th February, COMPAGNIE GENERALE DE GEOPHYSIQUE carried out geophysical surveys in the Mt. Tyndall (Tasmania) area for THE CONSOLIDATED SYNDICATE.

Three areas were surveyed, the Lake Rolleston Grid, the Lake Selina Grid and the Gooseneck Grid. These areas are located on a Cambrian compartment containing sediments and volcanics (Mt. Read Volcanics) striking north/south between the Dora Rift to the east and the Owen Rift to the west. East of the Dora Rift, the Precambrian basement outcrops as the quartzitic Sticht Ranges. West of the Owen Rift, the Cambrian is unconformably overlain by thick layers of the Ordovician Owen Conglomerate.

Several adits, pits and trenches have been dug as early as the end of the nineteenth century and showings of copper sulfides are known in all three areas. Sphalerite and galena are known at Gooseneck and galena at Selina. Pyrite, pyrrhotite and magnetite are common.

Previous geophysical surveys have been carried out by RIOTINTO (1957-1960), using airborne magnetometry, ground magnetometry, electro-magnetism, gravity and self potential in the Red Hills-Gooseneck Area, and by C.G.G. (1969-1970) in the Lake Rolleston and Red Hills areas using I.P., resistivity and S.P.

Since it is expected that any economic ore-body in this area

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will be within sheared Mt. Read Volcanics, then the aim of exploration is to define the limits of these volcanics and to outline structural features (faults, warpings, folds, ...).

All of the Rolleston area and parts of the Selina and Gooseneck areas are overlain by varying thickness of Pleistocene glacial overburden and conglomerate scree so that Apparent Resistivity measurements are the only method of outlining geological boundaries and structural features. Induced Polarization was used because it is the only geophysical method able to record anomalies over buried, disseminated, non-magnetic sulfides.

1. FIELD AND GEOPHYSICAL CONDITIONS

1-1. Location - Access - Terrain Conditions :

ROLLESTON GRID and SELINA GRID

These areas are located 5 miles south of Mt. Murchison and approximately 9 miles east of Mt. Dundas. The Rolleston Grid extends northward from Lake Rolleston to the southern limit of the Selina Grid which lies to the east of Lake Selina.

Access to the area is by a 17 mile track from the Murchison Highway that is suitable only for four-wheel-drive vehicles.

The Rolleston Grid is crossed by numerous moraine ridges covered with a low, heath-type vegetation. Movement along the lines is easy. Slopes on the Selina Grid vary from flat to very steep and the vegetation varies between button grass and dense rain forest so that movement along the lines is easy in some places and difficult in others. Progress was considerably slowed by steep slopes and thick vegetation on parts of lines 144N, 136N and 128N.

GOOSENECK GRID

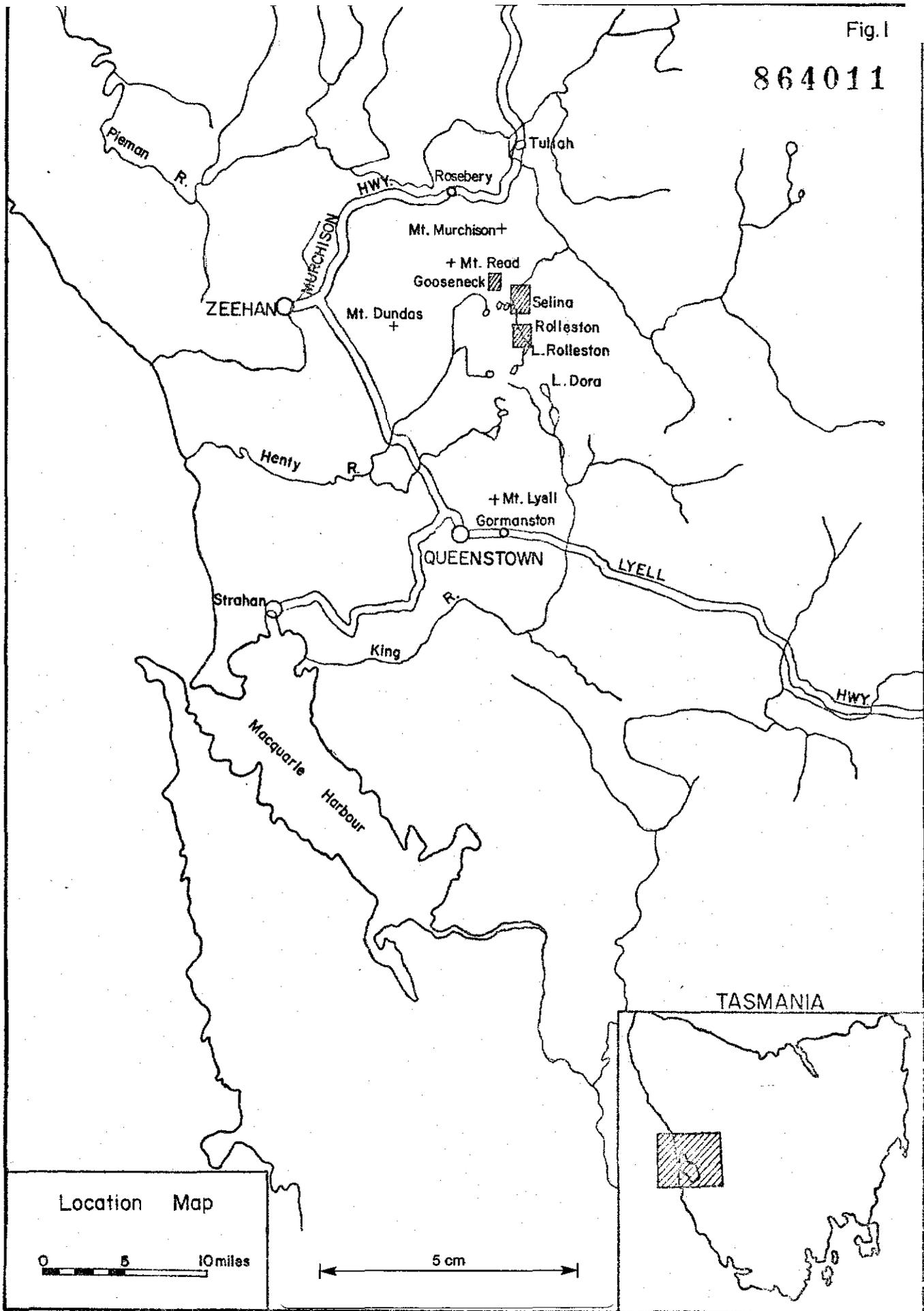
The Gooseneck Grid lies 6 miles south-southeast of Rosebery and 13 miles east of Zeehan. Access is by an 18 mile track suitable only for four-wheel-drive vehicles from the Murchison Highway.

The surveyed area is cut by a deep watercourse and covers the scree slope at the foot of conglomerate cliffs, so that movement along the lines is difficult in places. However the southern

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Fig. 1

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Location Map

0 5 10 miles

5 cm

TASMANIA

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part of the grid includes button-grass swamp and low ridges where measurements are easy and progress rapid.

1-2. Electrical Conditions :

ROLLESTON GRID

On the Rolleston Grid the bedrock is overlain by a Pleistocene glacial overburden of varying thickness having a resistivity of around 1000 Ω m.

For a current flow of 2 amperes the measured ohmic drops over a 50 foot receiving dipole ranged up to several tens of millivolts.

SELINA GRID

Some areas of the Selina Grid are overlain by glacial sediments while in other, more elevated areas the bedrock is exposed. Generally the overburden is shallow.

For an input current of around one ampere the measured ohmic drop using a 200 foot receiving dipole varied from several tens to several hundreds of millivolts.

GOOSENECK GRID

On the Gooseneck Grid conditions are similar to the Selina Grid conditions.

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2. FIELD OPERATIONS AND STATISTICS

2-1. Field Operations :

Field operations were carried out by three technicians and five field assistants. During the period 11th January to 14th February four days were lost due to rain and two days due to mechanical breakdown. An accident to one of the technicians caused the loss of 2 days.

On the Rolleston area four 6,560 foot current lines were used for the I.P. gradient array coverage.

A pole-dipole array was used on the Selina Grid with a potential dipole length of 200 feet and distances from current-input to receiving-dipole centre of 900 feet and 1,100 feet.

The pole-dipole array was also used on the Gooseneck Grid with receiving dipoles of 200 feet and current-input to receiving-dipole centre distances of 100 feet and 300 feet.

2-2. Statistics :

2-2-1. Composition of the Crew :

- One Party Chief : G. CLARIDGE
- Two Observers : P. BAXTER
- : A. IDIRIAN

After 23rd January, P. Baxter was replaced by B. DIGBY.

- Five Field Assistants; two supplied by C.G.G. and three by The Consolidated Syndicate.

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COMPAGNIE GÉNÉRALE DE GÉOPHYSIQUE

7.

GOOSENECK

- 2 Electrical Soundings
- 299 stations : 598 measurements of Apparent Chargeability,
Apparent Resistivity and Self Potential
- Total length of Profiles : 29,900 feet
- Duration of Survey : 10 days, of which 2 were
lost due to rain

3. LAKE ROLLESTON

3-1. Apparent Resistivity : (Plate 1)

Figure 2 presents 3 Electrical Soundings carried out on the Lake Rolleston Grid. From the surface downward the electrical markers are :

- a conductive overburden (100 ohm-m) less than 10 feet thick : soil
- a resistant layer ($\rho > 1000$ ohm-m) : moraine
- a conductive layer ($\rho < 25$ ohm-m) : glacial clay and/or weathered bedrock.
- a resistant basement ($\rho > 1000$ ohm-m).

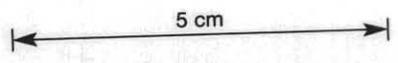
The depth of the resistant basement may reach 150 feet at E.S. 1, 250 feet at E.S. 2 and 300 feet at E.S. 3. A north/south conductive strip crosses the area (see Plate 1). Correlations from profile to profile are clearly visible - they are certainly due to structural features of the bedrock.

The Apparent Resistivity level is higher on the western resistant area than on the eastern resistant area. The eastern resistant area corresponds to the Cambrian volcanics; the western resistant area most probably corresponds to the Ordovician Owen Conglomerate

An east/west fault is visible between profiles 32N and 24 N - the throw of contours is 400 feet.

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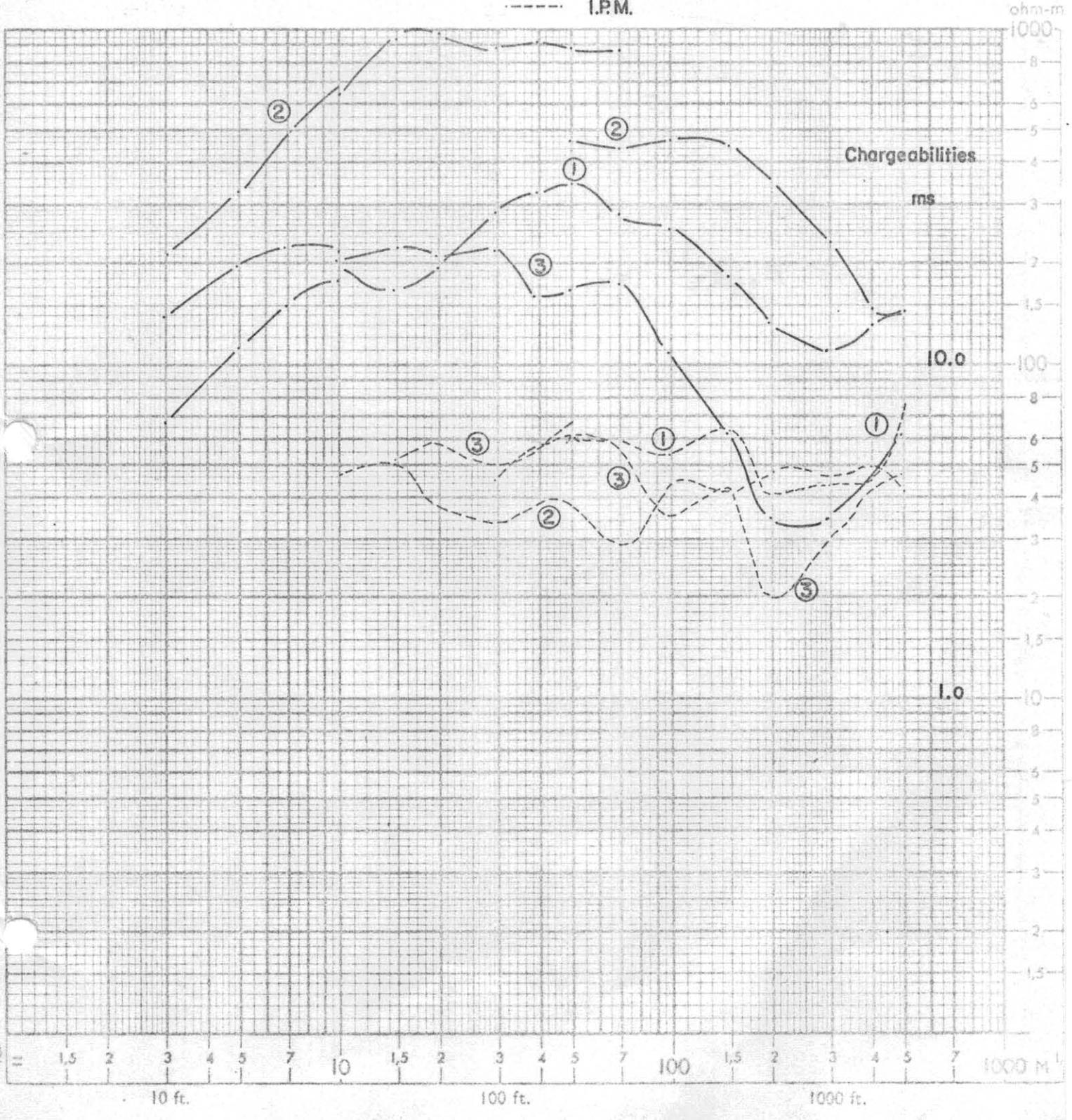
Fig 2



ROLLESTON GRID

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- E.S.① 24N-3300W
- E.S.② 16N-2000W
- E.S.③ 40N-3300W
- ρ
- - - I.P.M.



AB 2 = 1.5 2 3 4 5 7 10 1.5 2 3 4 5 7 100 1.5 2 3 4 5 7 1000 M

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3-2. Apparent Chargeability : (Plates 2 and 3)

The I.P. expanders and Electrical Soundings plotted on Figure 2 shows that the bedrock is polarized for current electrode separations larger than 2,500 feet. A current electrode separation of 6,560 feet was selected.

The amplitude of anomalies does not exceed 15 ms - the largest anomalies are located within resistant areas. A cancellation of a statistical correlation between Apparent Resistivities and Apparent Chargeabilities is obtained for $M_c = M + 5.60 C_{AB}$ where

M is the Apparent Chargeability

$$C_{AB} = \frac{AB}{2} = \frac{1000}{\rho} \text{ mho}$$

(ρ = Apparent Resistivity in ohm-m).

Plate 3 is a contoured map of M_c . It brings out a few narrow anomalies in the conductive areas and reduces the amplitude of anomalies in the resistant areas.

Anomaly numbers follow the 1969-70 survey series of numbers (see "Geophysical Survey in Mount Tyndall (Tasmania) E.L. 9/66 - December 8 1969 - February 1970).

A1 : A1 is the southern extension of the complex anomaly A1 detected in 1970. There are two main axes (see Plate 4) which are interrupted by the transverse fault.

A17 : A17 is the eastern extension of A1 south of the transverse fault. The anomaly thins out past profile 16N.

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A1 and A17 coincide with resistant areas. They do not appear to be due to a mineralization richer than the mineralization intersected in drillholes north from profile 40N.

A18 : A18 is located in the western resistant area, along the western edge of the conductive strip. A18 widens near the transverse fault. These features indicate that A18 may be due to a sulfide mineralization located near the Owen Conglomerate contact.

A19 : A19 appears to be the extension of A9 north of the transverse fault.

A20 : A20 is barely visible on Plate 2. The compensation (Plate 3) for the overburden influence confirms that there is definitely an anomaly which coincides with a conductive axis.

A21 : A21 appears to be the southern extension of A21 south of the transverse fault.

A22 : A22 is associated with a resistant axis within the conductive strip. There are two positive axes.

A23 : A23 is a small narrow anomaly which coincides with a conductive axis.

A24, A25 : A24 and A25 coincide with the high Apparent Resistivities on the western edge of the surveyed area.

The compensation for the influence of the overburden does not erase them. As mentioned in paragraph 3-1 the high Apparent Resistivities seem to indicate that the bedrock beneath the moraine is the Owen Conglomerate.

All Owen Conglomerate outcrops intersected by I.P. profiles in the Mt. Tyndall area are barren and do not give rise to I.P. anomalies. However, several anomalies due to disseminated sulfides in the Mt. Read Volcanics are located along the conglomerate contact.

The gentle Apparent Chargeability gradient suggests that I.P. anomalies A24 and A25 may be due to polarizable bodies located beneath the Conglomerate but this is merely an hypothesis. An accurate depth determination is impossible as only parts of the eastern flanks of I.P. anomalies extend over the surveyed area

3-3. Self Potential :

No S.P. anomaly has been detected.

3-4. Conclusions and Recommendations :

The detailed gradient array Apparent Resistivity coverage carried out on the Lake Rolleston Grid brought out a structural map of the area in spite of a moraine overburden thicker than 100 feet. Three main units are clearly visible:

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- a resistant area to the east corresponding to the Mt. Read Volcanics.
- a north/south conductive strip in the middle, corresponding to shales or perhaps the Gordon Limestone.
- a very resistant area to the west probably corresponding to the Owen Conglomerate.

A transverse east/west fault runs across the area between profiles 32N and 24N.

Ten I.P. anomalies have been detected - all of them, except three small narrow anomalies, coincide with resistant areas. There is no S.P. anomaly.

The main I.P. anomaly, A1, is the southern extension of an anomaly detected during a previous survey.

The southern extension of A1 does not show any variation when compared with the part which has already been drilled.

We would recommend drilling one drillhole on :

A20 : A20 is a weak anomaly but it coincides with a conductive axis. Its cause is unknown. A drillhole on A20 would determine the nature of the conductive marker and its significance with respect to possible mineralization.

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- Drill collar at 2,100 ft. west from the base line.
- Dip : 50° W.
- Length : 600 ft.

Drilling on A24 and A26 may be considered if their geological and structural setting seem favourable.

4. LAKE SELINA

4-1. Magnetometry : (Plate 5)

All magnetic anomalies are conformable - their axes are approximately north/south. Maximum amplitudes vary between 2000 and 2500 gammas. All magnetic bodies are almost outcropping. Floats of oxidized magnetite can be seen near Lake Selina on profile 160N.

The main magnetic anomalies are located along the edges of a Jukes Breccia body. Smaller, irregular anomalies are located above the Jukes Breccia body. Other anomalies are elongated and apparently due to stratiform bodies within the volcanics.

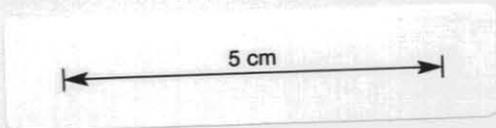
4-2. Apparent Resistivity : (Plates 6 and 7)

Electrical Soundings (see Figure 3) show that the conductive overburden is thin, even on the low ground - its thickness does not exceed 60 feet. The bedrock resistivity exceeds 1000 ohm-m except at some particular anomalies.

A main north/south conductive axis runs across the surveyed area. From profile 120N to profile 144N the anomaly increases and widens - it is wider and stronger for the largest pole-dipole array (Receiver 1 - Plate 6) and thus it is due to a conductive body, deeply rooted in the bedrock. The conductive anomaly does not show any interruption where it intersects the limit between the volcanics and the Jukes Breccia.

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Fig 3

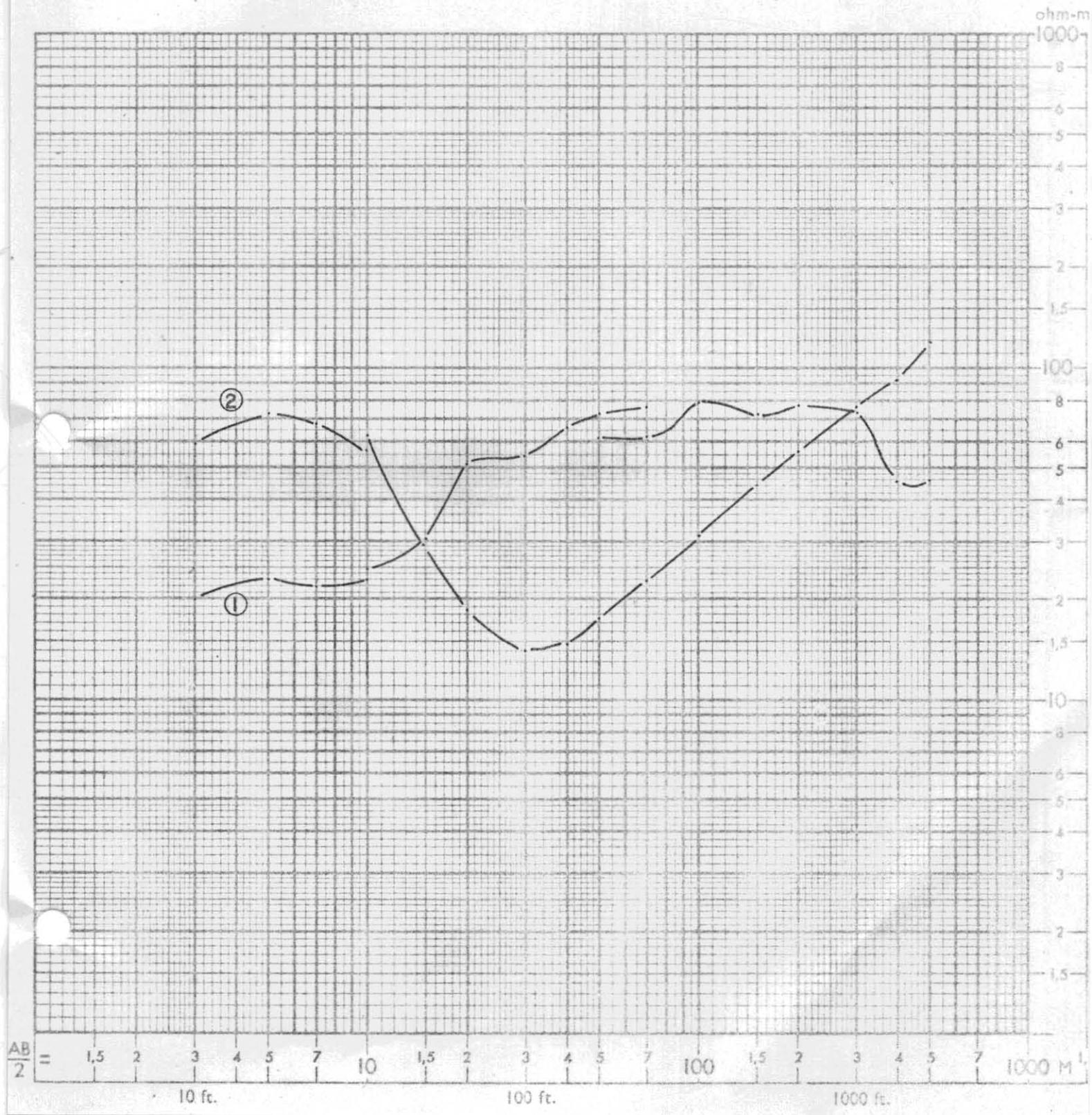


SELINA GRID

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E.S.① 300S of 3900W-152N

E.S.② 96N - 3500W



$\frac{AB}{2} =$ 1.5 2 3 4 5 7 10 1.5 2 3 4 5 7 100 1.5 2 3 4 5 7 1000 M

10 ft. 100 ft. 1000 ft.

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4-3. Apparent Chargeability : (Plates 8 and 9)

The general features of the Apparent Chargeability maps are very similar - a positive Apparent Chargeability axis coincides with the main conductive axis.

The I.P. axis is actually the northern extension of anomaly A1 (detected in 1970) between Lake Rolleston and the profile 88N.

The I.P. anomaly is due to an almost outcropping body about 300 feet wide. The amplitude of the I.P. anomaly increases northward, and intersects the southern limit of the Jukes Breccia body. It then becomes wider and more complex within the Jukes Breccia body. Several secondary I.P. axes appear near the southern limit of the Jukes Breccia. A secondary I.P. axis coincides with the magnetic axis which follows the western limit of the Jukes Breccia body.

Magnetic anomalies associated with strong I.P. anomalies are small; weak I.P. anomalies coincide with the main magnetic anomalies. therefore, the main I.P. anomalies are due to non-magnetic sulfides.

4-4. Self Potential :

A strong and narrow S.P. anomaly (-400 mv) is situated between peg 2250W and peg 2450W on profile 136N. It is probably due to an almost outcropping, narrow, pyrite lode.

There is no significant S.P. anomaly on any other profile.

4-5. Conclusions and Recommendations :

Anomaly A1, detected in 1970, extends at least 5500 ft. north from profile 88N. The percentage of polarizable conductive material is particularly large between profiles 112N and 136N. Between these profiles, a strong I.P. anomaly coincides with a strong conductive anomaly. The anomaly extends further north on the Jukes Breccia body but it becomes wider and more complex near the southern limit of the Jukes Breccia body.

We would recommend the following drillholes :

| <u>Location of Drill Collar</u> | <u>Dip</u> | <u>Length</u> |
|---------------------------------|------------|---------------|
| 136N - 2200W | 50°E | 500 feet |
| 136N - 2500W | 50°E | 200 feet |
| 128N - 2600W | 50°E | 500 feet |
| 120N - 2700W | 50°E | 500 feet |

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5. GOOSENECK

5-1. Apparent Resistivity : (Plates 11 and 12)

Electrical Soundings (see Figure 4) show that the total thickness of the overburden and weathered fringe does not exceed 100 feet. The resistivity of the overburden is about 30 ohm-m - the resistivity of the bedrock exceeds 1,000 ohm-m, except on some anomalies which coincide with shales.

The two Apparent Resistivity maps are similar. There are two main north-northwest conformable conductive axes. The western axis is intersected by a transverse fault between profile 56S and profile 64S. This fault is shown particularly well on Plate 12.

Both anomalies have a larger amplitude on Plate 11 which corresponds to the largest pole-dipole array. Therefore, they are due to conductive bodies rooted in the bedrock.

The eastern axis corresponds to the Red Hills mineralization - the western axis corresponds to a shale stratum.

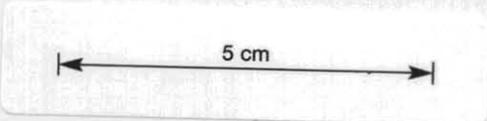
5-2. Apparent Chargeability : (Plates 13 and 14)

Two main anomalous axes coincide with the conductive axes. The eastern axis coincides with anomaly A1 defined in the 1970 report - the western axis coincides with anomaly A5.

A1 is due to disseminated sulfides and magnetite in acid volcanics. On profile 8S, the mineralization is more conductive and may consist of massive veinlets. A1 extends beneath the Owen Conglomerate south from profile 48S.

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Fig.4

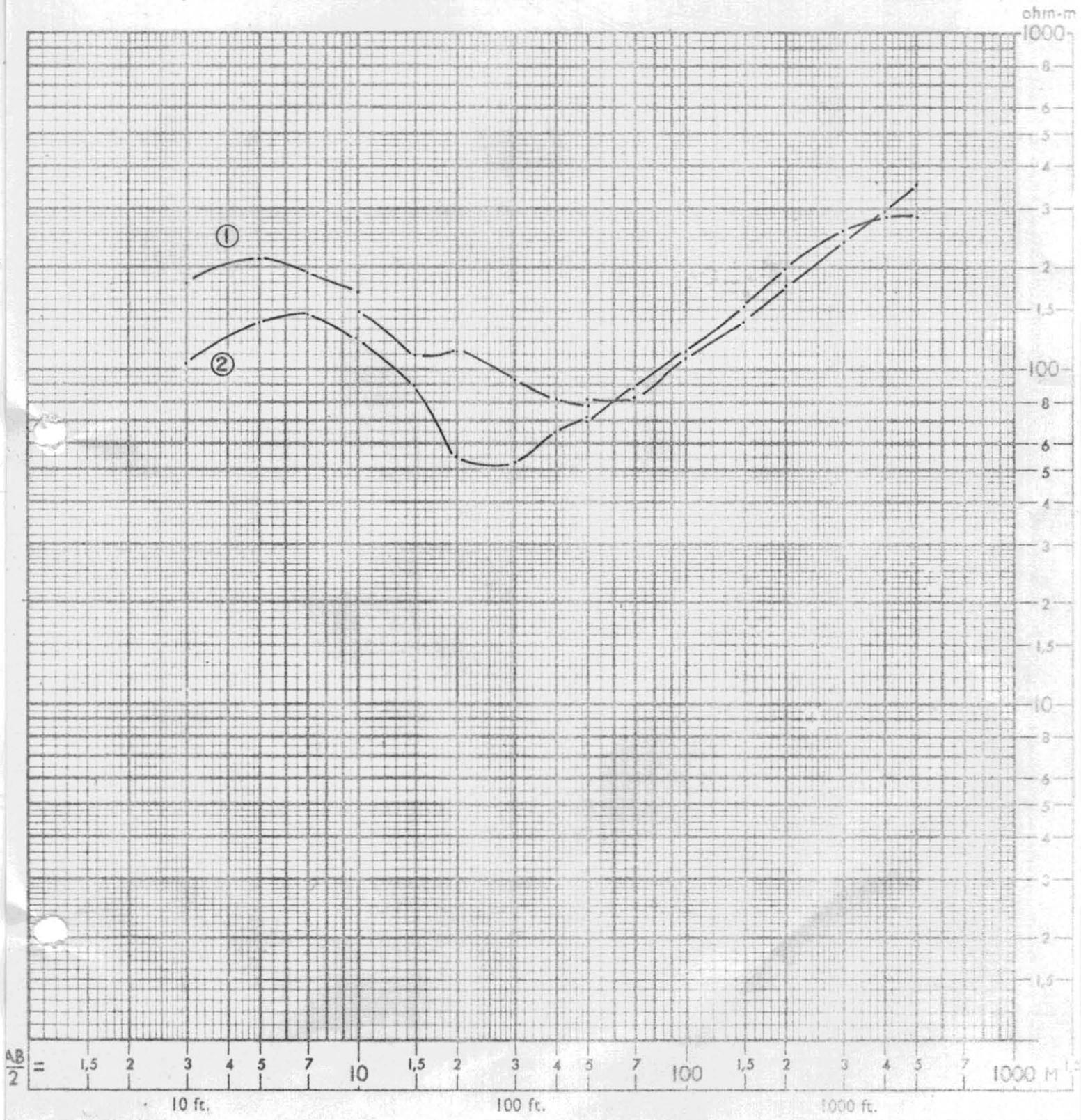


GOOSENECK GRID

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E.S.① 4800S-00

E.S.② 300N of 1100E-72S



A5 is due to black shales.

5-3. Self Potential :

There is no significant S.P. anomaly except on profile 8S where a -50 mv anomaly is associated with A1. This confirms that massive sulfide veinlets occur at A1 on profile 8S.

5-4. Conclusions and Recommendations :

The 1971 measurements showed that anomaly A1 extends at least over 4,500 feet from profile 8S southward. A1 extends further beneath the Owen Conglomerate.

Previous geological observations and percussion drillholes have shown that A1 is due to magnetite, pyrrhotite, pyrite and chalcopyrite disseminated in acid volcanics. The mineralization has an electrical continuity at profile 8S. The electrical continuity is most likely to be due to massive sulfide veinlets.

Anomaly A5 coincides with a black shale stratum.

We would recommend drilling a hole on the southern extension of A1:

- Drill collar at 2500E/32S
- Dip : 50°E
- Length : 700 feet.

The Party Chief

The Chief Geophysicist

G. CLARIDGE

G. OMNES

Brisbane, 10th June, 1971.

GEOPHYSICAL SURVEY AT LAKE ROLLESTON (TAS)

APPARENT RESISTIVITY

GRADIENT ARRAY: AB = 6560' MN = 50'



SCALE 1" = 200'

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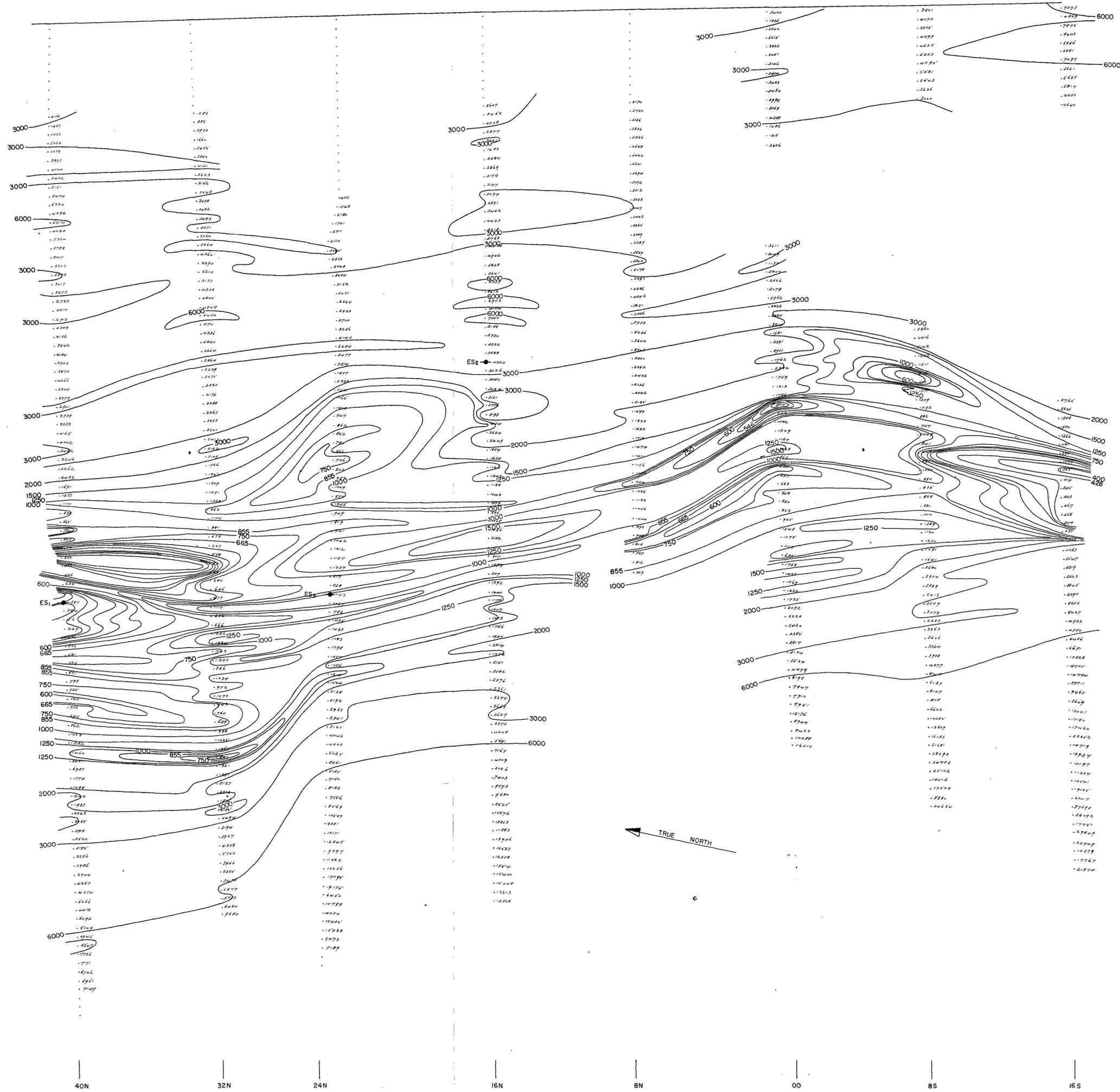
LEGEND

Apparent Resistivity in ohm-m

Apparent Resistivity Scale $\rho_n = \frac{6000}{n}$

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GEOPHYSICAL SURVEY AT LAKE ROLLESTON (TAS)

APPARENT CHARGEABILITY

GRADIENT ARRAY : AB = 6560 MN = 50'

TIME ON : 2 SECONDS TIME OFF : 2 SECONDS
INTEGRATION FROM 450 TO 1150 MS AFTER CUT OFF

SCALE : 1" = 200'



LEGEND

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Apparent Chargeability in milliseconds

C. G. G.

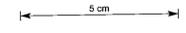
26 - 28 Manning Street, SOUTH BRISBANE



GEOPHYSICAL SURVEY AT LAKE ROLLESTON (TAS)

MAP OF $M_c = M + 5.60 C_{AB}$

GRADIENT ARRAY : AB = 6560 MN = 50'
TIME ON : 2 SECONDS TIME OFF : 2 SECONDS
INTEGRATION FROM 450 TO 1150 MS AFTER CUT OFF
SCALE : 1" = 200'



LEGEND

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$M_c = M + 5.60 C_{AB}$ in milliseconds

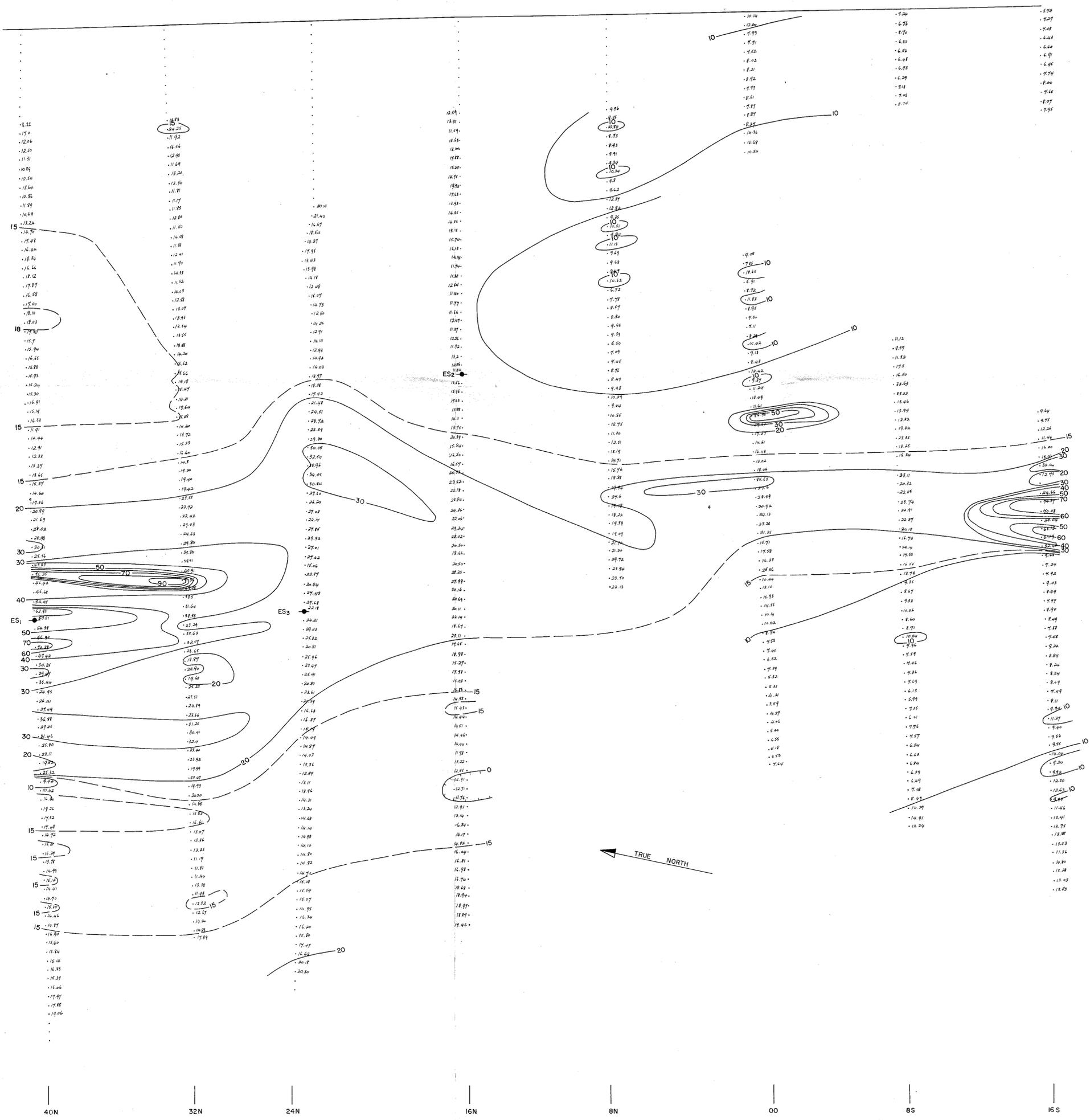
M = Apparent Chargeability

$$C_{AB} = \frac{AB}{2\rho_a}$$

ρ_a = Apparent Resistivity

C. G. G.

26-28 Manning Street, SOUTH BRISBANE



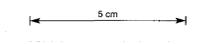
40N 32N 24N 16N 8N 00 8S 16S

GEOPHYSICAL SURVEY AT LAKE ROLLESTON (TAS)

COMPREHENSIVE MAP

SCALE : 1" = 200'

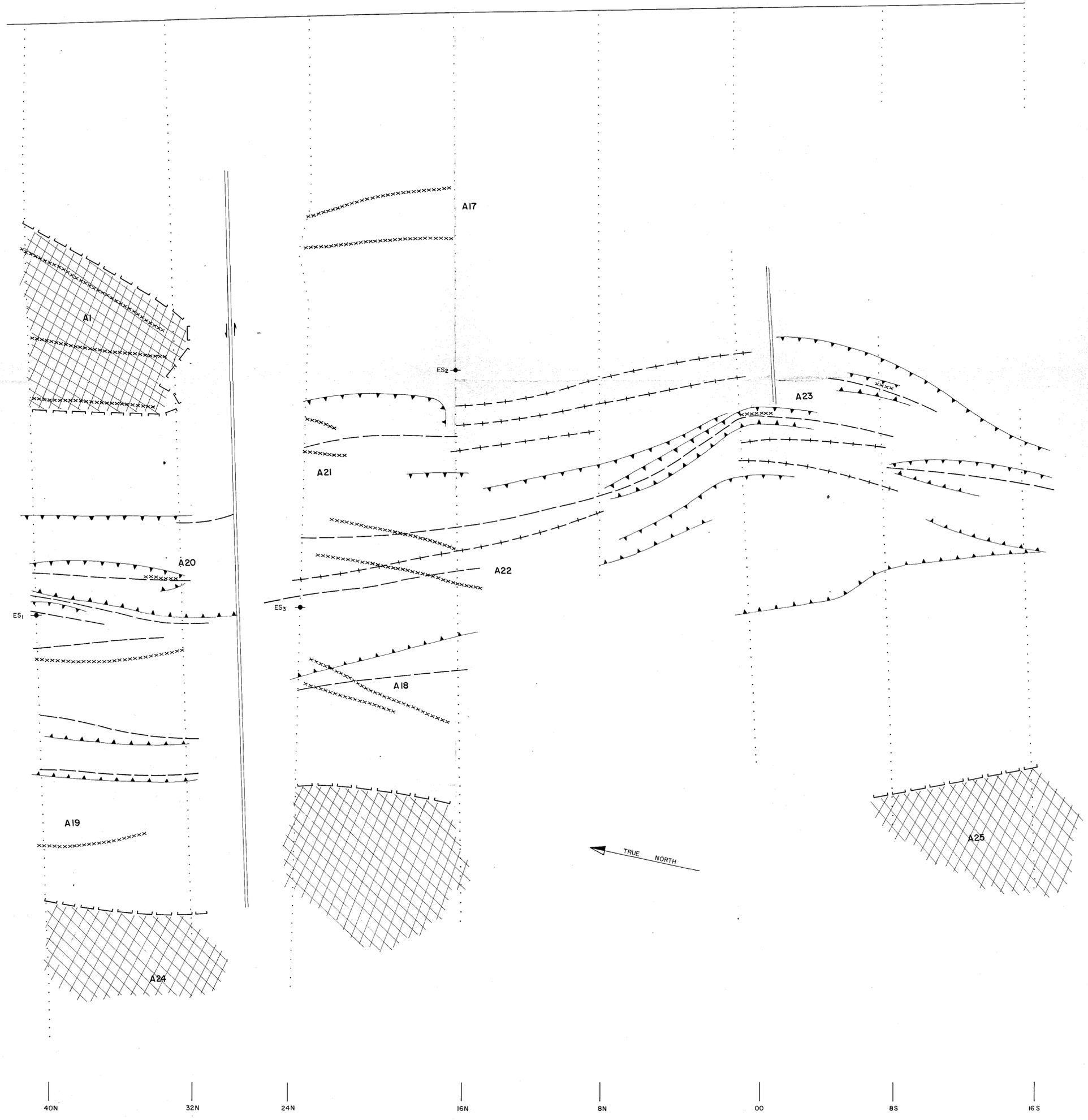
LEGEND



- Conductive axis
- - - Resistant axis
- (higher) / (lower) Gradient of resistivity
- ▨ Conductive area
- xxxxxxx I.P. positive axis
- ==== Electrical discontinuity
- (higher) / (lower) I.P. gradient
- A21 Anomaly number

864032

C. G. G.
26 - 28 Manning Street, SOUTH BRISBANE



GEOPHYSICAL SURVEY AT LAKE SELINA (TAS)

VERTICAL COMPONENT OF THE MAGNETIC FIELD

5 cm

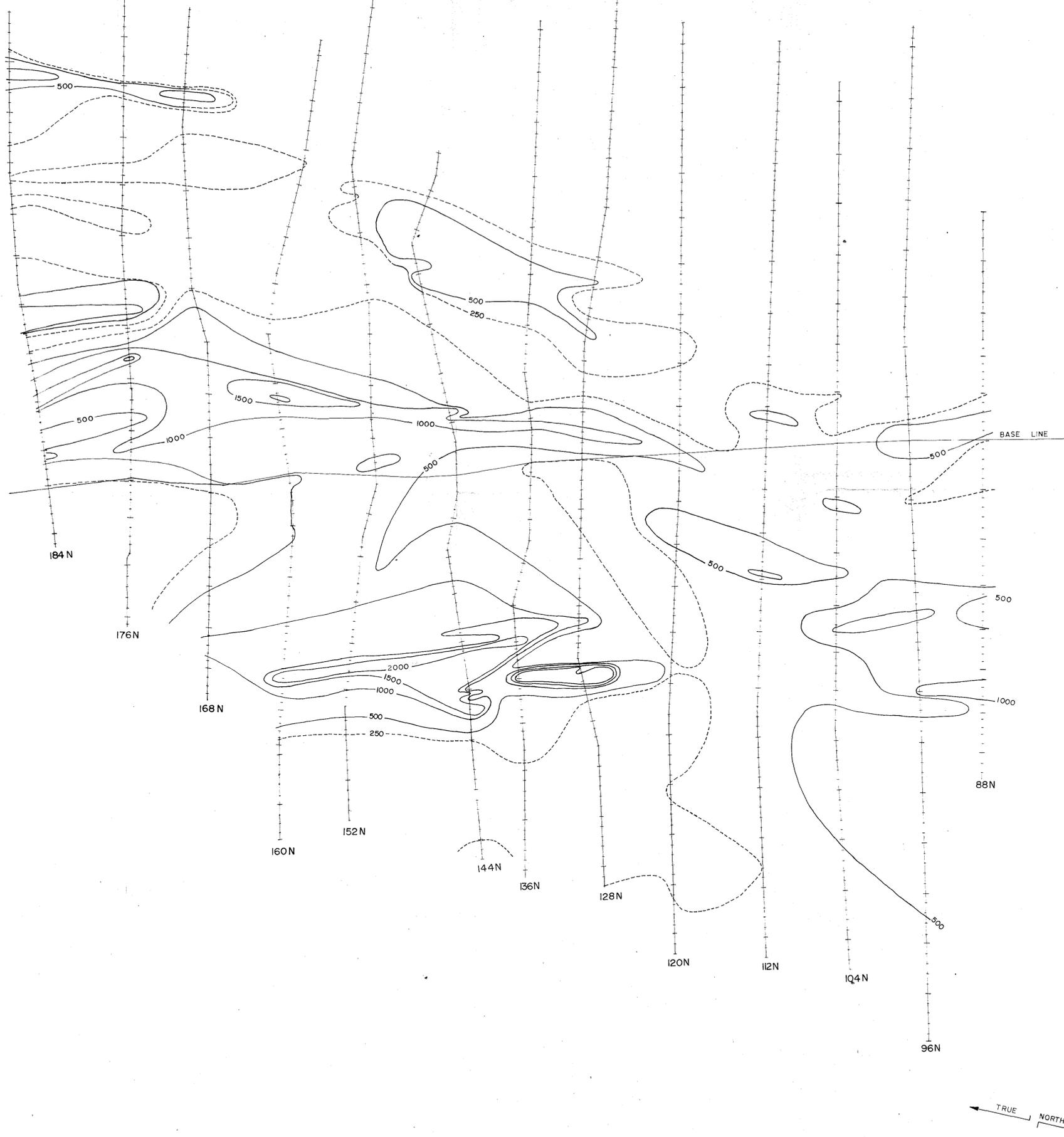
SCALE 1" = 500'
864033

LEGEND

1000
Isogam vertical field in gammas

C.G.G.

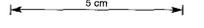
26-28 Manning Street, South BRISBANE



GEOPHYSICAL SURVEY AT LAKE SELINA (TAS)

APPARENT RESISTIVITY RECEIVER 2 POLE DIPOLE

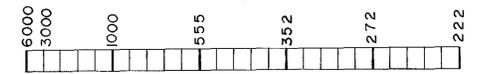
AM = 800' AN = 1000' MN = 200'
 TIME ON : 2 SECONDS - TIME OFF : 2 SECONDS
 INTEGRATION FROM 450 TO 1150 MS. AFTER CUT OFF
 864035
 SCALE : 1" = 500'

LEGEND 

••• Apparent resistivity in ohm.m.

555 555 ohm.m. contour.

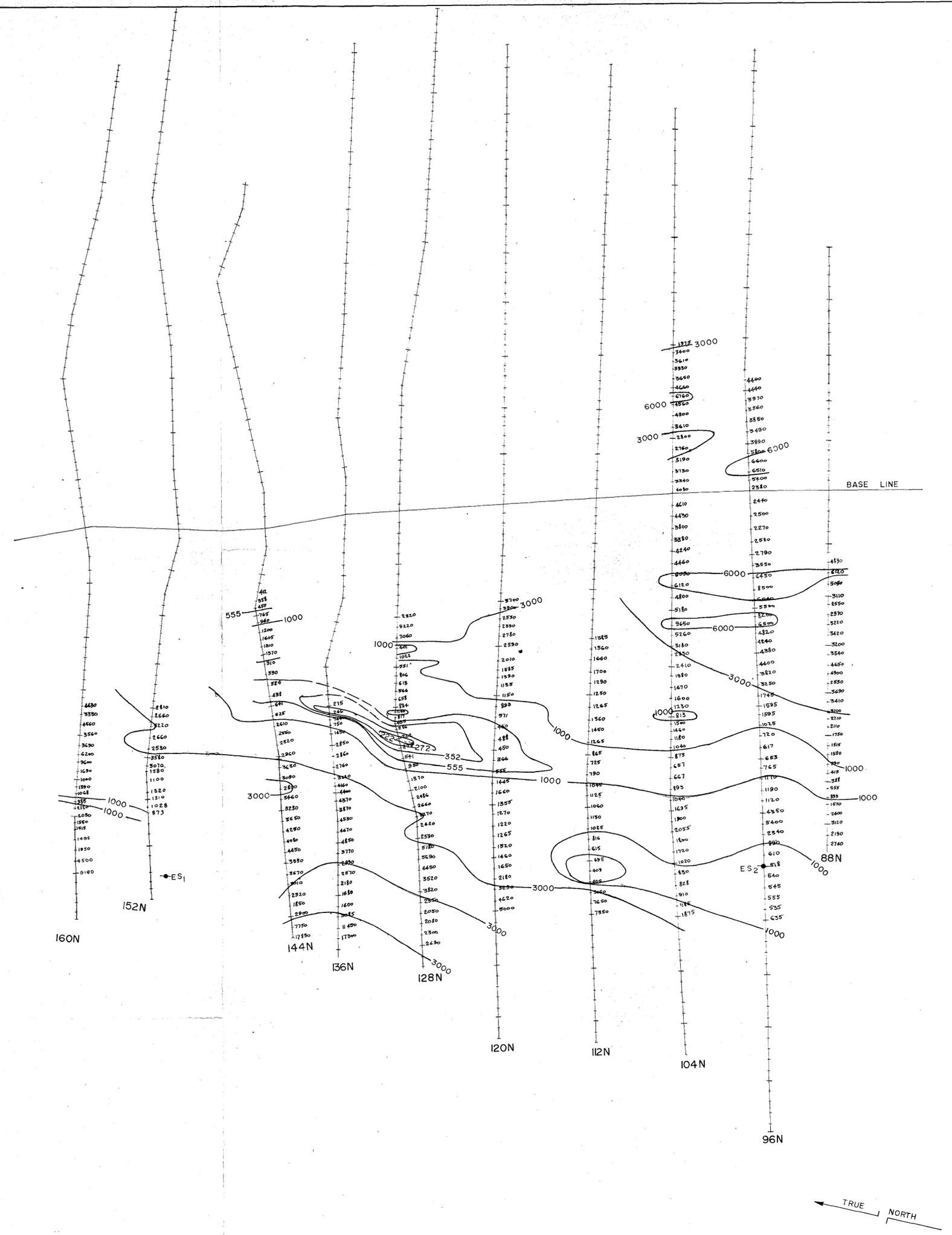
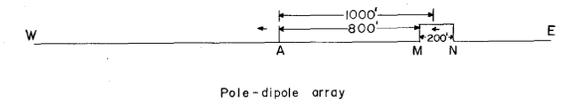
• ES₁ Electrical soundings.



Contour interval in ohm.m.

C.G.G.

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GEOPHYSICAL SURVEY AT LAKE SELINA (TAS)

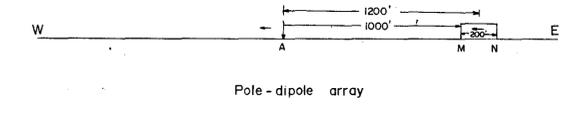
APPARENT CHARGEABILITY RECEIVER I. POLE DIPOLE

AM = 1000 Ft. AN = 1200 Ft.
 TIME ON : 2 SECONDS TIME OFF : 2 SECONDS
 INTEGRATION FROM 450 TO 1150 MS. AFTER CUT OFF

SCALE : 1" = 500
 864036

- LEGEND
- Apparent chargeability in m. sec.
 - 10 m. sec. contours
 - ES1 Electrical Sounding
- 5 cm

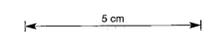
C.G.G.
 26-28 Manning Street, South BRISBANE



GEOPHYSICAL SURVEY AT LAKE SELINA (TAS)

APPARENT CHARGEABILITY RECEIVER 2 POLE DIPOLE

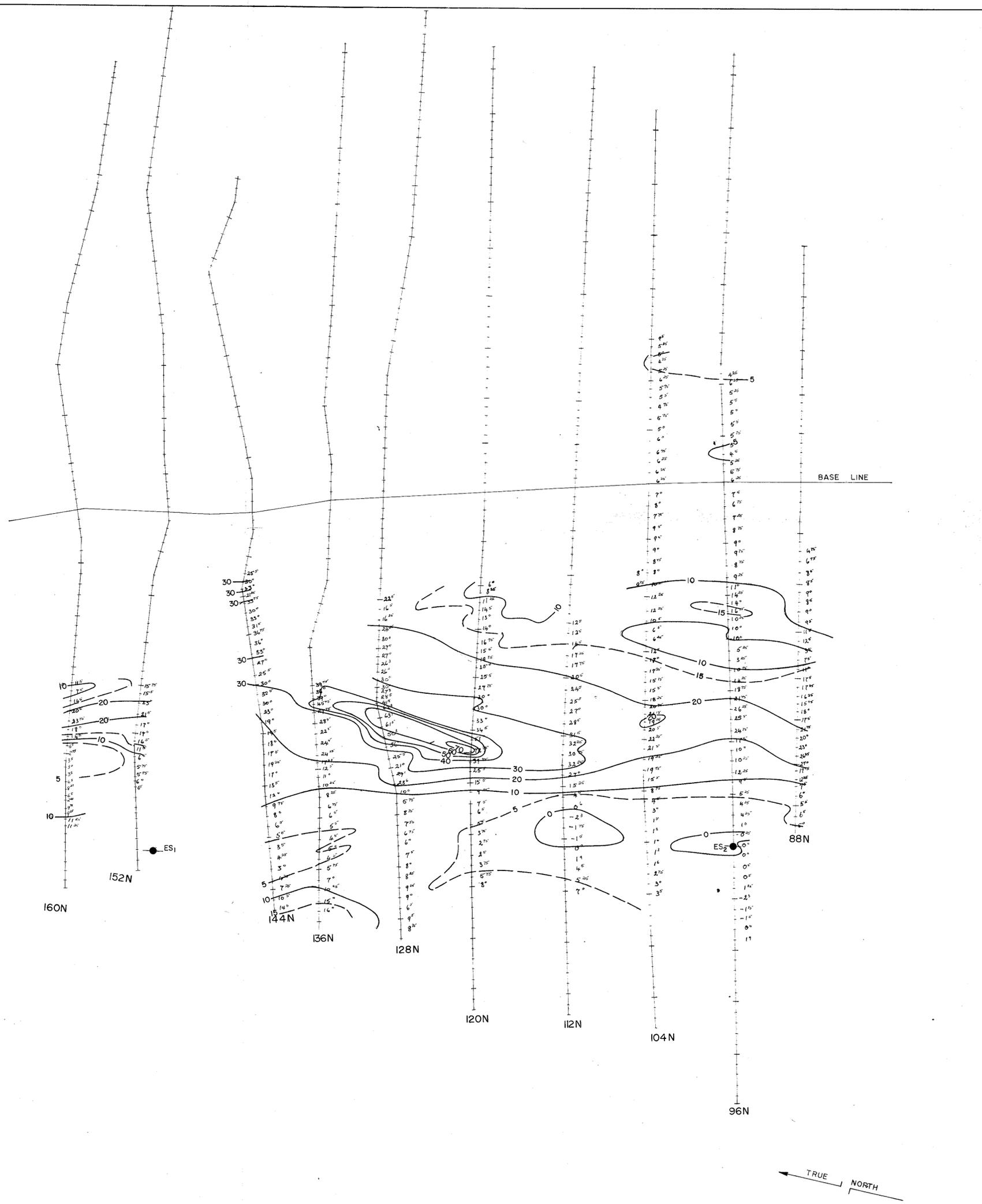
AM = 800 feet AN = 1000 feet MN = 200 feet
 TIME ON : 2 SECONDS TIME OFF : 2 SECONDS
 INTEGRATION FROM 450 TO 1150 MS. AFTER CUT-OFF
 SCALE : 1" = 500'



LEGEND 864037

- 27⁶ Apparent chargeability in m. sec.
- 10 m. sec. contours
- ES Electrical Sounding

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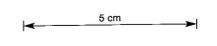


GEOPHYSICAL SURVEY AT LAKE SELINA (TAS)

COMPREHENSIVE MAP

864038
SCALE : 1" = 500'

LEGEND



- Conductive axis
- SPSPSPSP SP axis
- ←→ Positive magnetic axis
- ▨ Conductive area
- xxxxxx I.P. positive axis
- - - Geological limit

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GEOPHYSICAL SURVEY AT MT. TYNDALL (TAS) GOOSENECK AREA

APPARENT RESISTIVITY RECEIVER I

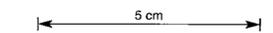
POLE DIPOLE

AM = 400 feet AN = 600 feet

SCALE 1" = 500'

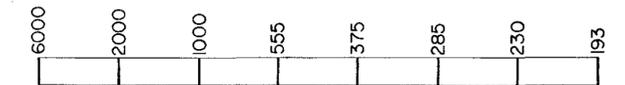
864039

LEGEND



Apparent resistivity in ohm-m.

555 ohm-m contour

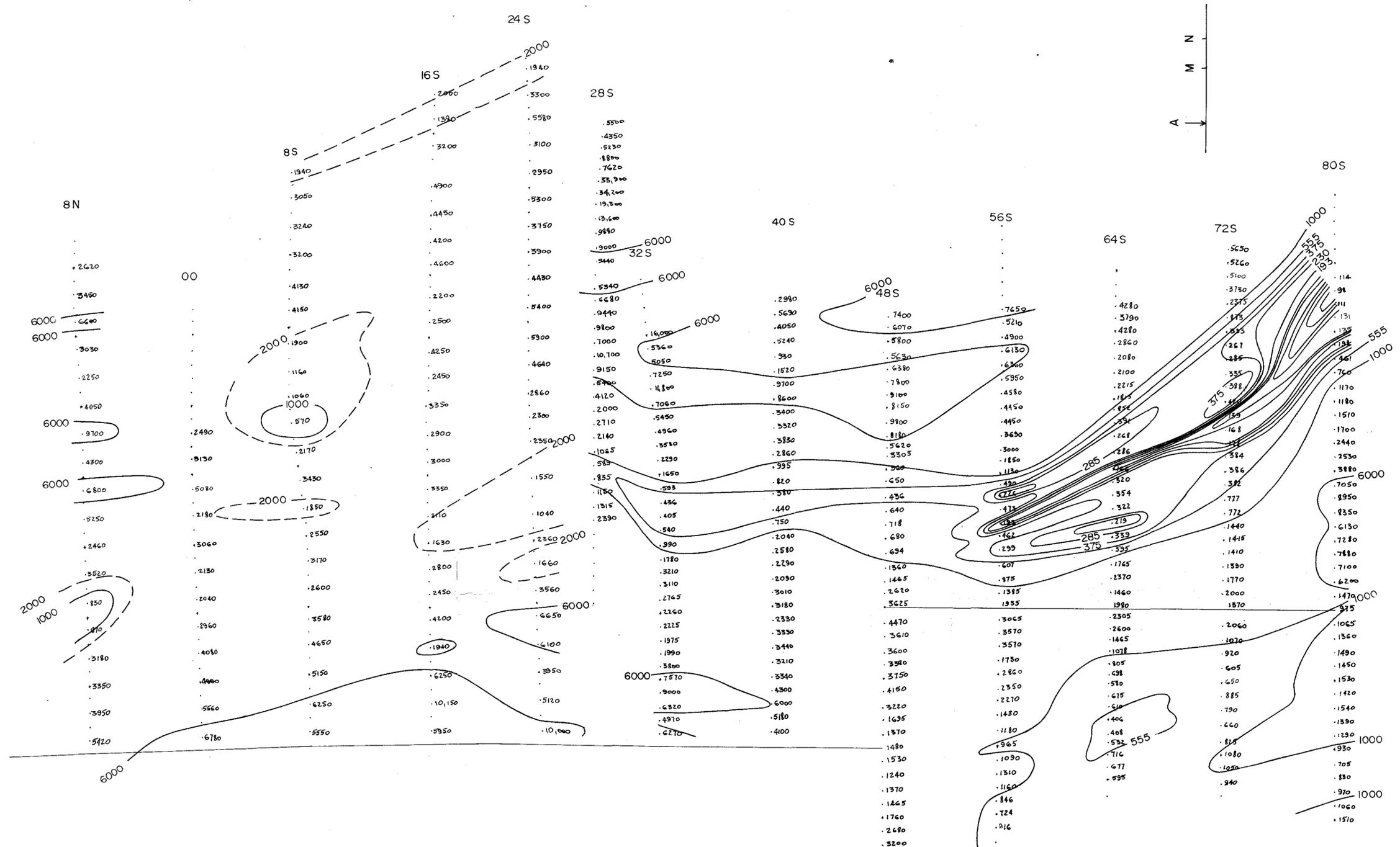
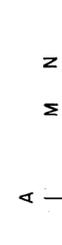


Contour interval in ohm-m.

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Pole-dipole array



GEOPHYSICAL SURVEY AT MT. TYNDALL (TAS) GOOSENECK AREA

APPARENT RESISTIVITY RECEIVER 2

POLE DIPOLE

AM=200feet AN=400feet

SCALE 1" = 500'

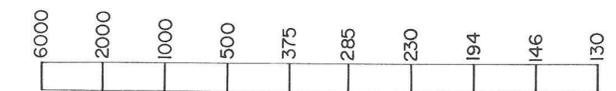
864040

LEGEND

5 cm

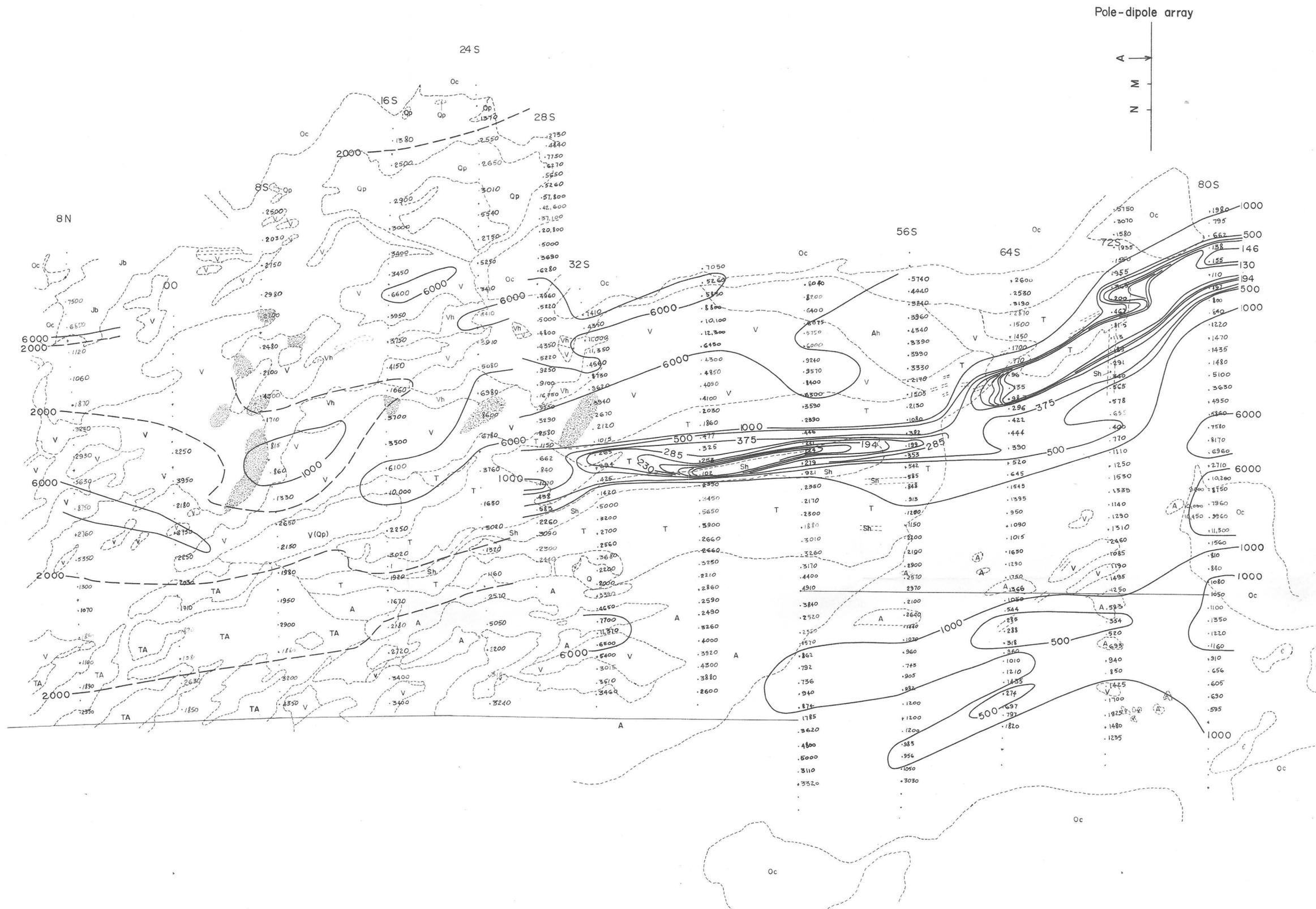
Apparent resistivity in ohm-m

500 ohm-m contour



Contour interval in ohm-m.

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GEOLOGICAL LEGEND

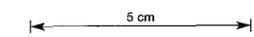
- A — Agglomerate — very chloritic
- Ah — Agglomerate — very haematitic
- Jb — Jukes Breccia
- Oc — Owen Conglomerate
- Q — Quartz Sericite Schist — sheared
- Qp — Quartz Porphyry
- Sh — Black Shales
- T — Tuff
- TA — Tuff and Agglomerate — both sheared
- V — Acid Volcanics — with haematite, chlorite, magnetite with pyrite and chalcopyrite — sheared
- Vh — Haematitic Acid Volcanics
- — Haematite, chlorite, magnetite containing pyrite and chalcopyrite
- Geological Boundary

GEOPHYSICAL SURVEY AT MT. TYNDALL (TAS) GOOSENECK AREA

APPARENT CHARGEABILITY RECEIVER I

POLE DIPOLE
AM=400feet AN=600feet
TIME ON: 2 SECONDS TIME OFF: 2 SECONDS
INTEGRATION FROM 450 TO 1150 MS AFTER CUT-OFF
SCALE 1" = 500'

864041
LEGEND

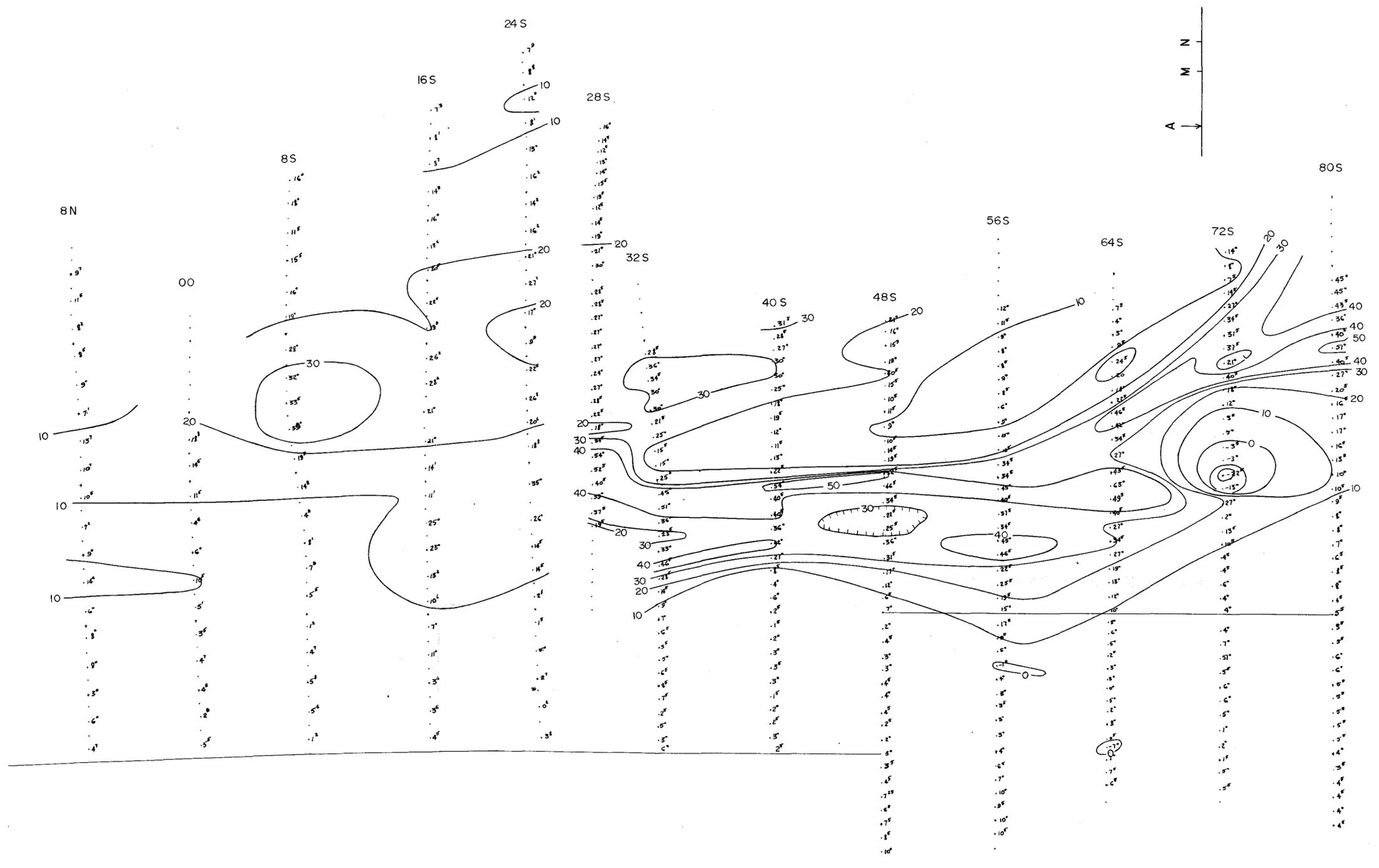
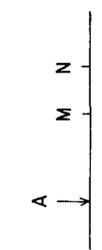


30° Apparent chargeability in m. sec.

10 m. sec. contours

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Pole-dipole array



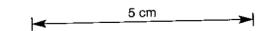
GEOPHYSICAL SURVEY AT MT. TYNDALL (TAS) GOOSENECK AREA

COMPREHENSIVE MAP

SCALE 1" = 500'

864043

LEGEND



- Conductive axis
- SP axis
- Electrical discontinuity
- IP positive axis
- Anomaly number

C.G.G.

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Magnetic North

