

2. Gradient Array I.P.

The grid was covered by three blocks of gradient array I.P. as detailed in Table 8.

TABLE 8

Spicer Grid Gradient Array I.P.

Array Number	Lines	Electrode co-ordinates (A.M.G.)	
		W	E
1	152S-184S	5353815N 386540E	5354310N 389160E
2	184S-216S	5353130N 387035E	5353500N 389200E
3	216S-248S	5352300N 387455E	5352690N 389420E

Lines 184S and 216S were repeated to check the overlap between arrays.

Contour plans of the chargeability and resistivity data are given in Figures 52 and 53.

Three zones of moderate-order chargeability anomalies were detected. A 53 mV/V peak at 224S, 2890E, was due to Lower Cambrian black shales. A 30+ mV/V zone along the western margin of the grid overlies Owen Conglomerate and is probably caused by a unit containing a high proportion of platy hematite.

The most important I.P. anomaly zone occurs between 152S, 2000E and 200S, 2420E, with a peak value of 39.5 mV/V against a background of around 14 mV/V. This zone corresponds with altered volcanics containing minor sulphide mineralisation, as described above (see Figure 45). The zone is widest on line 168S (about 150 m).

The only resistivity anomaly worth noting occurs near the western end of line 152S over swampy glacial cover. This zone is an extension of the R.T.A.E. Turam anomaly (Hutton, 1980, pl4) and is probably due to the swamp or a clay layer in the glacials.

3. Dipole-dipole I.P.

Dipole-dipole I.P. was carried out on line 192S between 2000E and 2600E to evaluate small gradient array I.P. anomalies at 2330E and 2230E as well as mineralisation found in a mine dump 80 m north of 2360E. The dipole spacing was 40 m and two superimposed setups gave a station interval of 20 m for better resolution (Figure 54). No significant anomalies were detected. Resistivity and chargeability data are given in Appendix J.

2.4.6 Conclusions

1. Geological mapping, geochemistry and geophysics all delineated a