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A MINERALOGICAL/GEOCHEMICAL APPRAISAL

OF HEAVY MINERAL CONCENTRATES

E.L. 15/76, DUNDAS, TASMANIA.

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KEYWORDS

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MINERALOGY
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MINERAL

GOLD
NICKEL
CASSITERITE
SULPHIDES
1978
HEAVY

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1. INTRODUCTION

There was no time to do statistical graphical work on elemental distributions, correlation coefficients, etc. so that correlation and comparison of compositions of various heavy concentrates is at the best, crude. Apart from this, without separating the various mineral components and analysing each phase the results would appear to be of doubtful value. In future if we do work of this nature it might be better to hire a Frantz Electromagnetic Separator to separate the ilmenite after having separated the magnetite with a hand magnet. Followed by removal of the limonite for separate analyses. This could be done using an ultrasonic separator.

As it happened, the spectrographic analysis (probably good enough for soil analysis) was not overly successful for determining Bi, Sb, W, Cr and Ti, in particular (Ref. 1) because of the abundance of iron present giving a multiplicity of intense iron lines which interfered with the important lines of some elements as above.

In the circumstances a limited assessment of the chemistry of heavy mineral species from the area is presented.

2. SUMMARY

The area where most 'economic' heavy minerals occur is principally the N.E. sector and it is suspected that these could be partly or mainly derived from mine dumps to the north adjoining the Renison Mines. Associated minerals are those from basic rocks such as augite, hypersthene, actinolite, ilmenite and magnetite which are all present as major quantities. Other minerals present are appreciable tourmaline associated with cassiterite occurrences and 'ultramafic' anthophyllite, steatite, chrysotile and metamorphic actinolite, epidote and garnet reflecting ultramafic rocks and metasediments of the associated watershed. The N.W. sector, directly west of the N.E. sector and separated from it by a large serpentine mass and Dundas Group sediments, is mainly an alluvial plain and similarly has captured some economic heavy minerals (cassiterite, chalcopyrite and sphalerite, etc.) from the Renison Bell workings to the north(?). The gangue mineral association is similar to the N.E. area but greater quantities of metamorphic and basic rock minerals are seen here, though a wider variety of metamorphics occur and the adjoining presence of granitic rocks may be shown by the presence of major zircon and of obvious apatite.

Old mine workings on the Cuni ultramafics are seen in the N.W. sector and copper, lead, silver and bismuth mines on the N.E. sector which may be responsible for traces and minor lead-zinc minerals in alluvial 'heavies'. Traces of gold are present in alluvial heavy mineral suites from both the above areas.

The central-north area just to the south of the N.W. sector is similar with abundant metamorphics (actinolite, tremolite and anthophyllite) but dolomite and a greater abundance of ilmenite and rutile. Augite and hypersthene from fresh basic rocks are less obvious.

The central-east sectors (northerly and southerly) are separated for comparison purposes by an unsampled corridor of relief of Oonah Quartzite. The alluvials appear to occur within or flanking an upfaulted block of Proterozoic, Concert Schist. The northerly sector shows very little of derived economic minerals but an abundance of metamorphic and basic gangue minerals. Basic rock minerals (augite, hypersthene and ilmenite) are more dominant and cassiterite occurs as a minor component. Trace of gold, chromite, bismuthinite, hemimorphite, pyromorphite and vanadinite appear. The southern sector shows a dominance of basic rock minerals (augite, hypersthene, ilmenite and rutile) and a near absence of metamorphics other than minor epidote and traces of anthophyllite and garnet. Traces of bornite and arsenopyrite occur. The southern sector shows few metamorphics other than fine grains in rock fragments. Similarly there are few basic and ultramafic minerals.

In spite of the inadequacies of the spectrographic technique, the chemistry bears out very roughly the mineragraphic results. With high tin, boron, some economic metalliferous sulphides, zircon and high chrome - though with no polished section available it was difficult to differentiate between chromite and magnetite present. Tungsten, present in large quantities does not apparently occur as tungsten minerals, wolframite or scheelite. Tungsten may be within the cassiterite lattice or more likely where heavily oxidised is absorbed on limonite/goethite present. This also goes for arsenic as arsenopyrite which could hardly be missed if it were present. To the south of the N.W. and N.E. sectors there is a dramatic drop in ore-forming elements and associated elements/minerals as per Co, Ti, Y, Zr, B, Cr, V, W, As, Bi, Sb and Zn. To some extent but more sporadically this occurs with In, Ni, P, Cu and Pb. Mn, Ca and Ba show very little variation throughout the analytical results and are presumably associated with constantly occurring gangue material such as limonite:goethite. P may be present in unseen apatite. Copper and lead occur in sporadically occurring sulphides and carbonates, etc. and some

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copper, lead and nickel may also be adsorbed ions on iron/manganese oxides. Also indium is partly within the cassiterite lattice and partly adsorbed on limonite and it shows no relationship with Cu present.

A slight upsurge in antimony and zinc occurs in the southern part of the central-east (south) area within and just south of old Pb, Zn and Ag workings. This continues with, in addition an increase in Zr and B in the south sector of sampling.

To obtain a better perspective of geochemical relationships the results should be subjected to rigorous statistical analysis (Q, R mode cluster analysis) etc. Though, in view of Dr. Beevers (the analyst) remarks, the very large abundance of iron present could have obscured the Bi, Sb, Cr and Ti spectrum in which case statistical analysis of the spectrographic results would be inconclusive for several elements.

From the continued mineralogical and chemical work, it would appear that the only areas justifying more detailed work are the north-east and north-west sectors just south of Renison Mines with the proviso that the heavy minerals recovered are not washed down from the Renison Mine dumps? Also possibly prospective is the southern part of the central-east (~ south) sector (refer map), but only from mineralogy.

3. MINERALOGICAL/GEOCHEMICAL SURVEY

3.1 Mineral Associations

N.E. Sector. Sample weights are variable (0.03 to 4.0 g). There is a great deal of rock material in the heavy mineral concentrates in this sector including amphibolite, mica and chlorite schist fragments. Some embedding by magnetite or other iron oxides has occurred to sink this non-heavy material which would otherwise have been mainly excluded from the heavy concentrates. The principal heavy minerals present in this sector are ilmenite, magnetite, limonitised magnetite and actinolite in places; in two cases augite is prominent and this would have washed in from an area outside the lease as no basic rocks are exposed here. The large amounts of cassiterite seen in some specimens are possibly 'wash' from dumps in the Renison mines just to the north of the lease. Tourmaline is common as a minor component in these concentrates. Other minerals present are traces to major spinel (chromite?), a little topaz, steatite, appreciable garnet (pink variety), hypersthene, traces of dolomite and corundum, traces to minor pyrite, pyrrhotite, chalcopyrite, epidote, anthophyllite and actinolite. The presence of anthophyllite, steatite and chromite suggest a serpentiferous (metamorphosed) source. Hypersthene and augite - a dolerite source. Metamorphic minerals, garnet, anthophyllite, actinolite and epidote (show variable grades of metamorphism of sources) also appear as well as tourmaline and topaz ancillary to tin mineralisation. Traces of fresh sulphide present would not be far from a sulphide source and much of the cassiterite is fresh. A trace of gold was seen in sample 004.

N.W. Sector. Sample weights (0.05 to 30.0 g). Again there is contamination with a whole host of rock fragments having iron oxide attachments. This includes quartz, chert, chlorite schist and actinolite schist. The

latter from surrounding metamorphic rocks. A moderate metamorphic trend is noted here with the presence of much actinolite, tremolite, cummingtonite and appreciable anthophyllite (from metabasics and meta-ultramafics?) - possibly from hornfelsed rocks. This is closer to source of basic intrusives, with a great increase in ilmenite, augite and hypersthene. Sulphides (chalcopyrite, pyrite and marcasite) are locally present and in about similar quantities and distribution to those in the N.E. block but freshness indicates proximity of source in the damp West Tasmanian climate. Less cassiterite, tourmaline and absence of topaz shows appreciable distance from exposed tin (or from the mine dumps?). Interesting, garnet becomes much less obvious which may indicate greater distance from a high grade contact metamorphism.

Central North. Sample weights (0.02 to 12 g). A maintenance of the dominance of medium grade metamorphic minerals - actinolite, tremolite, cummingtonite and anthophyllite is noted, but less dominance by basic rock minerals augite and hypersthene. Ilmenite retains its importance, as per the N.W. sector and leucoxene (pseudomorphs) and rutile have increased. The more durable basic rock heavy minerals have increased and this may indicate a greater distance from a basic rock source? The pneumatolytic minerals cassiterite and tourmaline have receded from prominence and metamorphic garnet is present only as a trace.

Central East (~north). Sample weights (0.05 to 10 g). Very much more ferruginised gangue material than for the previous groups and this has reduced the absolute amounts of heavy minerals present. There is much haematite present as haematite pebbles but not a great deal of chert, some limonitic quartz is seen. Much of the gangue material is amphibolite schist. Actinolite maintains a dominant position, as does augite but less of it and tremolite is a major component.

Rutile is dominant and much of this heavy mineral is from basic intrusives. Minor metamorphics such as garnet and staurolite (show fairly high grade contact metamorphic minerals) are seen and ultramafic (or mafic) hornblende, chrysotile and steatite are present in moderate amounts suggesting proximate ultramafic and basic masses. Magnetite has dropped to minor placement and sulphides pyrrhotite, chalcopyrite and sphalerite are fresh so that sulphide sources are close, and may be traced.

Central East (~south). Sample weights (0.05 to 15.0 g). Gangue mineral present is largely ferruginised; and chert, a little mica, amphibolite schist and a little steatite schist are seen. An increase in ilmenite is noted as is spinel (chromite?) and some olivine. Actinolite remains abundant in isolated specimens as does augite, hypersthene and tremolite but the diluting effect by 'loaded' gangue material has reduced the effectiveness of determining heavy minerals by eye. Sulphides (including arsenopyrite), tourmaline and cassiterite occur spottily and there are obvious signs of the proximity of ultramafics (olivine, hypersthene, antigorite and talc). Metamorphic minerals are relatively scarce except in rock fragments. Sulphides and tin occur spottily. Traces of bismuthinite and vanadinite occur. The presence of minor mineralisation is proximate but weathering of iron minerals (sulphides?) has ferruginised much of the heavy minerals released.

Southern Sector. Limonitised quartzite, quartz pebbles and less amphibolite, mica and chlorite schist make up the main 'loaded' gangue components. Similar to the Central East (~south) sector but actinolite and other metamorphic minerals are not much in evidence, though tremolite occurs spottily. Cassiterite is not seen and tourmaline occurs only in minor quantities. Minerals derived from basic rocks (ilmenite, hypersthene, augite and hornblende) occur sparsely, by number of samples, as do ultramafic steatite, anthophyllite and olivine. Minor

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sulphides are still evident. There are proximate metamorphic rocks (witness the rock fragments present) but as fine grained rocks; remote from metamorphic locus.

Remaining samples 360 to 622 in the sector represent resampling of other areas.

3.2 Mineralogical Chemistry

3.2.1 Antimony

High antimony in these specimens is inexplicable unless present as oxides or basic oxides of antimony in limonitic aggregates where these occur, or it may be present as tetrahedrite which occurs as fine inclusions in sphalerite. Arsenopyrite will accommodate up to 1.5 percent antimony only. It is significant that where Sb is reported (092) to be more than 10,000 ppm, only 0.05 g of sample was available, so the high result may be due to poor mixing during preparation of electrode for arcing.

High (1,000 ppm and more) antimony results appear in many samples from 004 to 064, and below sample 064 they become low (<30-500 ppm) except for 'high spots' at specimen 268 (also high lead), specimens 286, 288 and 290, 323, 334 and 350; and from 544 to 620 inclusive. This includes areas at the N.E. end of the prospect (E.L. 15/76) to the centre and the south-eastern area, mainly located in mine dumps existing on the Dundas group sediments.

There appears to be a direct relationship between antimony and zinc as mentioned above and this may be due to ore association.

3.2.2 Arsenic

Reported as traces or minor amounts in specimens 004, 117, 126, 167, 268, 358, 403, 461 and 620. Since arsenopyrite crystals are rarely seen (good crystal forms and characteristic metallic silver yellow colours were seen on occasions where it was identified), then it must be assumed that the arsenic to a great extent is present in other mineral phases. In the northern samples (early samples in list) down to 164 there is a rough sympathy with antimony and this indicates an association with hydrothermal veins and possibly the presence of tetrahedrite-tennantite series minerals as well as minor arsenopyrite; (proximate mineralisation may be indicated by an out-cropping porphyry). There is also high zinc, lead, tin and tungsten mineralisation in this zone.

In the oxidised sample arsenic may be present in mixed oxides.

3.2.3 Barium

Seen as traces of baryte in samples 089, 095 and 106 and minor in 268, yet reports as 300 ppm or less Ba. In the spectrographic analysis large amounts of Ba (>10,000 ppm) are reported in sample 028 where baryte remains undetected. Unidentified witherite may have been present in some oxidised samples but with mainly 0.1 and less percent present the chances are that barium occurs mainly as fine particles of baryte (or other) associated with limonite aggregates. The element is almost invariably antipathetic with lead(?). This would indicate a breakdown of primary minerals and in many cases adsorption of selected ions on limonitic material.

3.2.4 Bismuth

No sympathy with lead content of samples is seen but several 'high' bismuth results (200 ppm and

more) correspond with 'high' arsenic (e.g. 004, 017, 028, 035, 092, 124, 128 and McCooney) and nearly all samples come from alluvials in the block containing the Montezuma Fault zone, otherwise from two scattered positions to the west, roughly at a similar latitude. These correspond with unidentified particles of sulphide.

3.2.5 Boron

Relatively high (1,000 ppm and higher) to sample 052, intermittently high from 092 to 164 and with low results from 167 to 186, followed by few high results from 189 to 622 with McCooney Prospect having the highest result at > 10,000 ppm. The results correspond only very roughly with tourmaline present, or in some cases high B may correspond with no tourmaline seen in sample. However, high tourmaline present corresponds with the position of old mines on the Dundas Group prospects on Precambrian Concert Schist. From specimen 004 to 052 the presence of tourmaline can be equated with the presence of cassiterite and outcropping quartz porphyries but not so in the southern prospects, as per specimens 253 to 620 area where no intrusive quartz porphyries outcrop. Only a trace of possible alternative boron minerals (axinite) was seen in specimen 221.

3.2.6 Cadmium

All resulted are reported as < 3 ppm.

3.2.7 Calcium

Sympathy with other elements analysed does not occur. Where there is a rise in Y with a corresponding rise in Ca, this may be due to the presence of apatite. Calcium-rich minerals reported and present in the concentrate are extensive, including apatite, pyroxene and amphiboles.

Specimens that contain greatest abundance of Ca (> 10,000 ppm) contain appreciable actinolite and dolomite (specimen 128); tremolite (specimen 145); augite and actinolite (specimen 100).

Multiplicity of variable calcium minerals present would obscure sympathetic trends with other elements.

The calcium is of variable concentration but decreases abruptly after sample 164 down the mineral series and going north to south of the lease. A decrease in augite (but not of hypersthene which was classified as orthopyroxene in southern panned samples) takes place.

3.2.8 Chromium

Generally correlates with high cobalt and only very roughly with high nickel - insufficiently so to suggest that Co and Cr are in any sort of stoichiometric relationship in one mineral. Mineralogy suggests chromium is in the magnetite > limonite > ilmenite lattices, though as a leached product it is adsorbed on to limonite particles. Abundant chromium is pretty much confined to the northern part of the area and is most likely related to run-off from the serpentine masses (in spite of the element's relationship with cobalt).

Samples with Cr results of more than 10,000 ppm are 004, 031, 041, 098, 100, 115, 124, 126, 546 and 620. Many (041, 098, 100, 102, 115) are probably related (taking into consideration drainage directions of streams) to outcropping gabbros and serpentine and to magnetite or an unidentified chromium mineral weathering from these. The remainder are situated on Dundas Group sediments and it is difficult to assign an origin for them.

Chromium is often a contaminant in spectrography, from the grinding procedure of sample. This may account for some high results.

3.2.9 Cobalt

High Co (30 ppm and more) are from alluvium draining from the northerly gabbro > serpentine areas - these include samples 095, 098, 100, 102, 134, 141, 143 and 145. Others of high Co - 004, 047, 053, 115, 122, 124 and 149 are situated among mainly Cambrian, Dundas Group sediments.

Where high Co corresponds with high Cr as in 004, 098, 100, 102, 115, 124 the specimens are divided evenly among alluvium draining from gabbro/serpentine and from sediments. The cobalt would appear to be present mainly in the magnetite lattice or from magnetite altered to limonite - also some cobalt may be in ilmenite, unidentified chromite or spinel reporting with magnetite(?).

3.2.10 Copper

High coppers (1,000 ppm and more) are fairly common and only in sample 017 (2,000 ppm and in 065, 1,000 ppm) was chalcopyrite seen. Other places where copper is reported, the element is very likely associated with limonite, as fine chrysocolla or other copper silicate, or with wad.

3.2.11 Gold

All results are reported as < 3 ppm, yet a trace of gold was seen in specimens 004, 082, 120, 247, 251 and possibly 264 and 266.

3.2.12 Indium

From specimens 004 to 164 tin-rich zones, the relationship between In and Sn is sympathetic. After this the amount of In and Sn falls away to mainly < 5 ppm In and < 2 ppm Sn, except specimens 189 and 191 where a slight increase in values occurs and there is a sympathetic

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trend. Similarly a vague sympathetic In:Sn trend occurs from 566 to 622. Most specimens with a sympathetic Sn:In trend are in the N.E. sector of the lease flanking the northern part of the Montezuma Fault and in Dundas Group sediments south-east of Renison Bell mine. The remainder of sample locations from 566 to 622 are not detectable on the map.

It is interesting to note the antipathetic relationship between zinc and tin showing a virtual absence of distinctive zinc minerals, and for antimony and indium, and lead and indium, though there is a tendency for higher lead and antimony to occur in the high In zone of the N.E. block. Some In contents are higher than the average copper content in this zone. Similarly In is higher than Cr in places. Indium appears to be in the cassiterite lattice or in associated limonitic material in the N.E. block.

3.2.13 Lanthanum

Mostly reported as <100 ppm. Where greater as in samples 132 (300 ppm), 136 (200 ppm), then titanium (>10,000 ppm and 5,060 ppm) is high. Since rutile is present as a major component the La could be in the rutile mineral lattice, yet lanthanum normally sympathetic with yttrium, remains low (30 ppm).

3.2.14 Lead

Some occasional high leads (>10,000 ppm) occur in samples 065, 268 and 302 in the S.W. sector of the N.E. block of the lease, close to the Hecla (copper-bismuth-lead) and Ramsdale (lead-copper) mines and may have washed in as oxidised sulphide products from dumps on the latter mine and judging from the limonitic nature of most of the panned samples, the element could have been present in limonite-stained cerussite associated with limonitic aggregates. Specimen 268 comes from a stream

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west of Mt. Dundas and close to and east of the nearby Sth. Comet (lead, zinc) mine in the Dundas Group of sediments. The last in the list of samples high lead specimen shows only chalcopyrite as its 'economic' mineral component but since many of the minerals present are limonite stained it is possible that the lead is present as unidentifiable cerussite (altered weathered galena), adjoining and downstream from a group of mines (lead-silver), e.g. Comet Mine, Maestries Mine, etc.

Lead generally shows more isolated high result (1,000 ppm and more) in the N.E. block lease areas than elsewhere, shows a rough sympathy with other ore-bearing elements (Cu, Sb, Sn) and is rather insensitive to the chemical trends. In a general way though, lead shows some sympathy with Ba. Presumably all occurrences of lead-zinc are from acid volcanics associated with the Proterozoic Oonah Quartzite and Cambrian Dundas Groups of the area.

3.2.15 Manganese

Mainly present around 100 to 300 ppm and shows occasional highs (1,000 to 10,000 ppm). There are more specimens containing $>10,000$ ppm in the N.E. tin-rich part of the lease than elsewhere. In the majority of cases the manganese is present as wad mixed with limonite as dense aggregates. Where little limonite occurs as in samples 035, 041, 082 and McCoony Prospect, the most obvious common mineral in all these samples is ilmenite which as manganilmenite(?) can accommodate large amounts of Mn. No obvious manganese minerals are seen. Alternatively a red to pink garnet occurs in specimens 035, 041 and 082 and this may well be spessartite and this could explain part of the high Mn present.

3.2.16 Mercury

Most results are reported as <30 ppm. The one exception is specimen 100 (100 ppm). Unless there

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are traces of unidentified tetrahedrite present, it is difficult to explain this trace of mercury present.

3.2.17 Molybdenum

Below limits of 'quantitative' determination in that all results are reported as < 3 ppm.

3.2.18 Nickel

High nickel (50 ppm and above) is correlatable with high chromium except in samples 063, 095, 132, 151, 286 and 290. However, 095, 132, 286 and 290 drain off adjoining serpentine/gabbro bodies and there nickel is supposedly in the magnetite lattice or with limonite.

3.2.19 Niobium

All results are reported as < 20 ppm.

3.2.20 Phosphorous

This is moderately high in most samples from 004 to 064 from whence (in numerical sample progression) it declines to 304 and then rises again from 307 to 622 and McCooney Prospect. High P does not correspond with high Ca or with high La or Y. Where it is highest at > 10,000 ppm no apatite is seen, but traces of minerals occur in samples 124, 130, 132 with minor in 134 and traces in 296 and 463. There is little correspondence of analytical results with the apparent presence of apatite.

3.2.21 Silver

Mainly < 0.1 ppm, but increases appreciably and sporadically all the way through specimens (in numerical sequence) after specimen 124 reaching to 20.5 and 30 ppm at 136 and 141 probably from material eroded from surrounding Dundas Group sediments but interestingly where both lead and copper were low. Similarly for the group of specimens 180 to 191 adjoining the Pb, Ag and Pb-Zn mines in the eastern area, west and

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north-west of Mt. Dundas. Appreciable Ag (maximum 30 ppm) recurs all the way through the assay list to sample 584 occurs within the areas around known Pb-Zn prospects in the eastern part of the lease, delimiting the Proterozoic 'window' which flanks the west side of the Montezuma Fault. Silver is notably lacking from the N.E., Sn rich area, the N.W. area around and west of Cuni, and there are patchy areas in the eastern area where Ag is < 0.1 ppm. High silver (20 or more ppm) is occasionally geochemically associated with high Pb (>10,000 ppm) in sample 302, or high Sb (1,000 or more ppm) in samples 461 and 564. No obvious mineral source occurs except that seen in sample 302 (presence of abundant chalcopyrite and pyrite). With the exception of sample 564 all other specimens contain appreciable limonitic material where Ag possibly resides in an oxidised form. In 302 it would be present preferably in the chalcopyrite lattice. Silver content of silicates is only to 15 ppm in some garnets (Wedepohl) which is recorded as having the highest content of Ag, for a rock forming mineral.

3.2.22 Tin

High indium sympathises roughly with tin (but not with zinc) and is therefore likely mainly in the cassiterite lattice. Boron is also high in the series from 017 to 034. Then further down the series, isolated high tins occur from specimens 080 to 092, 102, 122 to 134, where B is not particularly high. High Pb and Zn correspond very generally with high Sn in the north-east Dundas prospect indicating similar proximity to the base metal mineralisation here.

3.2.23 Titanium

This is high all the way through the series and only where more than 10,000 ppm titanium is present will the element be discussed. Many specimens drain off Dundas Group sediments (004, 017, 024, 026, 028, 031,

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035, 041, 047, 056 and 063) from the northernmost part and south of the area (and 544, 546 and 557) from Silurian and Cambrian sediments. Draining from gabbro are 052, 053, 128, 138, 145, 155 and 164. Most specimens containing titanium may be accounted for from the ilmenite and/or rutile present. Where analysis does not check with titaniferous minerals present (e.g. specimens 024, 031, 034, 061 and 128) the titanium is possibly with abundant limonitic aggregates present. Once again errors in reading the densities of titanium spec. lines may account for this. In one or two places (e.g. sample 143) chromite may have been counted as ilmenite although this would be invalid if the crystal forms were plain to see. Some spinel may have remained undetected since no polished section investigation was attempted.

3.2.24 Tungsten

Generally high (1,000 to > 10,000 ppm) in the samples of the N.W. and N.E. parts of the lease, i.e. the tin area, and falls off to < 50 ppm thereafter except for spot 'highs' (100 to 500 ppm). The high results stop at sample 149 close to the northern boundary of the lease. There is a rough sympathy between tungsten and tin in the samples and also of tungsten and indium. The close association of tin, tungsten and indium suggests that they occur in the same mineral cassiterite, but cassiterite will not contain tungsten in its lattice, so that the tungsten if present is there as wolframite or scheelite (not seen in the panned concentrates) or as possible wolframite inclusions in cassiterite(?), but more likely as inclusions or intergrowths of cassiterite within wolframite which was not identified in open dish. Absence of scheelite was proved using S/W ultraviolet light whence no bluish fluorescence was seen and no characteristically whitish resinous scheelite crystals. Particularly high results (> 10,000 ppm) are noted in sample 118 from Quarternary, Older Alluvium in the north-west part of the lease and adjoining the Cuni area.

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Chromium, tin, antimony and lead vary from 1,000 to 3,000 ppm for this sample. This would appear to be a convenient low-lying drainage area from neighbouring old mine dumps. Tungsten shows more consistent highs (1,000 to 10,000 ppm) in samples from the north-east, high tin area. Sample 034 (> 10,000 ppm) adjoins two small quartz porphyry outcrops.

3.2.25 Vanadium

Any vanadium present will most likely be substituting for Ti^{4+} in magnetite. Titanium and vanadium are roughly sympathetic in a few places where titanium is high and this may be due to the presence of abundant magnetite with minor ilmenite and absent rutile (e.g. specimen 128). Generally speaking, there is no sympathy between Ti and V. More often than otherwise high magnetite corresponds with high V (samples 004, 034, 038, 041, 100, 102 and 115 etc.). There are also the exceptions (samples 035, 124 and 126) where V is high and magnetite is present in minor quantities. Only in a few specimens (100, 102, 115, 132, 134 and 143) do high vanadium (300 ppm and more) correspond with proximity to basic/ultrabasic rocks.

3.2.26 Yttrium

Progressing through analyses, the content is quite variable until sample 202 and varies from < 10 ppm to 500 ppm and to some extent is correlatable with high tin and boron. Yttrium cannot be correlated with any other single element analysed in this exercise, though the element is commonly present in minerals where Ca occurs and with F if present.

3.2.27 Zinc

Only traces of minor amounts of sphalerite were detected in samples 063, 089, 247, 262, 284, 302, 403 and 561 scattered over the eastern part of the area mainly to the immediate north-west of Mt. Dundas, partly in Concert schists.

3.2.28 Zirconium

Present in zircon where it occurs, yet where zircon reports at 5,000 ppm in samples 004, 017 and 3,000 ppm in sample 026, then no zircon was seen in the sample. The mineral was missed in other samples and was probably finely embedded in chlorite schist or other heavy rock fragments present. In one or two places where the mineral was present in major amounts, zirconium was reported in small amounts (300 ppm).

3.3 Chemistry - As A Comparison by Catchment Sectors

The north-east sector is characterised by high W, Sn, Ba, As, Ti, Zn, P, B, In and Sb and some sporadically high Cr. This would indicate a high pneumatolytic/hydrothermal zone (Sn, B; W, As, Sb, In, P) and proximity of ultrabasic (Cr), basic rocks (Ti), and granitic material (Zr). Only sporadic high Pb, Zn, Cu and Ba shows the lower but still high proximity of economic sulphides, but no sedimentary Pb, Zn, Cu sulphide deposits. Indium does not sympathise with B, Zn and Pb or the presence of tourmaline and is possibly distributed among cassiterite and different Fe³⁺ rich mineral phases. High yttrium does not sympathise with insensitive (to chemical methods) La. It could be associated with Ca in apatite but apparently is not (cf. Y with P content). It is however, roughly coincident with Ti and Zr and may be in the lattice of ilmenite or zircon, substituting for Fe³⁺ of Fe⁴⁺. Boron and tin both decrease dramatically below 053 (numerical sequence) in the southerly part of the sector possibly indicating a modification of catchment source.

North-west Sector. It is suggested that cobalt here is in the magnetite/spinel lattice. It sympathises roughly with chromium and is sporadically high (compared with the north-east sector) compared with other sectors of the area. As for the north-east sector - In, Pb, Sb, Sn, Zn, Cr, Ti, W, Bi and As, all remain relatively high, indicating a mineralised catchment of such an area. There is however, a slight drop in W, As, In, Pb and a dramatic fall in Y but little difference for P. If the W and In are in the cassiterite lattice (as suggested for the north-east sector) there is a variation in composition of cassiterite for the N.W. sector as the relationship of

W, In and Sn is not maintained except sporadically. An increase in distance from sulphide shedding areas (or change in drainage pattern) and also a phasing out of cassiterite and tourmaline may indicate a greater distance from pneumatolytic tin-producing area to the north of Zeehan. There is a continued proximity to eroding ultramafic and basic rock sources (with maintenance of Cr, and Ti, V 'high').

Central-north Sector. A number of scattered samples which were taken from an area just south of but falling between the N.E. and N.W. sectors with some run-off from serpentine and Dundas sediments. All elements as indicated show a median decrease compared with more northerly areas - In, Pb, Sn, Sb, Zn, As, Bi, Cr, Ni, Ti and V. This indicates an increase in distance from all mineralisation areas of sulphides and tin, and grades into the southerly areas of absence of the ore forming elements. P, Y, W and B show few scattered high results among a background of low results. This may indicate a decrease in apatite (P, Y) and remoteness from pneumatolysis.

Central-east (north). Abrupt decreases occur in In, Sb, Sn, As, Bi, Zn, W, As, Zr, P, B and Cr, shows catchment from less mineralised areas including ultramafic minerals, sulphides and tin except for sporadically 'high' As, In, Pb, Zr, P and B; also Mn shows sources of sporadic mineralisation or of waddy/limonitic absorptive material which adsorbs the metal ions in solutions from weathering sulphide zones. Depending on whether Zr is from granite direct or is secondary from sediments there would still indicate a greater remoteness from a granite source as are shown by 'lower' P and B. Barium (from sediments?) continues with a number of sporadically high results in a background of low results but this is not linked with Ca content.

Central-east (south). Continuing the pattern as per central-east (north) but sporadic increases in Pb, Sb, Zn and W. These are associated with limonite from isolated pockets of eroding mineralisation. The slight upwards increase

in Ni, Bi and As background may indicate the presence of significant pyrite (confirmed by mineragraphy). According to the mineragraphy there is appreciable tourmaline in this area and chemistry shows a very slight increase in Sn background.

Southern Sector. Compared with the central-east sectors, there is a sharp but spasmodic increase in Zr, and as granite is not seen to occur in the area this may be due to catchment from Silurian sediments - but still the zircon appears very fresh(?). Titanium, as an increase in ilmenite and/or rutile has also increased slightly (confirmed in mineragraphy). Increase in P and B are not obvious since there is no greater apatite and tourmaline content. No obvious immediately proximate serpentiferous mass could account for the increase in Cr in specimens 323 and 350 although the doleritic Mt. Dundas (source of increased Ti) is on the eastern relief side of drainage. Nickel is very low here and indicates absence of ultramafics and/or virtual absence of pyrite.

Appreciable chromium present in samples 323, 334 and 350 is from limonite, spinel or magnetite in the specimens and originates from not too distant serpentine masses, though nickel in these specimens is minimal, titanium is high indicating proximity to basic rocks (from Mt. Dundas, dolerite), as is zircon (from micaceous sediments). High boron of some results does not equate with minimal observed tourmaline present.

Without statistical analysis it would be difficult to follow any definite trends between central-east, north and southern sectors and possibly these sectors should not have been included in separate sectors for appraisal of chemistry.

Resampling. From 361 to 622 are rarely marked on the map of sample plots and many are from a later resampling programme of already sampled areas.

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APPENDIX I

027

223028

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===== 272-5733
=====

8th March, 1978.

Mr. P. Macnamara,
P.O.Box 196,
ZEEHAN.
TASMANIA.

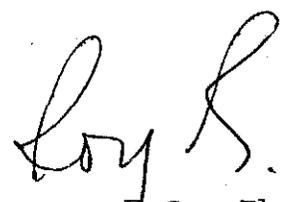
Dear Phil,

At long last you have got the results for batch A2218 in your possession. The samples in this batch were quite the filthiest we have ever had to deal with by emission spectrography, mainly due to the very intense iron lines in the spectrum. With many of the elements, in particular Bi, Sb, W, Cr, Ti, the results are nothing more than an educated guess since very intense iron lines are either sitting on top of, or are closely adjacent to the analyte lines.

Certainly the values for these elements would not be up to the normal precision for Emission Spectrography, and if you really want more precise values we should choose a different technique -- perhaps XRF for some or AAS where suitable.

If ever you are over in Adelaide I would love to show you the problems your samples created. Anyhow, I do hope you can get some value out of the results. As you will see, there is nothing left of some samples, and very little of others, but we do have quite a lot of samples with sufficient sample left for more work if required. We have all the samples packed ready for shipment to Peter in Sydney but will await your further instructions before sending them.

Kindest regards,

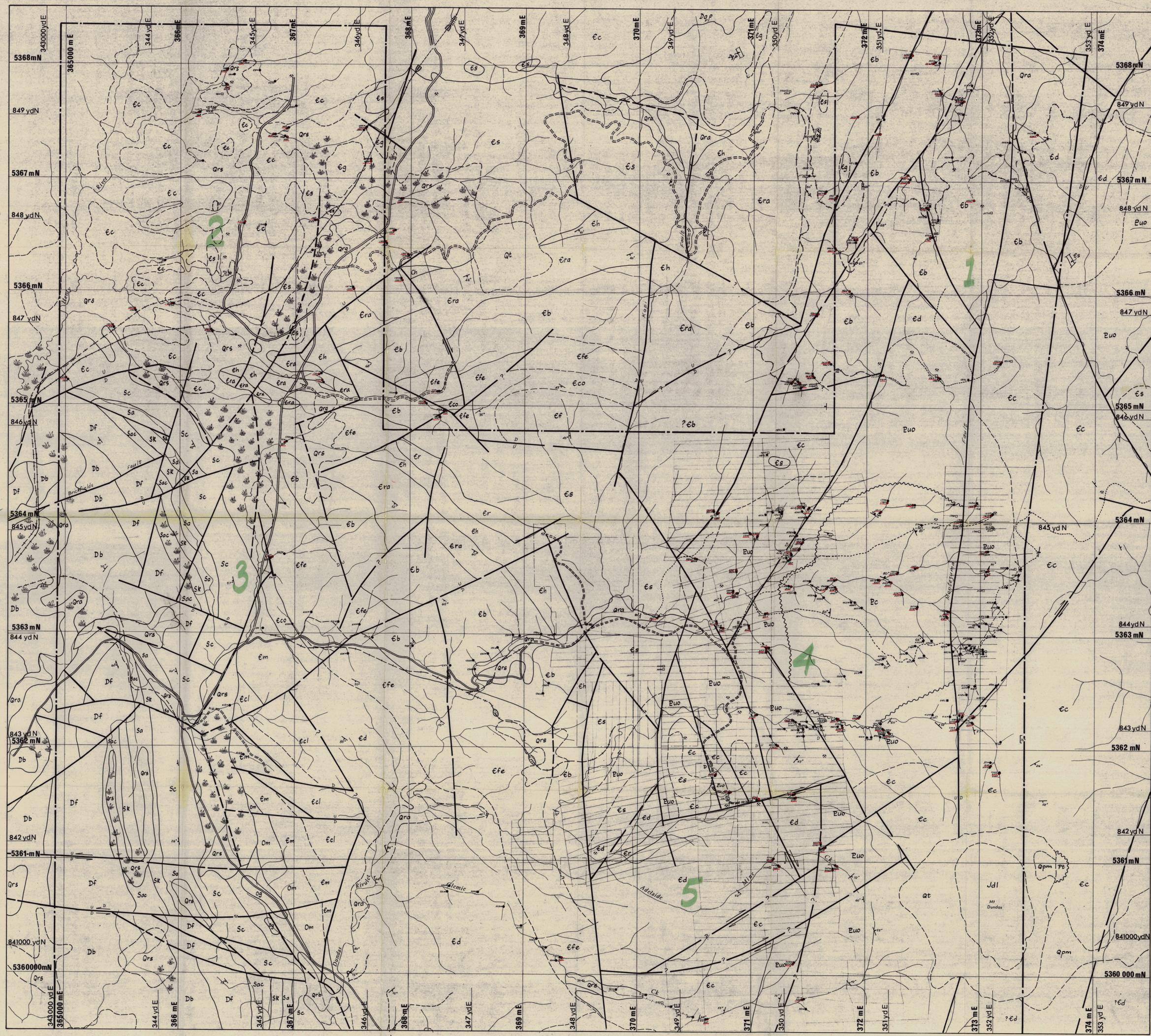


J.R. Beevers, B.Sc., Ph.D.
Managing Director.

cc. Mr. P. Curtis

APPENDIX II

SEE FICHE 2



REFERENCE

Qra	Alluvium
Qrg	Gravels
Qc	Conglomerate talus
Qd	Dolerite talus
Qrs	Older alluvium marsh deposits, downwash etc.
Qpf	Fluvioglacial & lacustrine deposits
Qpm	Moraine

DEVDIAN PERMAN QUATERNARY

P	Pt	Zeehan Glacial Formation
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DEVDIAN PERMAN

D	Db	Bell shale
D	Df	Florence quartzite

SILURIAN

S	Sac	Austral Creek siltstone
S	Sk	Keel Quartzite
S	Sa	Amber slate
S	Sc	Crofty quartzite
S	S	Unassigned

ORDOVICIAN

O	Og	Gordon limestone
O	Om	Moira sandstone
O	Oz	Mt Zeehan conglomerate

CAMBRIAN

E	Em	Misery conglomerate
E	Ecl	Climo Formation
E	Ef	Ferriflow Formation
E	Eco	Comet Formation
E	Efe	Ferrifields Formation
E	Eb	Brewery Junction Formation
E	Era	Razorback conglomerate
E	Er	Hodge slate
E	Eh	Red Lead conglomerate
E	Ed	Dundas Group unassigned
E	Ec	Ormside Creek Formation
E	E	Cambrian unassigned

PROTEROZOIC

P	Euo	Danah quartzite & slate
P	Ec	Concord schist
P	Es	Wynle schist

IGNEOUS ROCKS

Jdl	JURASSIC Dolerite
-----	-------------------

CAMBRIAN DEVDIAN

D	Dz	Quartz porphyry
E	Eg	Gabbro, norite and dolerite
E	Es	Serpentinite and pyroxenite

- Stream sediment (M)
 - Panned concentrate (P)
 - Rock chip, float (R, F)
 - Limonite (L)
- Underlined numbers correspond to analysis sheet

Dundas 1:25,000 223038

PACMINEX PTY. LIMITED

78-1305

DRAINAGE SAMPLE NUMBERS
 PETROGRAPHY OF HEAVY MINERALS
 IN PANNED CONCENTRATES
 E. I. 15/76 - North Sheet
 DUNDAS, WEST TASMANIA

SCALE 1:10,000
 DRAWN P.M.M./P.H.
 DATE
 REVISED

K 555-6

029

223031

APPENDIX III

SEE FICHE 2

APPENDIX IV

TABLE 2

MINERAL ABUNDANCE BY AREAL SECTOR, DUNDAS

(a) N.E. SECTOR (specimens 004-089)

<u>Major</u>	<u>Minor</u>	<u>Trace</u>
Actinolite	Anthophyllite	Apatite
Augite	Chalcopyrite	Cerussite
Cassiterite	Chrysotile	Gold
Garnet	Epidote	
Hypersthene	Pyrite	
Ilmenite	Rutile	
(+ leucoxene)	Steatite	
Magnetite	Zircon	
Spinel	Sphalerite	
Tourmaline		

(b) N.W. SECTOR (specimens 090-145)

<u>Major</u>	<u>Minor</u>	<u>Trace</u>
Anthophyllite	Apatite	Antigorite
Actinolite (great increase over NE block)	Chalcopyrite	Baryte
Augite (great increase over NE block)	Cassiterite	Gold
Hypersthene (more pervasive than NE block)	Chrysotile	
Ilmenite (+ leucoxene) (much heavier than NE block)	Cordierite	
Magnetite	Corundum	
Siderite	Cummingtonite	
Tremolite	Garnet	
Zircon	Marcasite	
	Rutile	
	Tourmaline	
	Tremolite	
	Pyrite	
	Pyromorphite	
	Sphalerite	

(c) CENTRAL NORTH SECTOR (specimens 147-164)

<u>Major</u>	<u>Minor</u>	<u>Trace</u>
Actinolite	Apatite	Antigorite
Anthophyllite	Augite	Chalcopyrite
Dolomite	Cassiterite	Corundum
Ilmenite (+ leucoxene, more so than for NW)	Pyrite	Epidote
Magnetite	Sphene	Tourmaline
Rutile (increases over NE and NW)	Spinel	Hypersthene
Tremolite	Zircon	Garnet
	Cummingtonite	

(d) CENTRAL (northerly) EAST SECTOR (Specimens 167 to 223;
251 to 257 and 284 to 304).

<u>Major</u>	<u>Minor</u>	<u>Trace</u>
Actinolite	Augite	Apatite
Antigorite (1 spl only)	Chrysotile	Biotite
Chalcopyrite (1 spl)	Epidote	
Dolomite	Garnet	
Ilmenite (but less for NE and NW blocks)	Hornblende (1 spl only)	
Leucoxene	Marcasite	
Magnesite	Pyrrhotite	
Orthopyroxene/hypersthene	Sidierite	
Pyrite	Spahlerite	
Tremolite	Sphene	
Rutile	Steatite	
	Tourmaline	
	Zircon	

033

(e) CENTRAL (southerly) EAST SECTOR (specimens 225 to 249; 260 to 281 and 307 to 310).

<u>Major</u>	<u>Minor</u>	<u>Trace</u>
Actinolite (1 spl only)	Antigorite	Apatite
Augite	Arsenopyrite	Bismuthinite
Dolomite	Baryte	Chrome mica
Epidote	Biotite	Cordierite
Hypersthene	Cassiterite	Gold (?)
Ilmenite (increase in abundance) + leucoxene	Chalcopyrite	Hemimorphite
Limonite	Corundum	Pyromorphite
Magnetite	Cumingtonite	Siderite
Olivine	Garnet	Sphene
Orthopyroxene	Monazite	Vanadinite
Pyrite	Pyrrhotite	
Pyrolusite	Rutile	
Specularite	Sphalerite	
Spinel (chromite?)	Steatite	
	Tourmaline	
	Zircon	

(f) SOUTHERN SECTOR (specimens 315 to 360)

<u>Major</u>	<u>Minor</u>	<u>Trace</u>
Augite	Anatase	Anthophyllite
Hypersthene	Epidote	Arsenopyrite
Ilmenite	Pyrite	Bornite
Magnetite	Sphene	Garnet
Olivine	Steatite	Hornblende
Rutile	Tourmaline	Pyrrhotite
Spinel(?)	Zircon	
Zoisite (1 spl only)		

N.B. Major (5 volume % or more)
 Minor (1 to 5 volume %)
 Trace (<1 volume %)

APPENDIX V

SEE FICHE 2

