the exposure is fresh, the rock is seen to be comprised of hypersthene, augite, plagioclase and titano magnetite, with replacement by uralite, chlorite, saussurite and epidote. In Costean 3600N, the rock is deeply weathered and has a pseudoclastic texture similar to that in the intrusive rocks of the Basaltic Complex. It appears to be related to the Serpentine Hill Complex and was probably tectonically emplaced at the same time. Other similar gabbroic bodies are thought to occur in the area but are not exposed.

A stock of altered greisenised adamellite occurs at Pine Hill, south-west of GAP (plan TAS/2/1421). This body, together with porphyry dykes extending radially from it, forms the top of a buried granitic body which is known to underly the Renison Bell Anticline (Appendix III). The country rocks, including ultramafics and gabbros, have been metamorphosed up to 1300m from Pine Hill.

The dyke like body of altered microgranodiorite, mapped in Wilbur Creek along a probable fault zone, is thought to be related to the Pine Hill body. The rock is silicified and made up of almost equal proportions of quartz, plagioclase and chloritised mafic minerals.

5.1.4. Structure

Folding

A structural interpretation of GAP must be regarded as highly tentative on account of the problems of correlation of lithologies and poor exposure. Good exposures in Costeans 1000N and 1200N, on sections of the Ring River and on the Ring River Road, allow for an appreciation of the complexity of folding. The meso and micro fold structures have been measured where possible, and these, together with cleavage bedding relationships and facing directions, have led to a definition of the macro fold pattern.

The dominant fold structures are the Renison Bell Anticline to the west, and the Huskisson Syncline to the east. These structures plunge to the north-north-west in the northern part of the

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