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## REGIONAL GEOPHYSICAL ASSESSMENT HENTY RIVER EL 10/85

for  
OCEANA TASMANIA PTY LTD  
by  
DR. D.E. LEAMAN

**OPEN FILE**

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## SUMMARY

Regional gravity and magnetic data in the Henty River region (EL 10/85) have been inspected. Although coverage and geological information are both limited it is possible to show that much of the region is non anomalous and properties are consistent with reasonable inferences and structures.

The geophysical data do show that the disrupted offset form of the Firewood Siding Fault is produced by transverse faults which can be traced across the area. The Firewood Siding Fault has a complex history and was active in Early Ordovician time as shown by the local thickening of siliceous rocks south of it. The fault and its splays tend to reproduce a nearly E-W orientation which mirrors a possible fundamental feature a little north of the present fault. All transverse structures and changes in magnetic character are affected by this feature.

The Queensberry mineralisation possesses a recognisable signature even though data are limited. The lode system or mineralised volume may be larger than previously thought and is clearly related to a Devonian rejuvenation of the Firewood Siding Fault.

Other magnetic anomalies cannot be adequately explained. The largest positive features each contain a fault segment but the affected volume is quite large and not a fault surface skin. It is possible that a negative anomaly on the inferred controlling structure, which also terminates fault zone anomalies, may represent alteration. These features, and the inferred axis of control are the only prospective parts of EL 10/85 and ground work is required to assess the materials actually present. All occur at surface or very shallow depths.

## INTRODUCTION

EL 10/85, known as Henty River, is located midway between Queenstown and Zeehan. The Henty River bisects the licence area (52 sq km) - see Figure 1. Oceana Tasmania Pty Ltd also holds a mining lease (27M/83) of 32 ha centred on the old Queensberry Mine workings (see Figure 2). The lease is completely surrounded by EL 4/78 (Cyprus Minerals Australia Co) which is, in turn, enclosed by EL 10/85.

The region is rugged, variably vegetated and relatively inaccessible. Ground-based exploration is inevitably difficult and costly. The area has not previously been explored in detail. There are several reasons for this. The mapped geology is hardly encouraging with respect to any mineralisation models, access is difficult and few hopeful indicators have been defined, with the exception of Queensberry itself, which might justify exploration effort. This historic, wholly negative, appraisal has clearly influenced other explorers. Oceana Tasmania Pty Ltd developed a similar view in the first year of the licence grant and relinquished it in December 1986.

After the statutory reserve period, in which there were no expressions of interest in the area, Oceana Tasmania Pty Ltd was asked by the Department of Mines whether it wished to reconsider its relinquishment. In view of its Queensberry holding and a proposal to settle, by regional analysis, the exploration potential of the area it was decided to reclaim the licence. The licence was renewed as EL 10/85 but application was not made until April 1987 and fees were not paid until May 14. The writer finds this history fascinating and unusual.

Work proposed for the effectively renewed licence included literature review, assessment of magnetic data for alteration or mineralisation responses, analysis of gravity data for regional structural setting and possibly some self potential tests. The self potential tests have not yet been undertaken.

This report details evaluation of the available gravity and magnetics data bases with respect to features within EL 10/85. No attempt has been made to exclude the area covered by EL 4/78 or lease 27M/83 from this study since these must be considered integral to overall understanding of EL 10/85.

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PREVIOUS EXPLORATION

Several workers have commented on, or described, the Queensberry lode system which has yielded 790 tonnes of high grade lead ore (McIntosh Reid, 1927). This lode system has not been rated particularly highly by Forsythe (1968), Lennox (1970) or Goldner (1985) although each felt that the presence of even small ore bodies may indicate other more economic mineralisation nearby.

Detailed work at Queensberry has suggested that the lode system may be more extensive than McIntosh Reid thought but the extent of the mineralisation is not known.

No significant geophysical surveys, or regional data analysis, have been attempted previously either in the vicinity of the Queensberry Mine or in the surrounding areas.

Gross, general analysis of the Henty Region has been reported. While not exploration, as such, it has been summarised on page 4.

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## GEOPHYSICAL ASSESSMENT

## DATA:

Regional aeromagnetic and gravity data have been used for this assessment.

The magnetics data presented in Figure 3 are a segment from the Mines Department 1981 survey (see Corbett et al, 1982 and Leaman, 1986a). The magnetic field is relatively stable across EL 10/85 with several minor (< 5 nT) features and three primary deviations. While there is no immediately obvious explanation for the N-S feature (C), features A and B lie at fault intersections as mapped by Baillie (1977). There is no direct correlation between field character and exposed lithologies and it may be concluded that the gross characteristics of the field - values reducing to the north and east - are regionally induced either by materials outside the survey segment or by magnetic units at considerable depth.

The gravity data, drawn from the Mines Department TASGRAV and MT READ data bases, also present a diffuse pattern (Figure 4). This is to be expected in an area with a nominal station spacing of 1 km and where there are significant gaps. There is, however, a gross correlation with the aeromagnetics in that pronounced N and E gradients are evident. In all other respects the patterns differ and there is no direct correlation between the geology (e.g., Figure 2) and Bouguer anomaly. The Firewood Siding Fault, for example, is not reflected and the broad, relatively positive feature extending from the SW correlates with only part of the exposed Cambrian rocks north of the fault.

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**REGIONAL INTERPRETATION:**

A regional structural and stratigraphic setting for EL 10/85 has been established by Leaman (1986 a,b). The elements of the regional deductions are summarised by Figures 5 to 8. Neither profile set applies centrally to EL 10/85 but the situation north and south of the Firewood Siding Fault is sampled. Thus Figures 5 and 6 reflect the Siluro-Devonian cover and possible gross folding south of the fault and Figures 7 and 8 the virtual exposure of the Cambro-Ordovician north of it.

The overall setting established by this gross evaluation is important to any detailed review of a small area such as EL 10/85. In the gravity case strong positive anomalies crest within it and there are major gradients induced by structures of crustal scale. These need to be appraised and removed (or understood) prior to specific study since their three dimensionality and scale would diminish the reliability of any regional - residual separation. Similar, but subtler, problems apply in the magnetics case.

The strong gravity anomalies and gradients are due to the existence of a thick Cambrian trough fill beneath the Henty River region. The axis of this trough lies across EL 10/85. The gradients are generated by the contrast between the trough materials (Dundas Trough) and the crystalline basement to the east. Figures 5 and 6 offer alternative structural relationships which are not critical to this study. Both involve significant mafic and intermediate volcanic sequences at depth. Only the gravity models adequately scale the section.

Regional trends inferred from the gravity and magnetic data available in 1986 were presented by Leaman (1987) and related to mineralisation. The mineralised sites shown were derived from extant but incomplete prospect lists then in preparation by the Mines Department (e.g., Bamford and Green, 1986). The Queensberry Mine lies close to an E-W corridor and there are intersecting NE and NW trending features (Figure 9). This crude representation was based on subjective analysis of regional data. The magnetic data base has not been upgraded but the gravity coverage was improved in the Henty area during the 1986-7 summer season (refer Mount Read and TASGRAV data bases - Mines Dept). Trend definition is poor near the coast due to increasing subtlety and the need for more comprehensive analysis. There is, however, much evidence to suggest that the mineralisation near Queensberry is no accident. What is not clear, in Figure 9, is the precise focus for source fluids or alteration.

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**LICENCE AREA STUDY:****Gravity**

The gravity data (Figure 4) reveal that the axis of the positive anomaly lies a little east of the Professor Plateau. Values diminish rapidly northward across the Professor Range but increase southward. The Firewood Siding Fault has no impact on the observed responses. Other significant features (gradients) occur east of Melody Creek and near Marchfly Ridge. Available geological mapping offers no explanation for these observations.

Since the data is not site specific and the station spacing is relatively coarse no mineralisation responses can be detected and the treatment has been limited to a structural setting study based on three profiles. The principal objective of the gravity study was provision of a local understanding of structure and units.

The models shown in Figures 10 to 12 present shallow detail and mantle forms only. Regional presumptions illustrated in Figures 5, 6 and 7 have been built in but not reproduced in all cases. The general basement - trough relationships and the regionally inferred lithological content are fundamental to all profile fits. The calculated shift differential required (approx 55 mGal) is consistent with current whole crustal knowledge for W and NW Tasmania using the observed Bouguer anomalies.

**Profile 5343:**

This profile samples the section south of the Firewood Siding Fault. It is not well defined but there is sufficient information to show that the crestal anomaly is compound (Figure 10).

The model reveals that the Permian cover at the western end of the profile is not significant. The Tertiary cover is too far west to include. The shift differential shows that the pattern is not directly related to the trough section. The subtle deviations are related to fold wedges of Ordovician materials. This is more clearly demonstrated in other sections. Although scaling is not particularly reliable the thickness of Siluro-Devonian cover immediately south of the fault probably does not exceed 1000 to 1500 m. The profile fit was not perfected due to uncertainties and gaps in the observed data. The negative deviation at about 371 000 mE is unexplained but may represent data errors or thick pockets of low density materials in the river bed since the stations upon which this feature is based lie in the valley notch.

**Profile 5345:**

This profile transects the Queensberry block, the Cambrian section north of the Firewood Siding Fault and part of the covered section south of the fault. The profile is well defined (Figure 11). The smooth gradient across the Queensberry block and east of the Henty River (372 000 mE) is essentially the regional gradient generated by the

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basement - trough relationship and the mantle. Its smoothness shows that the Queensberry block is either homogeneous at this scale or gravimetrically insignificant - probably the latter. The use of basic densities about 2.74 t/cu m for the Cambrian rocks demonstrates a largely sedimentary section with minimal amounts of igneous materials.

The abrupt negative anomaly near Marchfly Ridge is not explained but certainly appears to be a data problem. The gentle step reduction to the east reflects the presence of Ordovician siliceous rocks at the Sisters (to the north) and within the covered fault block to the south (from 371 to 373 000 mE). These are essentially three dimensional effects not well reproduced on this profile.

The small negative depressions at the western end of the profile (SW of the Firewood Siding Fault) can be wholly explained by pods of siliceous Ordovician rocks beneath the Silurian cover. Various other explanations were sought, including granite effects or simple light unit folding but none consistently satisfy the shift differential, regional or structural requirements.

#### Profile 5348:

This well-controlled profile (Figure 12) extends across known Ordovician rocks of the Professor Plateau and the Cambrian sequence to the Henty River. There are no concealment problems.

It will be noted that the character of the profile is very similar to 5345 (Figure 11) and it is easily proven that the western negative response is related to the Ordovician rocks. These are shown to be very thick locally and to thicken across the Firewood Siding Fault and its splays.

The Cambrian sequence is essentially sedimentary and relatively homogeneous. The slight over sharpening of gradients near 367 000 mE can be explained by an increase in shale content or equivalent lithological variation. Its thickness and lateral extent must be limited and it would seem to extend no further east than 368 000 mE. A small fault may be inferred.

The poor fit near the eastern end of the profile suggests inadequate allowance for the form of the northern end of the Sisters conglomerate, incorrect mapping of the western conglomerate boundary or poor gravity field definition. The observed and modelled profile shapes are consistent suggesting that the 1 km spacing may be too coarse to properly define the local response.

The gravity analysis, though basic, does show that the features defined in the regional coverage can be wholly explained in terms of the mapped geology. No abnormalities need be invoked. No mafic successions occur nearby. The absence of any response differential across the Firewood Siding Fault and the thick localised Ordovician accumulations related to it imply an active zone and growth fault development during deposition of these rocks.

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## Magnetics

The magnetic field at some 150 m elevation (Figure 3) is relatively smooth across EL 10/85 but the presentation is dominated by three obvious isolated features and one comparable, but subdued feature.

The regional perspective (Figure 8) shows that the positive rise is very broad and its crest is west of EL 10/85 (and west of the gravity axis). The eastern gradient, with its minimum near the Zeehan Highway, is the dominant feature of the area under study and the source is clearly a significant magnetic volume. Yet, inspection of the field (Figure 3) and the profiles (Figure 13) shows that no exposed block, sequence or unit can be directly correlated with any magnetic feature. No exposed lithology carries a measurable contrast at 150 m.

This property inference confirms the gravimetric implication; there are no masses of igneous rocks - especially intermediate or mafic - within hundreds or thousands of metres. Inspection of the profiles also suggests that the smooth profiles for 5343 and 5344 N are what would be expected from continuation of other profiles to equate depth of cover by Siluro-Devonian rocks. A profile for 5342 has not been included in Figure 13 since the herring bone effect evident in Figure 3 near this northing implies a survey compilation error. Continuation of profiles 5346 to 5350 would not, however, yield profiles 5343 and 5344 if the anomalous features were included. This suggests that these features are not original or pre-Devonian since any Cambrian lithology, for example, would recur south of the fault and a subdued magnetic response would be evident. Indeed, two of the magnetic features (A, B) are fault-related and occur either along mapped fault segments or at fault intersections. Both have extensions, slightly subdued, along the local faults and both are related to the Queensberry fault block. These relationships indicate a Devonian, or last movement, age for the features and their sourcing property changes.

Feature C, extending N-S, is not as simply explained with existing geological knowledge. Figure 17 presents a compilation of detailed trend inferences and known or implied structures within EL 10/85. The mapping of Blissett and Gulline (1962) has been coupled with that of Baillie (1977). This shows that two probable faults extend southward along the flanks of anomaly C and down the valley of Malcolm Creek to 5347 000 mN. This fault block presents a mapping mismatch across the map edges and the zone should be reviewed. The magnetic response, however, is quite clear; the same material is involved throughout but is truncated or, possibly, offset from anomaly B.

It is thus possible that all three features represent fault zone shatter, alteration or oxide introduction effects. The anomalies appear comparable and occur independently of the particular exposed lithologic couplings. There is no evidence of lithologic association. Examination of the anomaly forms and couplets suggests magnetisation by induction only.

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Three profiles (5343, 5345, 5348) have been reviewed in more detail for structural content and comparison with the gravity interpretation. The constraints discussed above have been incorporated and shift differentials consistent with regional profiles applied.

Two basic solution styles were tested. The first presumed that the general form of the profiles, as exemplified by 5343, represented the negative tail of a positive feature to the east while the second presumed the presence of a deeply buried, slightly magnetic, folded sequence of Cambrian rocks in the same way as suggested regionally (Figure 8).

It was found that either presumption could generally account for the gradients observed with realistic geometries and contrasts. Only the second, however, could account for the subtle concavities in the profiles. This is a critical requirement since the first option presumed no magnetic contrast west of 372E and this precludes the possibility of allowing negative contrasts to account for concavity. The second option is thus preferred and more readily accounts for effects north and south of the Firewood Siding Fault.

#### Profile 5343:

A solution is suggested in Figure 14. The gradient is fully accounted by a 2500m section of slightly magnetic, folded Cambrian rocks. Uplift or displacement at the Henty Fault accounts for the field east of the profile. The local minimum response is overstated which may reflect 3D effects, some contrast in the fault zone or a greater terrain clearance. The latter explanation is most likely, due to the ravines occupied by the Yolande River and tributaries, although the fault corner termination of the eastern rock suite (off profile) would generate a 3D rise in values. The model is relatively sensitive to changes in shape of the upper surface but the actual depth is uncontrolled. The overall form of the modelled body is consistent with dip patterns in gross terms.

#### Profile 5345:

Similar comments apply to the solution offered for profile 5345 (Figure 15). The body is certainly terminated by fault splays from the Henty Fault. The synclinal, or down thrown, portion east of 370 to 371000E may relate to the surface faults in this region. The offset or step at 365 to 366000E is apparently related to the Firewood Siding Fault and, as the model is very sensitive to upper surface changes, would imply a throw of only 500m at this depth in this unit. Due to the effects of fault growth during deposition of Ordovician rocks (see previous discussion - gravity) the present day, post Cambrian, throw is much greater. This profile does not adequately sample anomaly A (Figure 3) but provides a section of anomaly B. Review of this feature suggests a bulk contrast of around 0.001 cgs and a relatively tabular east dipping source at least 200 m thick. Since the source is not considered to be a discrete

unit but a chemical change in the fault zone it may be presumed that this segment or offset of the Firewood Siding Fault - as it trends NNW - dips east. This may also be inferred by the shift of the anomaly east of the fault bend and its truncation by the E-W segment at 369000E. Detailed examination has shown that anomalies A, B and C cannot be explained by point sources or very thin, single tabular, sources. A substantial rock volume is involved or the sources are multiple within that volume.

**Profile 5348:**

A solution for this profile is presented in Figure 16. The basic elements are as discussed in other profiles. The source for anomaly C, however, appears to dip west (compare anomaly B). This behaviour would be consistent with a source along a fault (unmapped by Baillie, 1977) in the valley of Malcolm Creek. Such a fault could extend from structures mapped by Blissett and Guilline (1962) to the offset across Marchfly Ridge (see Figure 17).

It may be useful to speculate on the origin and disruption, or discontinuity, of anomalies B and C. It has been argued that each is related to fault zone alteration. Yet regional compilation shows that extensions of the faults implied, or mapped, do exist. Indeed Grieves prospect (365000E, 5350000N) occurs along trend from anomaly B. The anomalies suggest that only parts of the structures involved continued an active existence or remained conducive to fluid transfer, perhaps due to local lithocontrasts.

In a more regional context E-W linears can be recognised in the magnetic field at about 5347000N and such a feature could separate anomalies B and C as well as terminate anomaly A. It is also possible to suggest a feature, parallel to the fault mapped along the northern edge of the Queensberry block, through the saddle in anomaly A (363500E, 5346000N) and between anomalies B and C (Figure 19). The nominal E-W features could well locate the fundamental control on structures in this region and also influence fluid movement.

## MINERALISATION RESPONSES

Little mineralisation is known within the region of EL10/85. This may reflect absence, limited exploration or difficult terrain.

The lode mineralisation at Queensberry Mine is associated with fault and fracture systems and may represent deposition during movements of the Firewood Siding Fault. These movements may have been late and the mineralisation could have been sweated from deep in the Cambrian section and introduced along the fault. Or it could be Devonian and granite-related.

The Queensberry Mine region possesses a typical lead mineralisation response (compare Leaman, 1987). This consists of a subtle positive effect in a variably depressed field. The second part of this equation is not especially obvious in Figure 15 but it is present. The first part, with amplitude of perhaps 6 nT, is evident. The affected area, though not well defined by the available survey with a line spacing of 500m, appears to extend between the two approximately E-W faults forming the Queensberry block at about 366500mE. The only possible comparable response appears to be at 370800E, 5347200N which lies on, or near, the feature which separated or generated anomalies B and C.

Possible alteration responses are suggested at 5346300N from 364-365300E and in the vicinity of 368000E, 5346800N. All can be associated with the same dislocating feature or trend (see Fig 17).

It is difficult to assess any of these responses given the patchy geological knowledge, absence of recognisable units with established properties and no definitive understanding of the material sourcing anomalies A, B and C.

It is clear, however, that the volume affected at the Queensberry Mine probably contains a good deal more mineralisation than the 790 tonnes recovered.

It may also be that anomalies A, B and C represent indicator alteration in this area of subtle contrasts. Diffuse or disseminated targets may be involved. An alternative view might suggest that the cross trend which fed or terminated A, B and C may be the essential element. Whichever view is taken prospectivity in EL 10/85 is limited to a band no more than 1500m wide until surface examination or geochemistry suggests otherwise.

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## CONCLUSIONS

Examination of regional gravity and magnetic data across EL10/85 has suggested.

1. Gross features of both the gravity and magnetic fields can be explained by what is known, or directly inferrable using these fields, of the geology. Geological detail is insufficient to account for more subtle features at the present time. Much of the licence area is non anomalous in terms of present geological information coupled with the regional geophysics.
2. The Queensberry mineralisation is associated with rejuvenation of the Firewood Siding Fault where a segment of the fault trends a little north of east and intersects another fracture system. The mineralisation possesses a normal West Tasmania lead mineralisation signature and the mineralised volume may be larger than previously thought. The magnetic response extends across the local fault block.
3. Three fault zone anomalies can be recognised which are apparently due to zonal alteration. None are continuous and north trending segments are dominant. All can be linked by a trend or structure a little north of east. This orientation appears to be critical and the fundamental basement controls on structures in this region may have this orientation. A chemical or magnetic origin of the fault zone anomalies has yet to be established.
4. The Firewood Siding Fault as presently exposed is a rejuvenation of deeper features which may now be located up to 2 km to the north. Analysis of forms and thickness of Lower Ordovician siliceous units indicates growth fault activity for at least part of its history.
5. The region appears to possess limited exploration potential beyond the influence of the features described above. This may reflect data coverage or the characteristics of the particular, predominantly sedimentary, lithologies.

## RECOMMENDATIONS

1. Most of the licence area should be relinquished. The areas south of EL4/78, north east of 369000E, 5348000N, north west of 367000E, 5347000N and south east of 370000E, 5346000N should not be renewed. This will reduce the EL to 19 sq. km.
2. Materials in anomalous areas (A, B, C - Figure 3) should be sampled and tested for magnetic susceptibility.
3. Some local soil geochemistry in these regions is also advised. Gold and arsenic should be included in the analysis.
4. The anomalous features should be reviewed in light of results from recommendations 2 and 3.
5. Soil geochemistry should cover the possible trend control east from anomaly A with particular attention to the area between anomalies B and C (E) and the possible target at 370800E, 5347100N (D).
6. No further geophysics is advised until other indicators confirm the possible presence of mineralisation or the origin of anomalies A, B and C is adequately understood.

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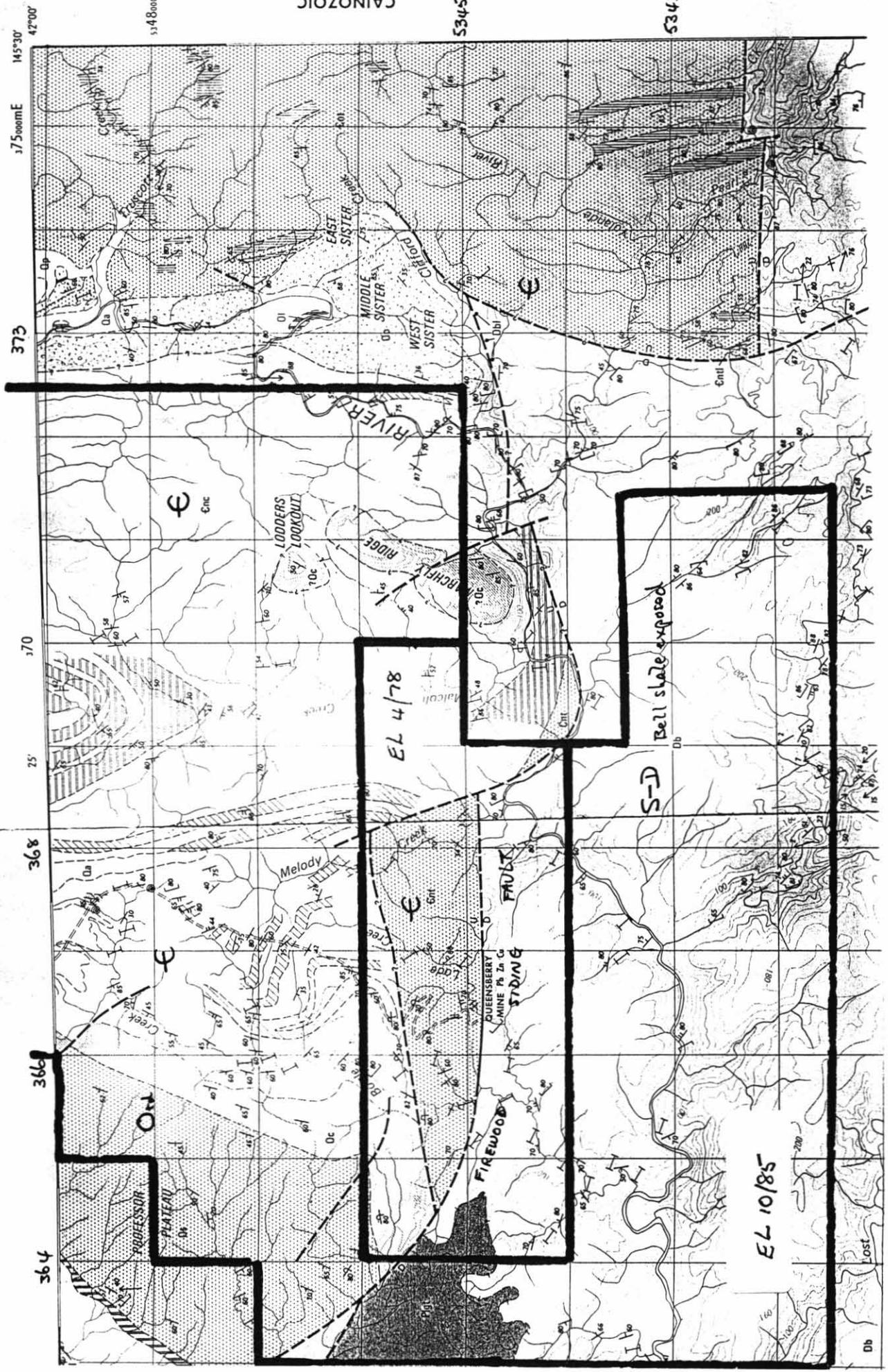


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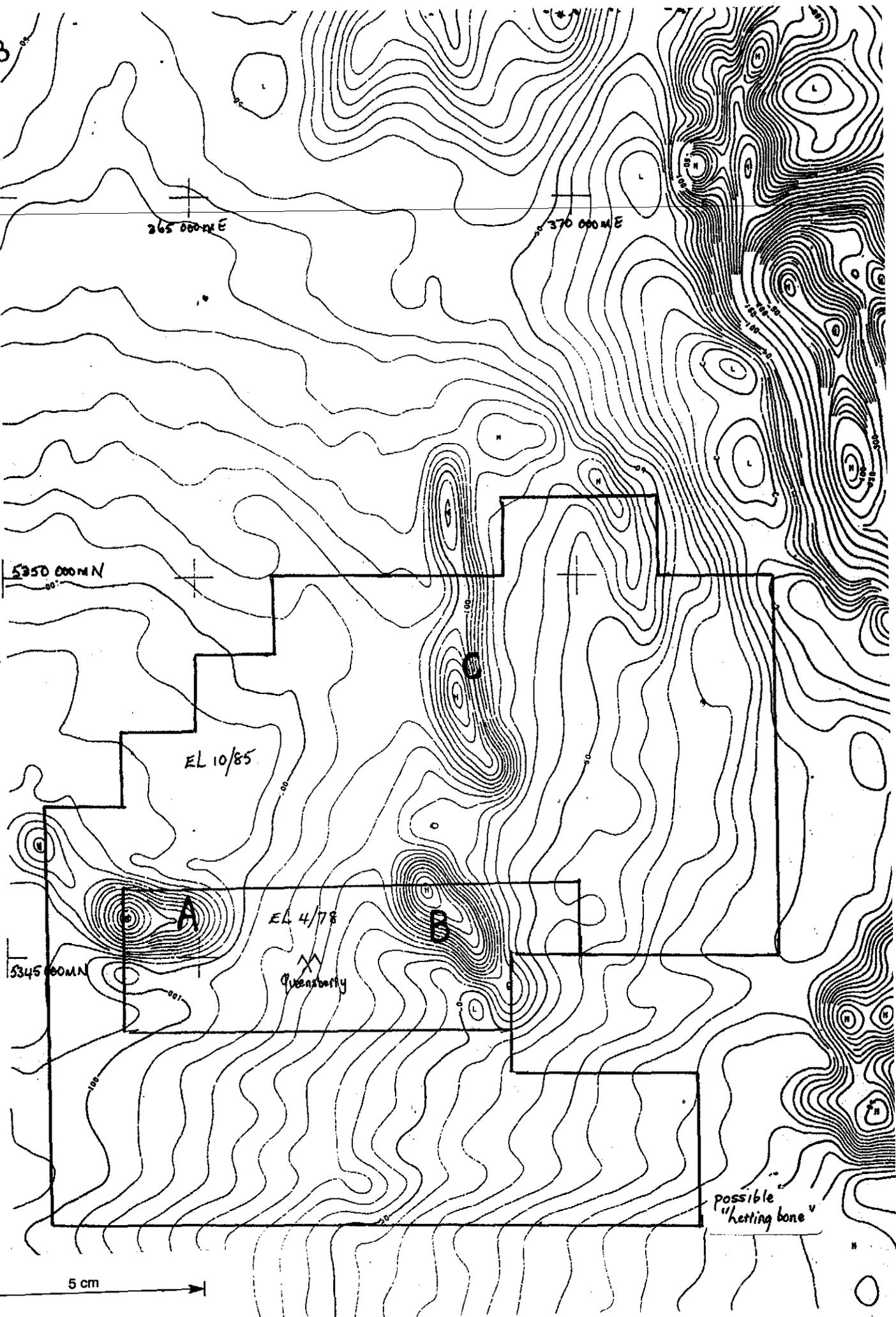
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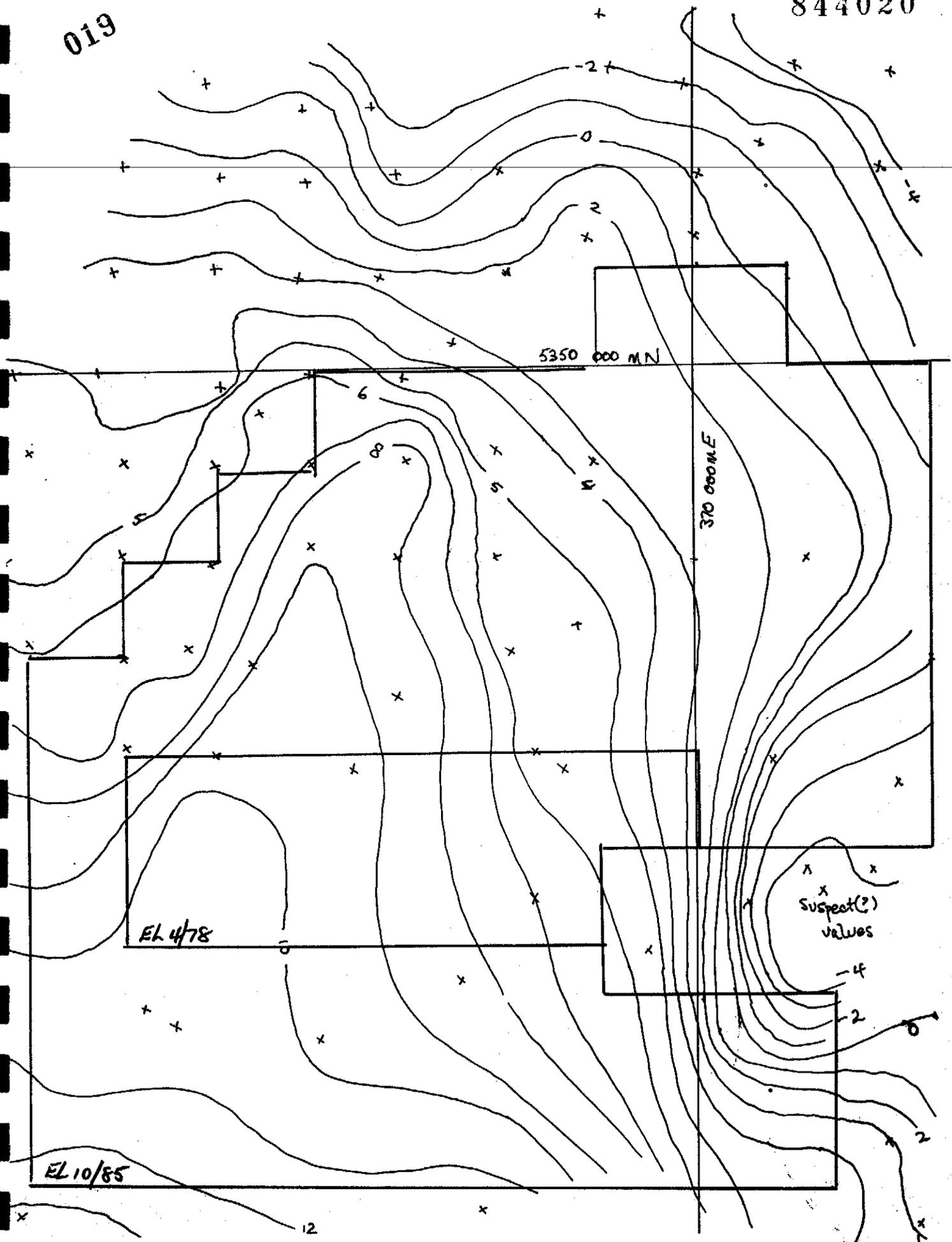


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EL 10/85 MAGNETIC FIELD (150 m aeromagnetic data) from 1981  
Mines Department West Tasmania Survey. 1:63360

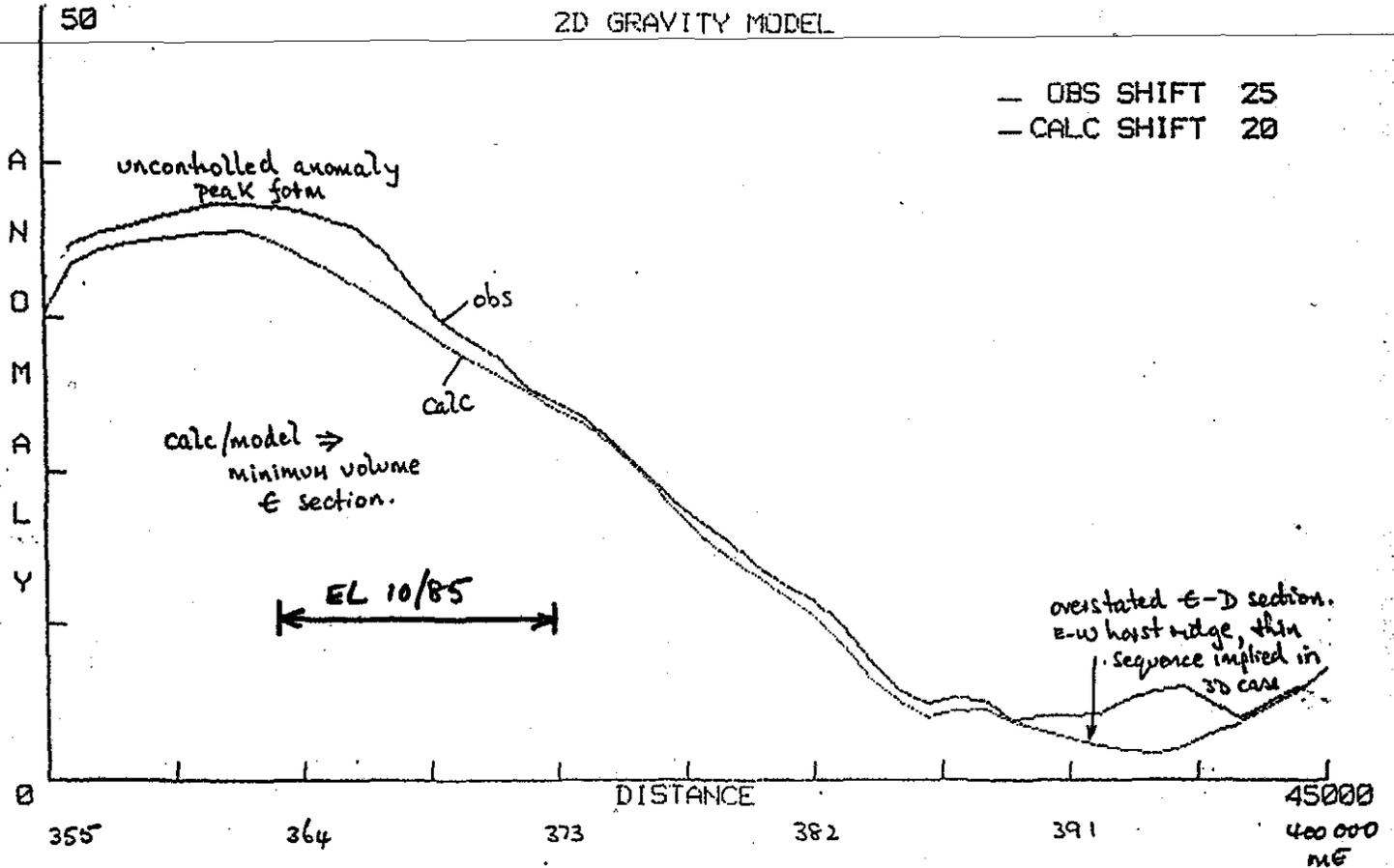
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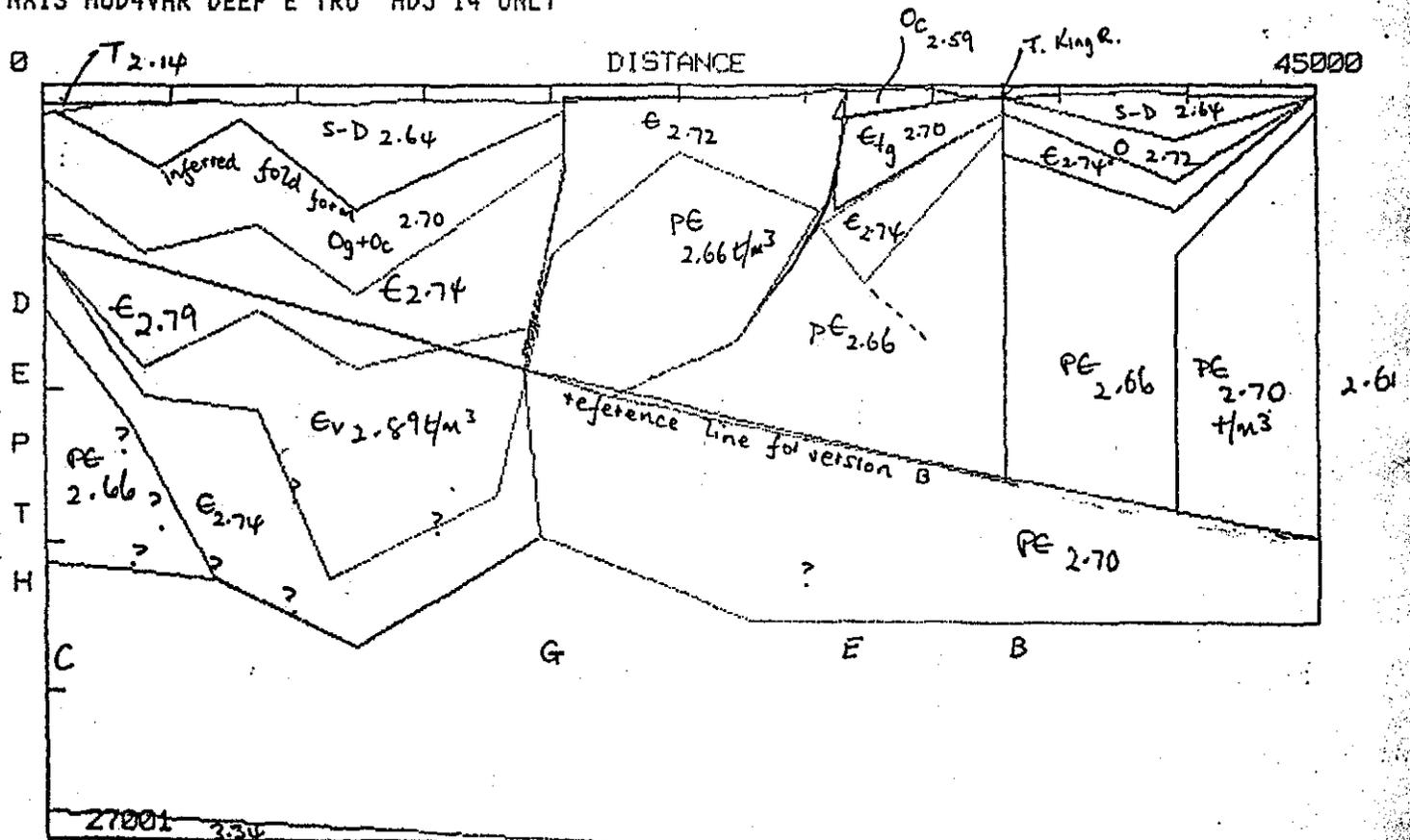


EL 10/85 BOUGUER ANOMALY in mGal - density 2.67 t/cu m  
(from Mines Department gravity data bases)

FIGURE 4



BOUGUER ANOMALY 700M LINE 343N 355-400E LYELL  
NX13 MOD4VAR DEEP E TRO ADJ 14 ONLY

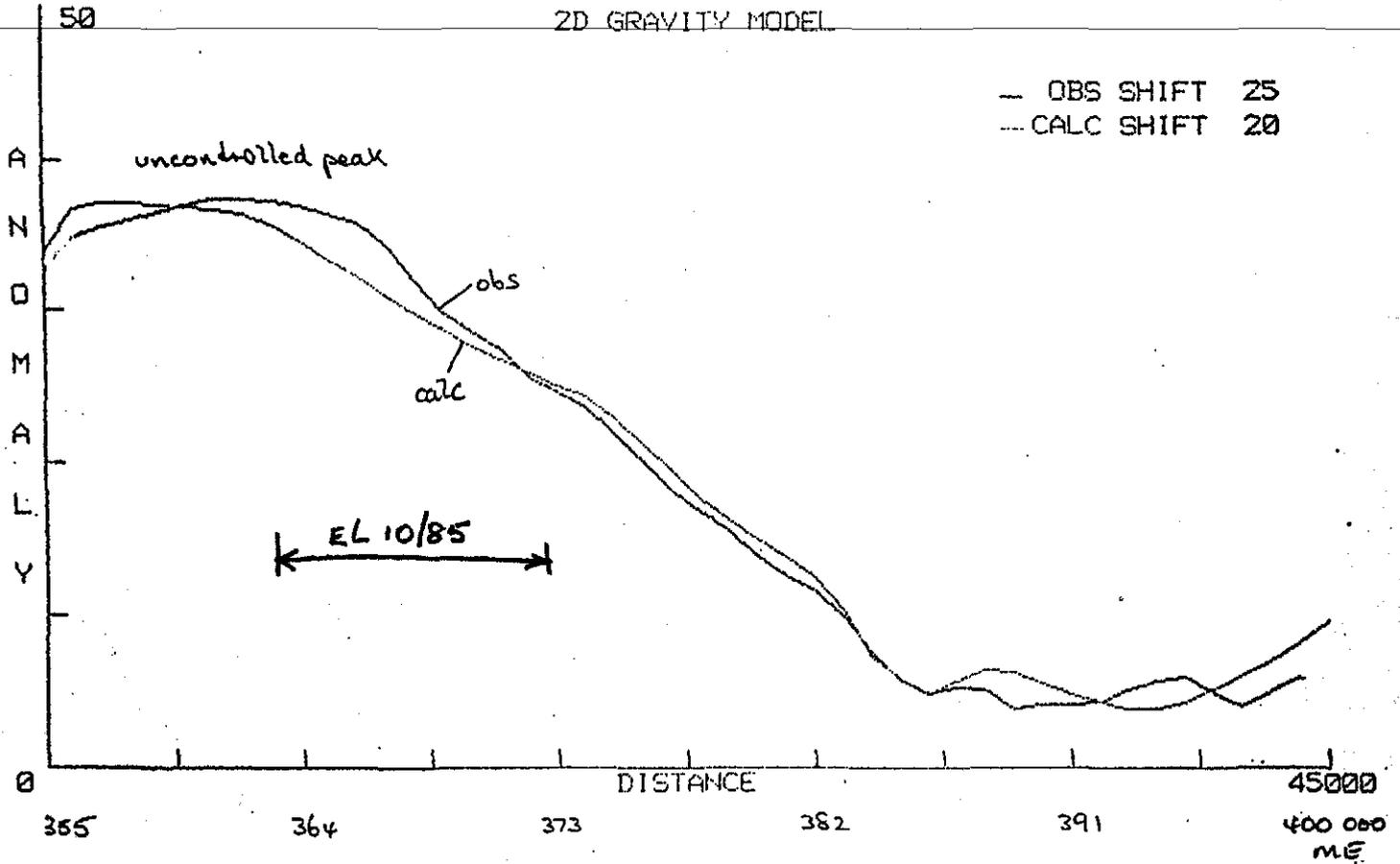


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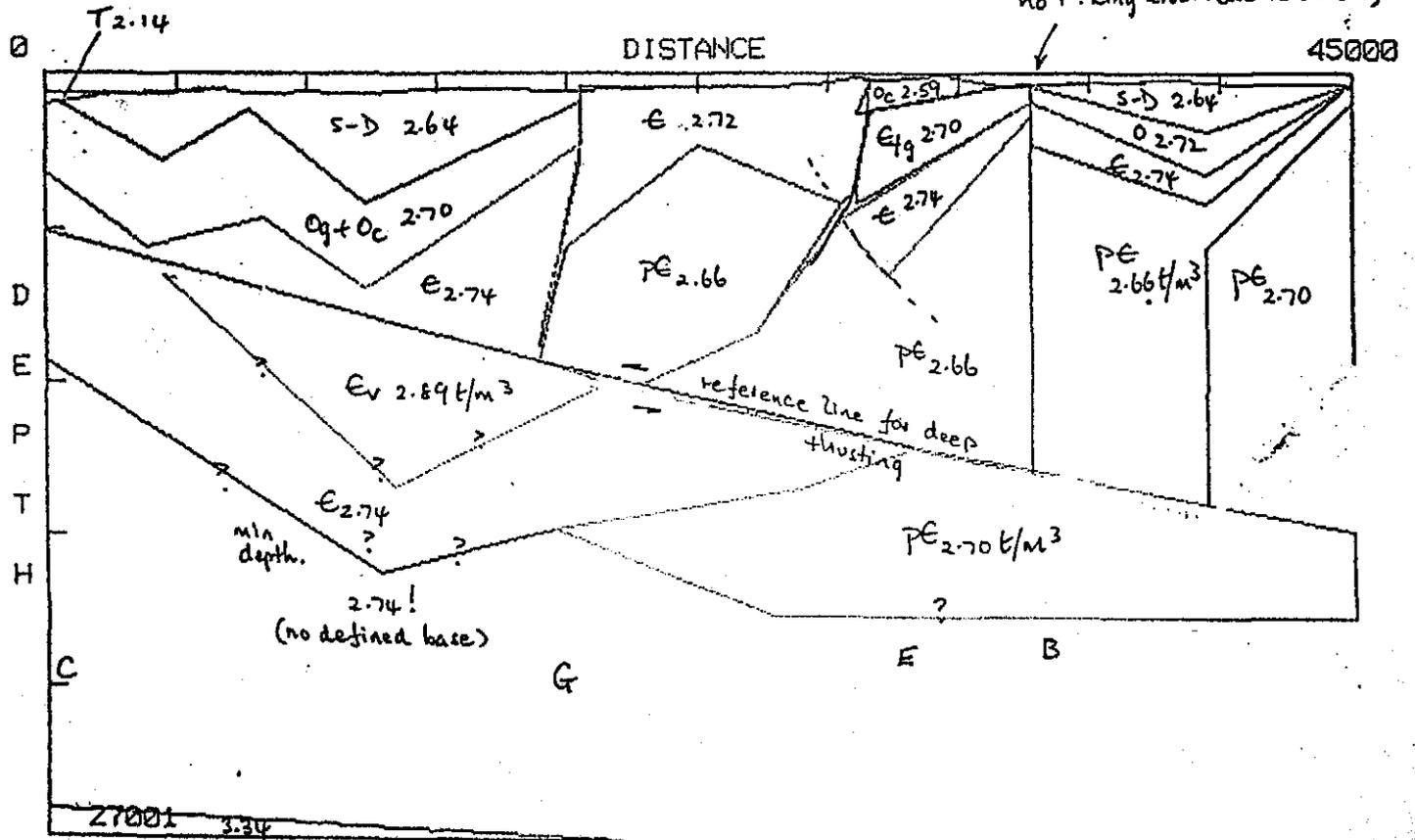
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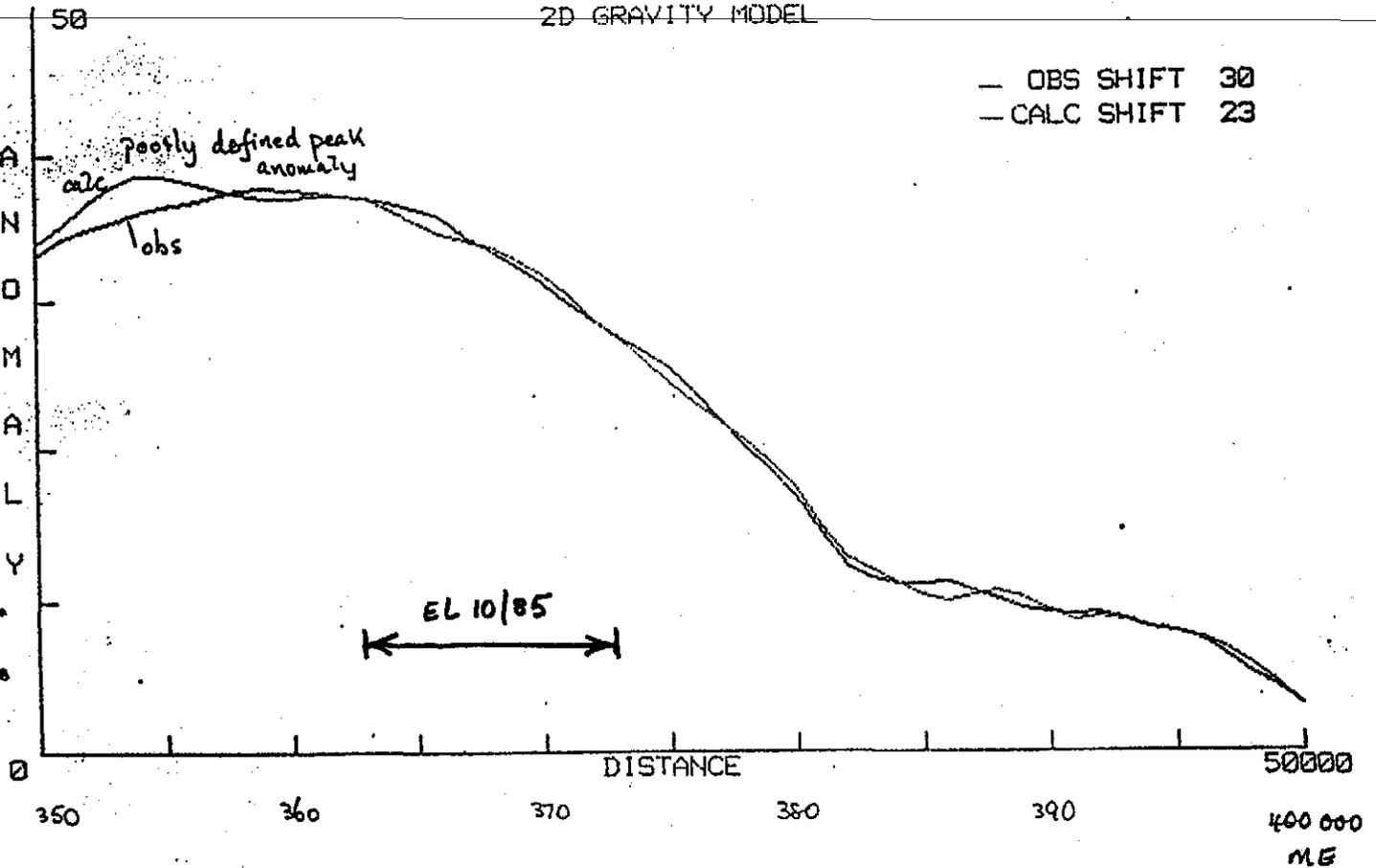
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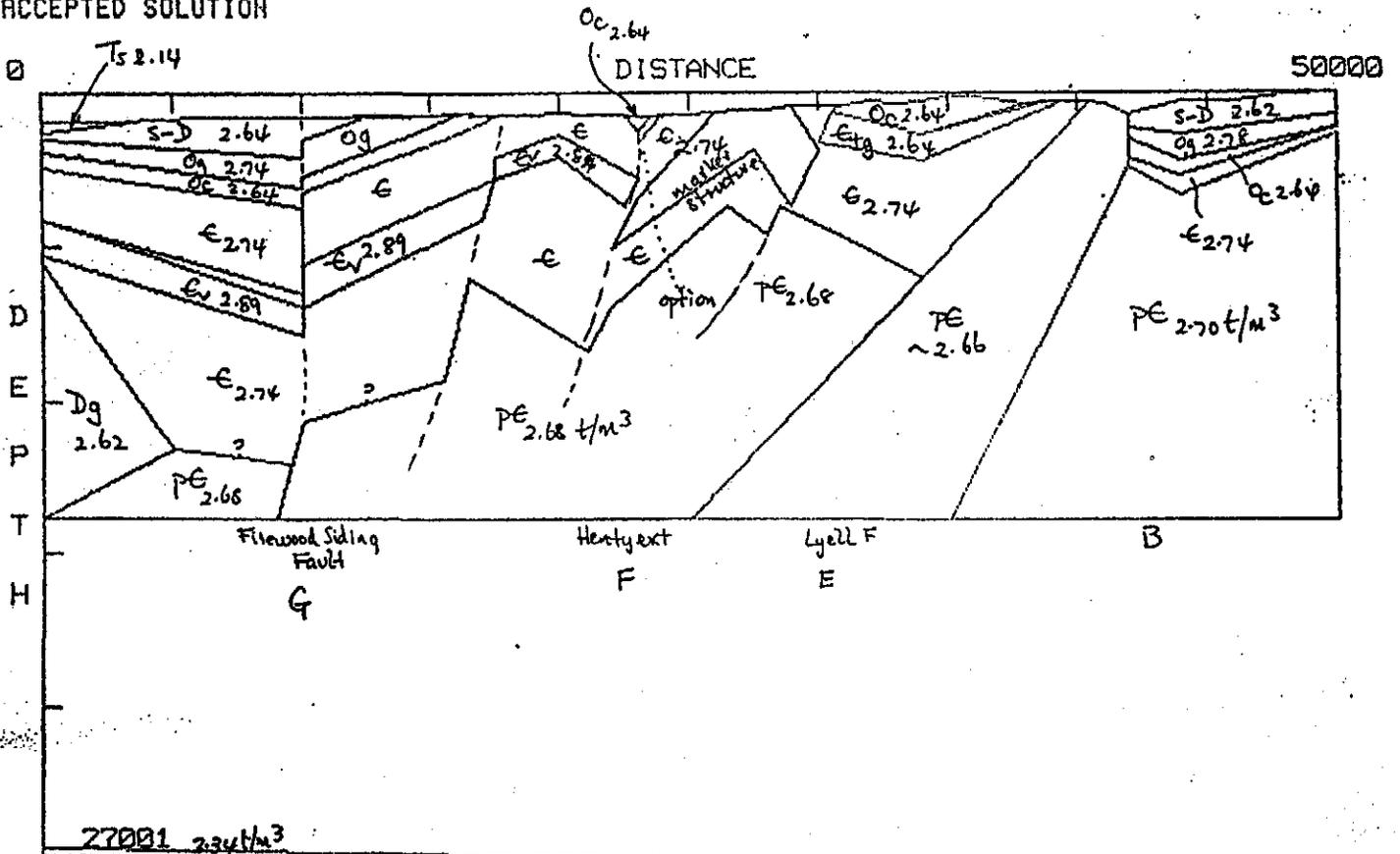
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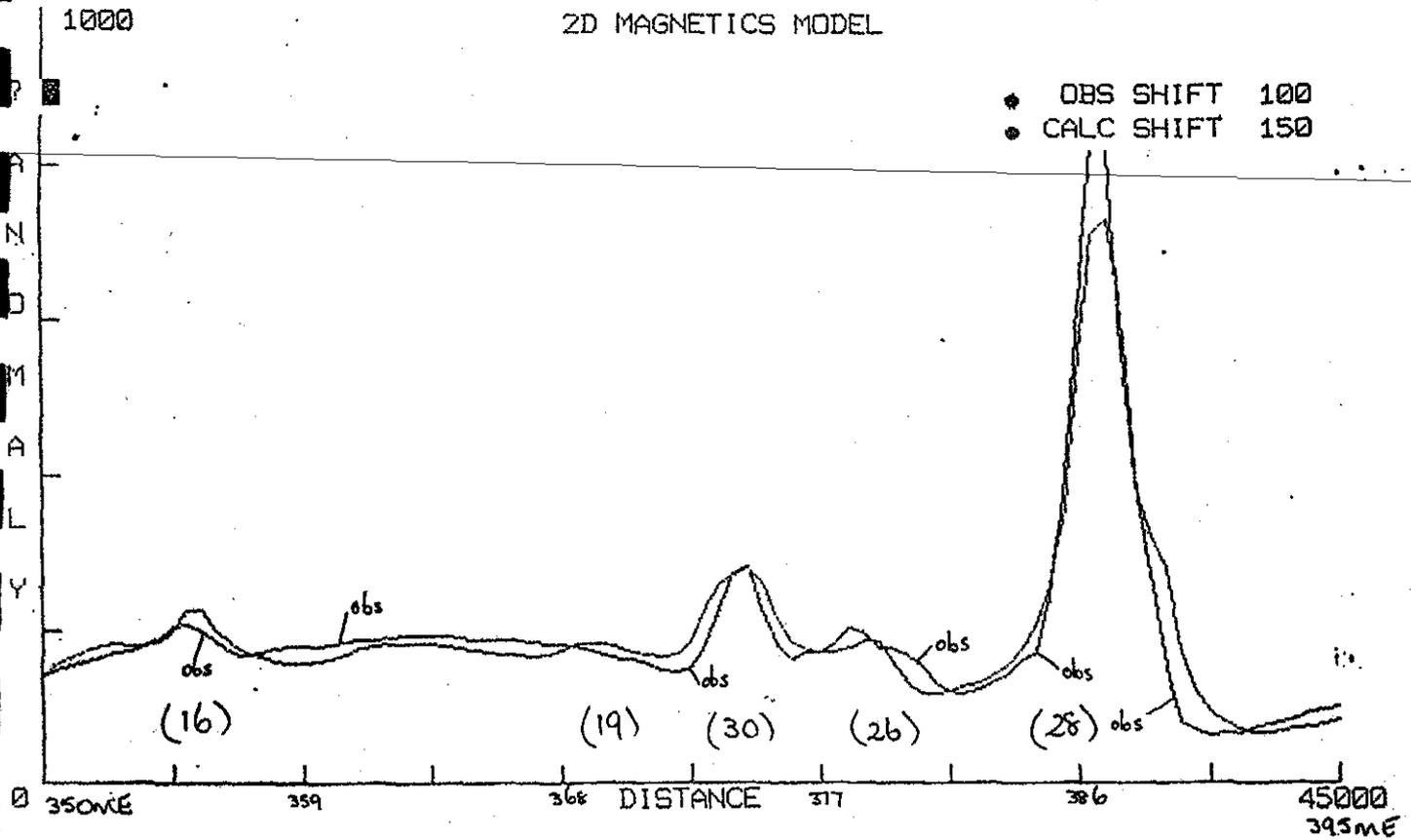
## 2D GRAVITY MODEL



LINE 350250 MN 350-400 ME  
ACCEPTED SOLUTION



2D MAGNETICS MODEL



LINE 1030 REFERENCE LEVEL 1000 M

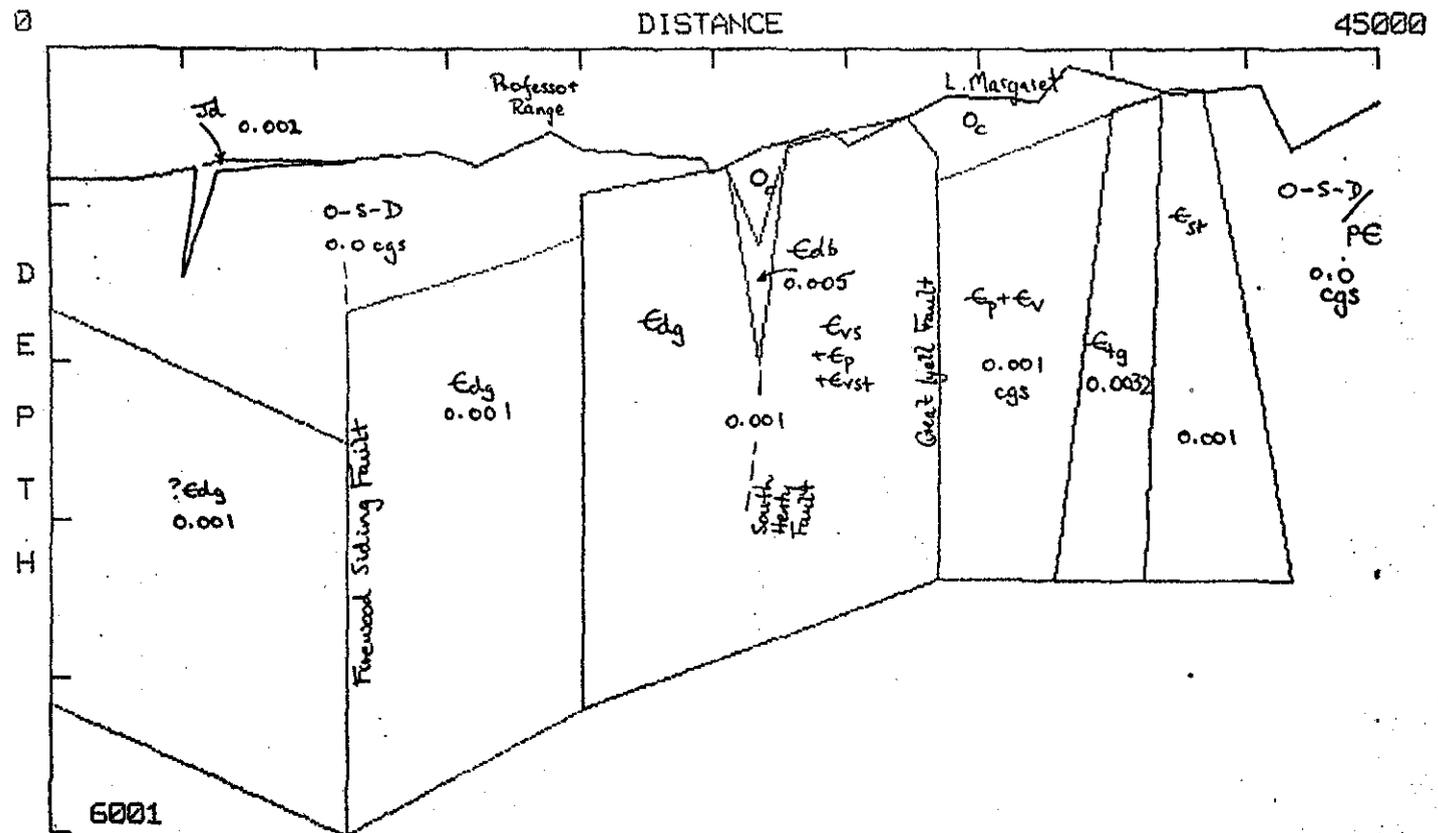
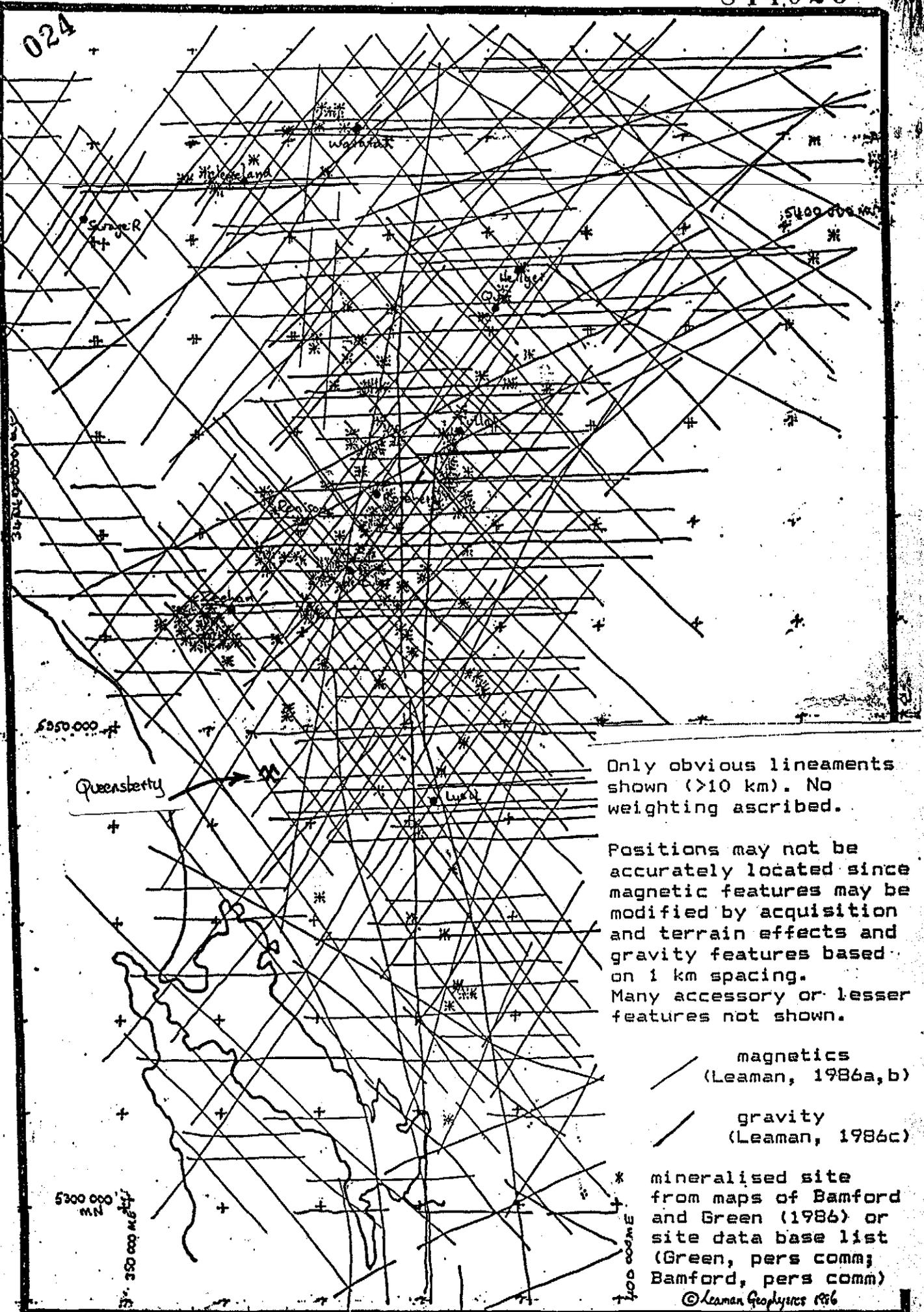


FIGURE 8

LINE 1030 - 2D REGIONAL INTERPRETATION

( ~ 5350 250 MN )

024



REGIONAL TRENDS IN WESTERN TASMANIA (from Leaman, 1987)

5 cm

025

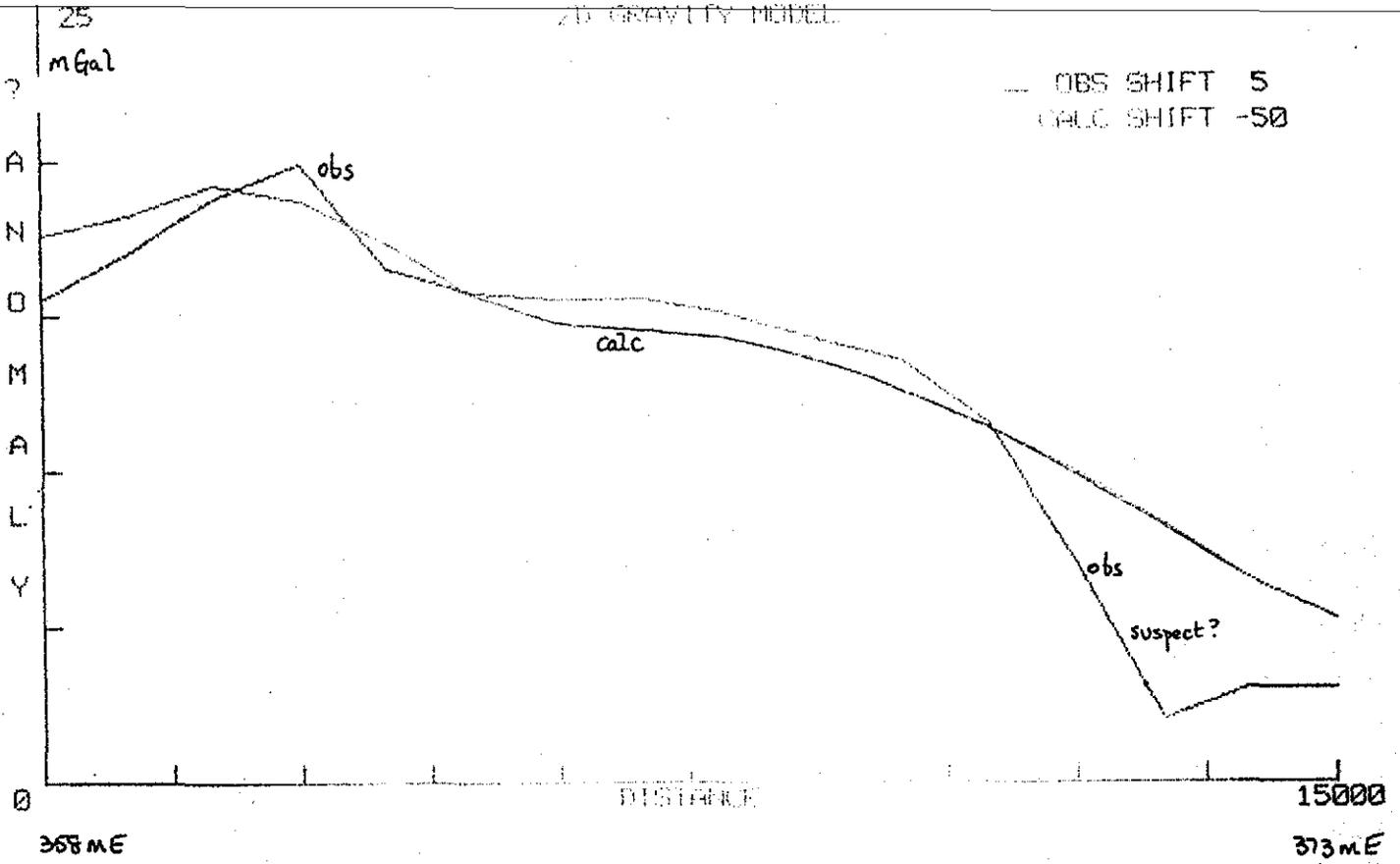
# LEAMAN GEOPHYSICS

Survey Review, Specification, Reduction, Interpretation  
Wide Experience Most Methods  
Specialties:- Gravity, Magnetics, Seismic Methods

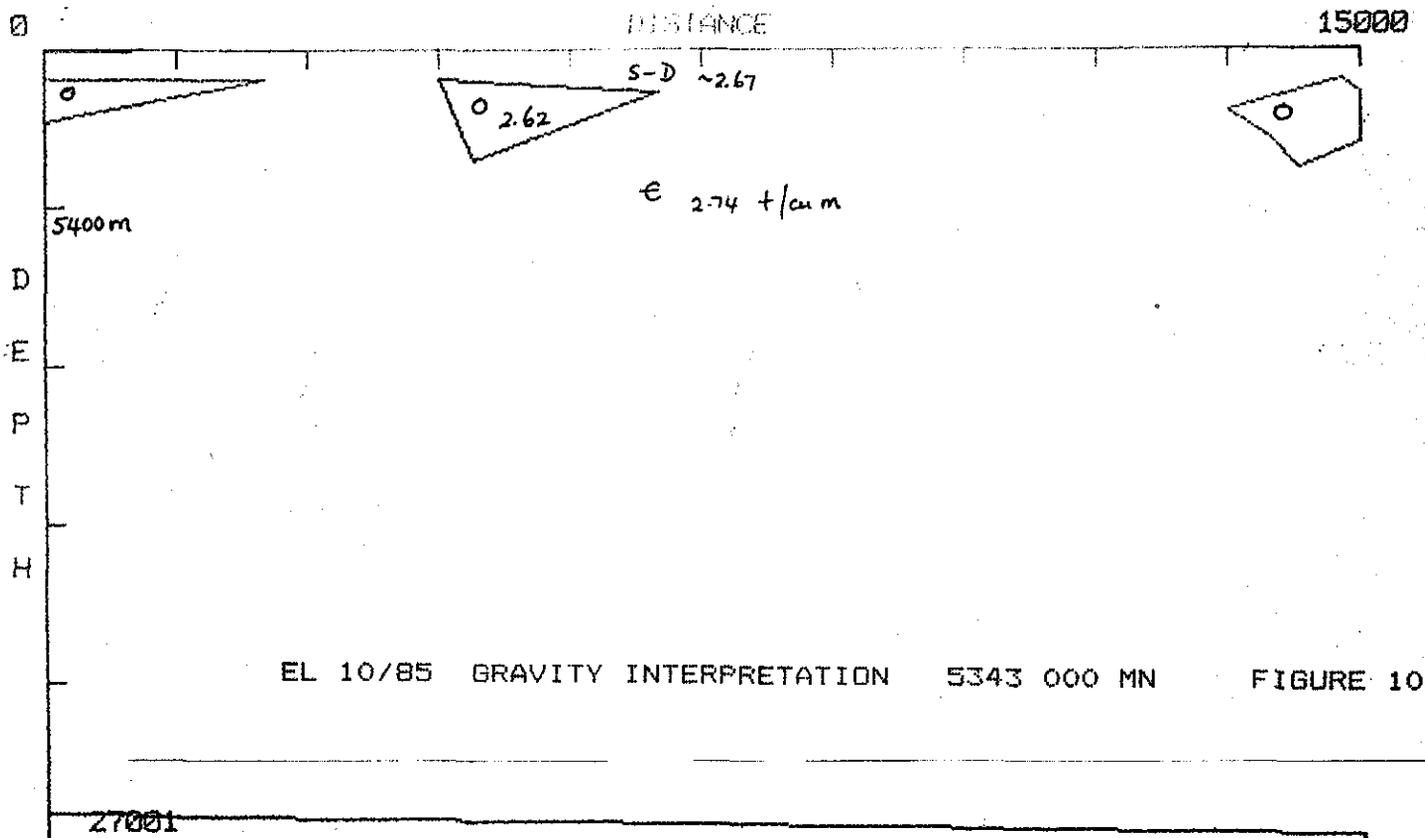
844026

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LINE PARAMETERS - ORIGIN, LIMIT, INCR : 0 15000 1000



HENTY RIVER 5343N 358-373E  
ADJ 5 8

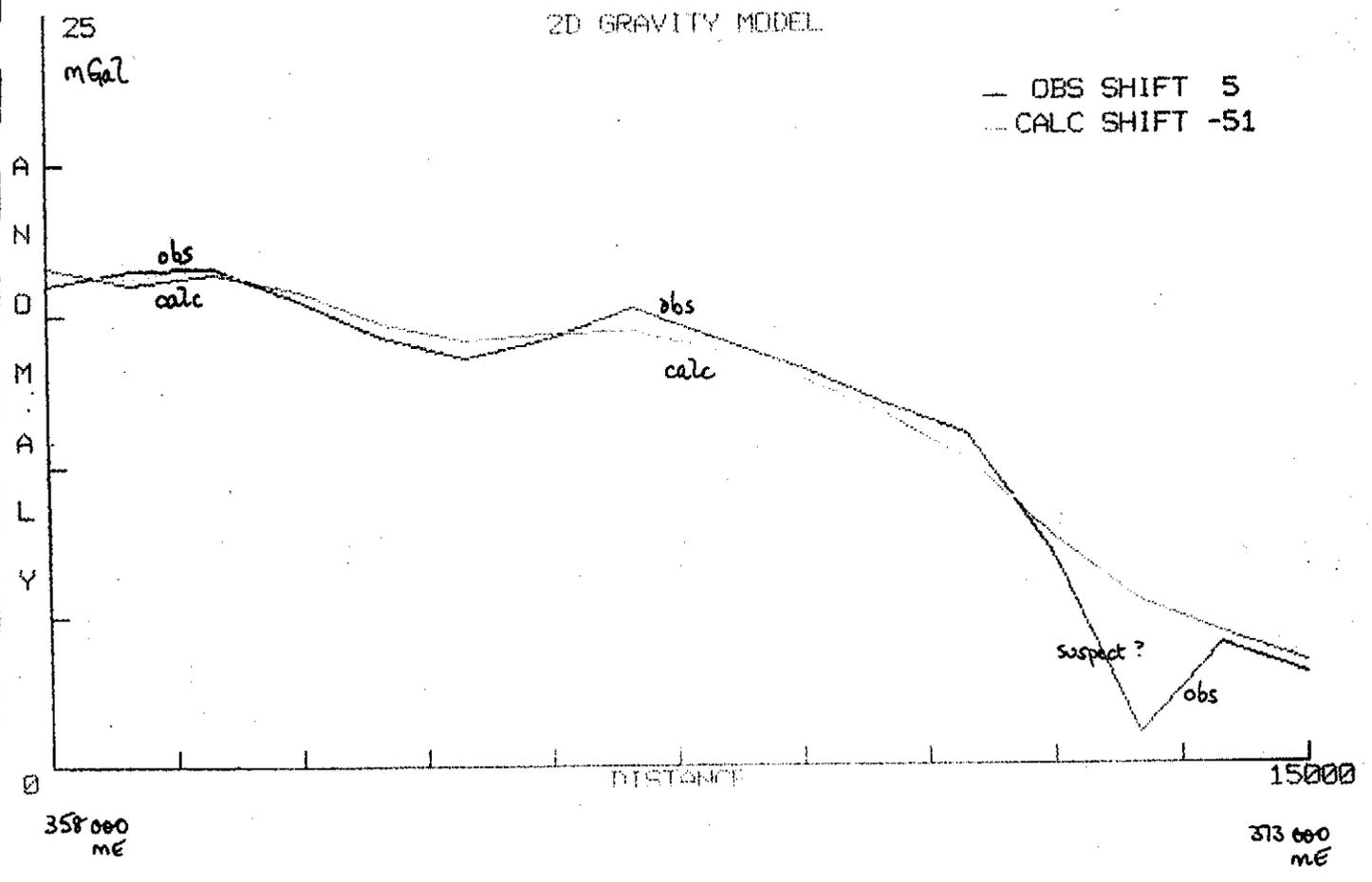


# 026 LEAMAN GEOPHYSICS

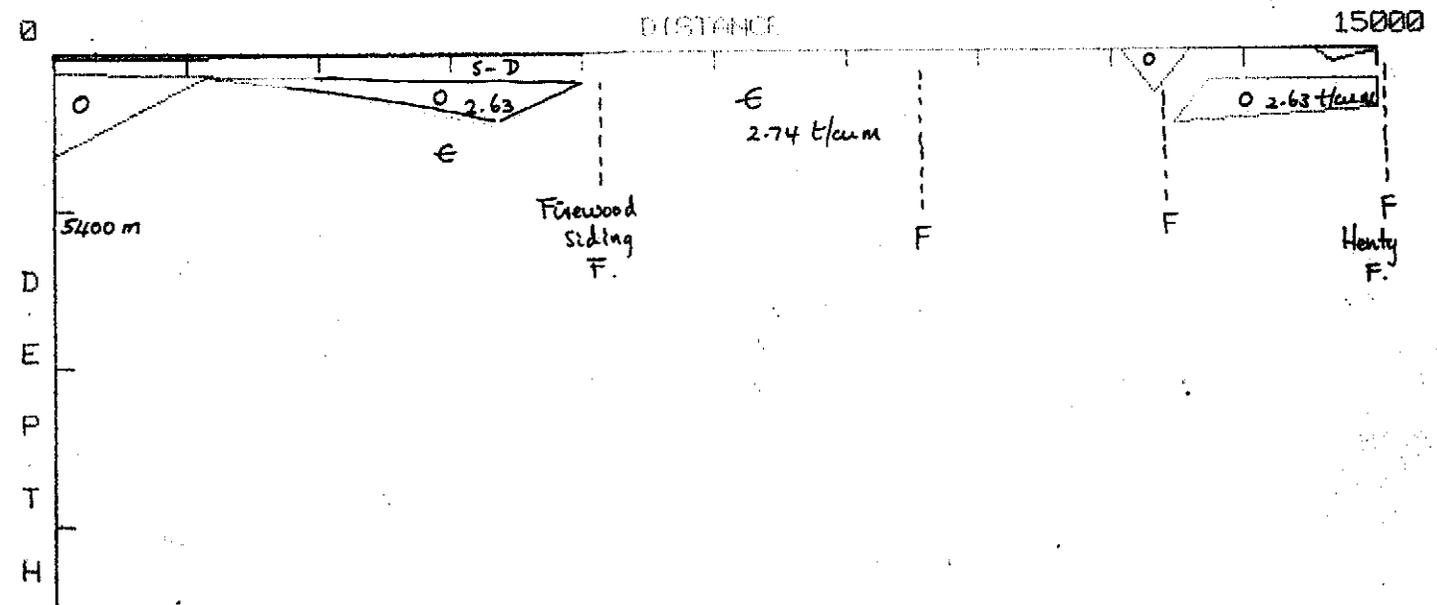
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LINE PARAMETERS - ORIGIN, LIMIT, INCR : 0 15000 1000



HENTY RIVER 5345N 358-373E  
ADJ 8 + 10 ORD IN FLT BLOCK S OF FAULT



027

# LEAMAN GEOPHYSICS

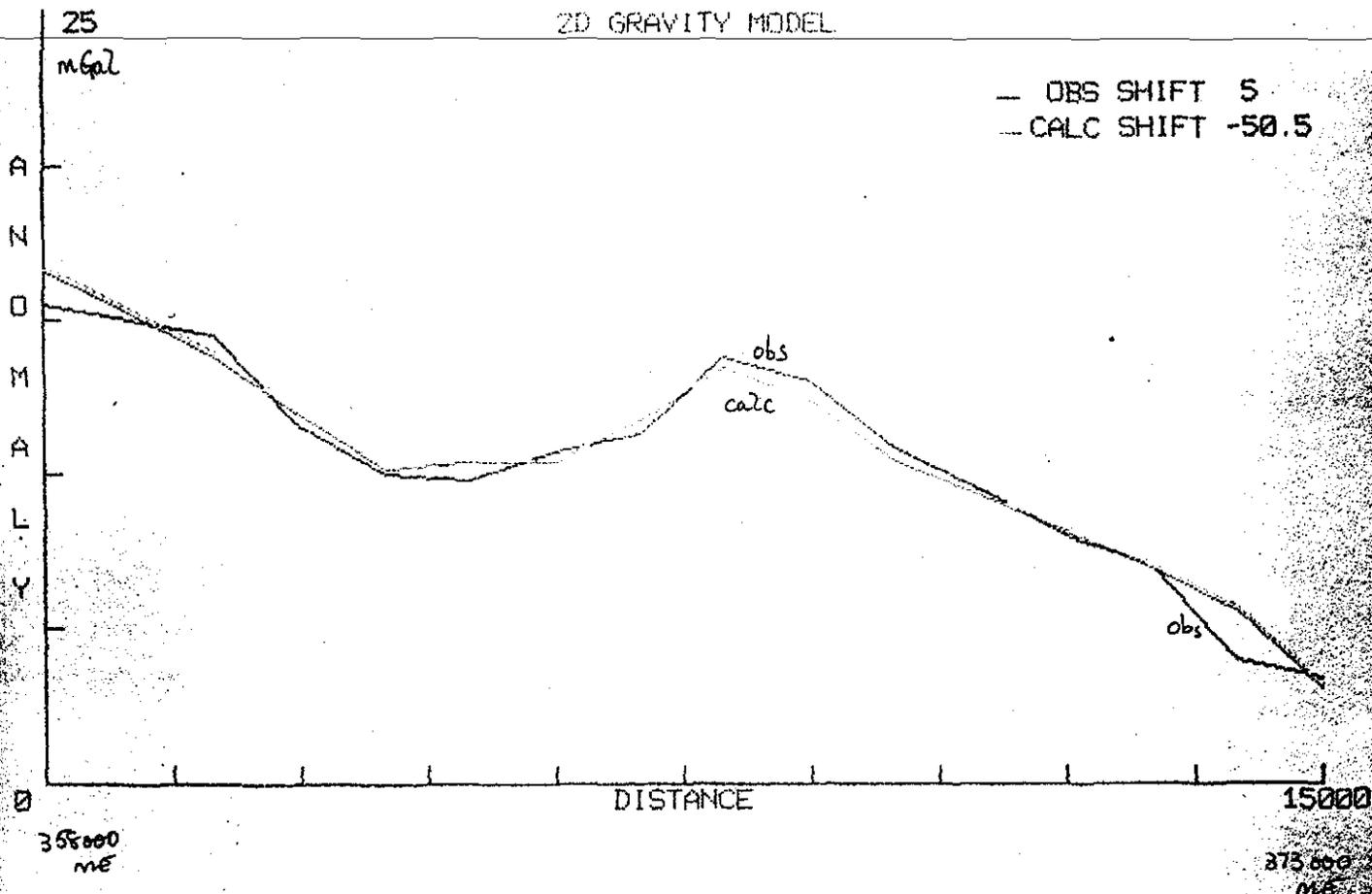
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844028

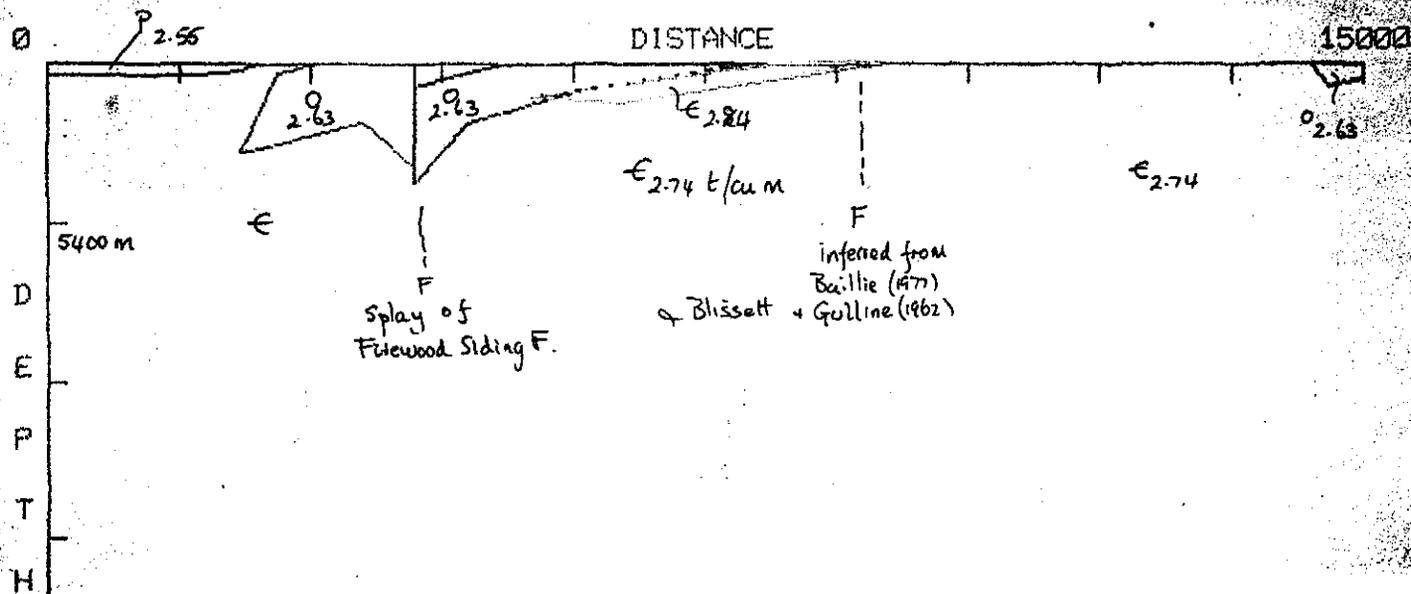
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HENTY RIVER 5348N 358-373E  
ADJ 10 K5=0.065 FOR E CONTRAST

## 2D GRAVITY MODEL



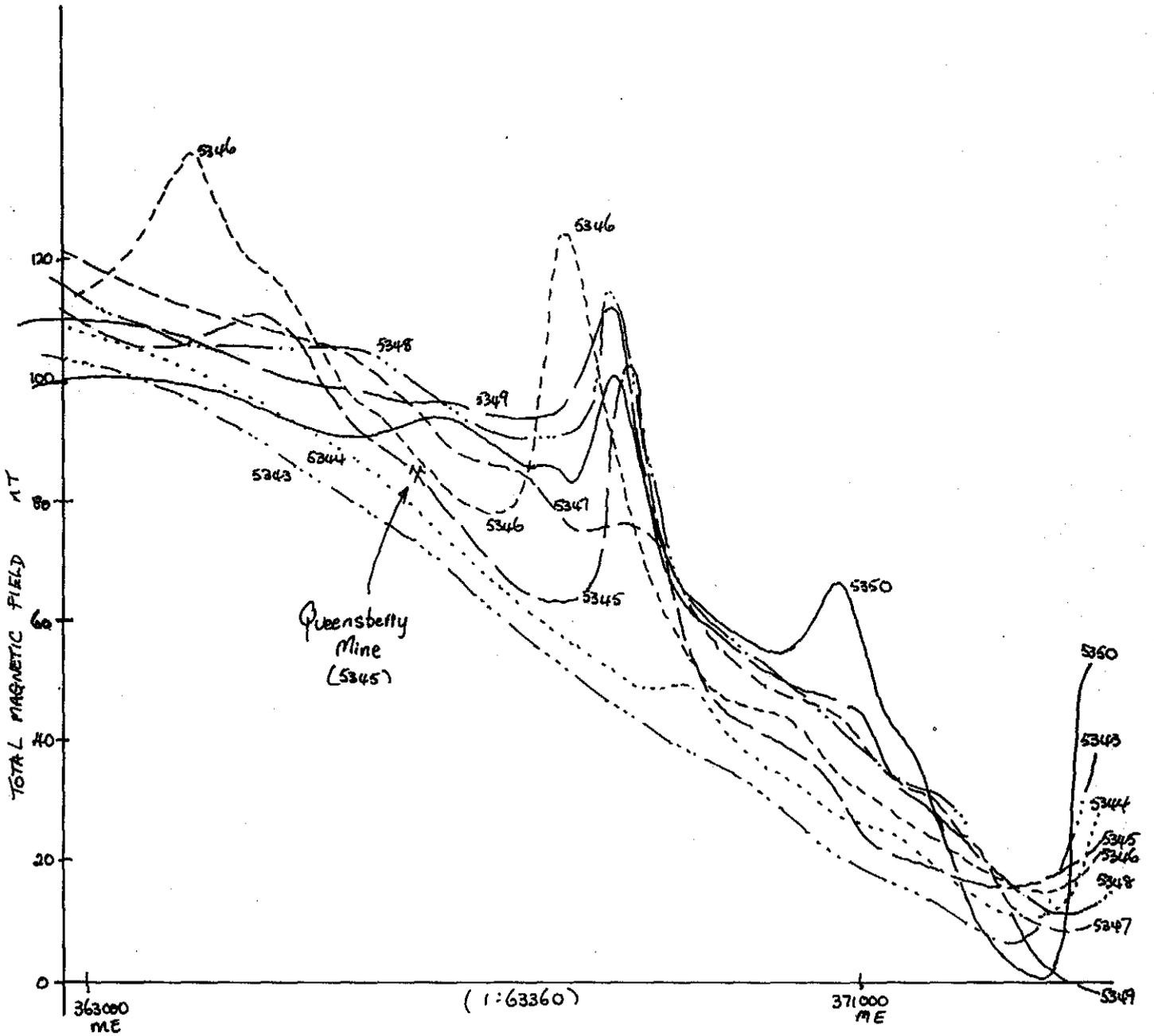
HENTY RIVER 5348N 358-373E  
ADJ 10 K5=0.065 FOR E CONTRAST



EL 10/85 GRAVITY INTERPRETATION 5348 000 MN FIGURE 12

028

- 5343 - - - - -
- 5344 ······
- 5345 ————
- 5346 - - - - -
- 5347 ————
- 5348 ······
- 5349 ————
- 5350 ————



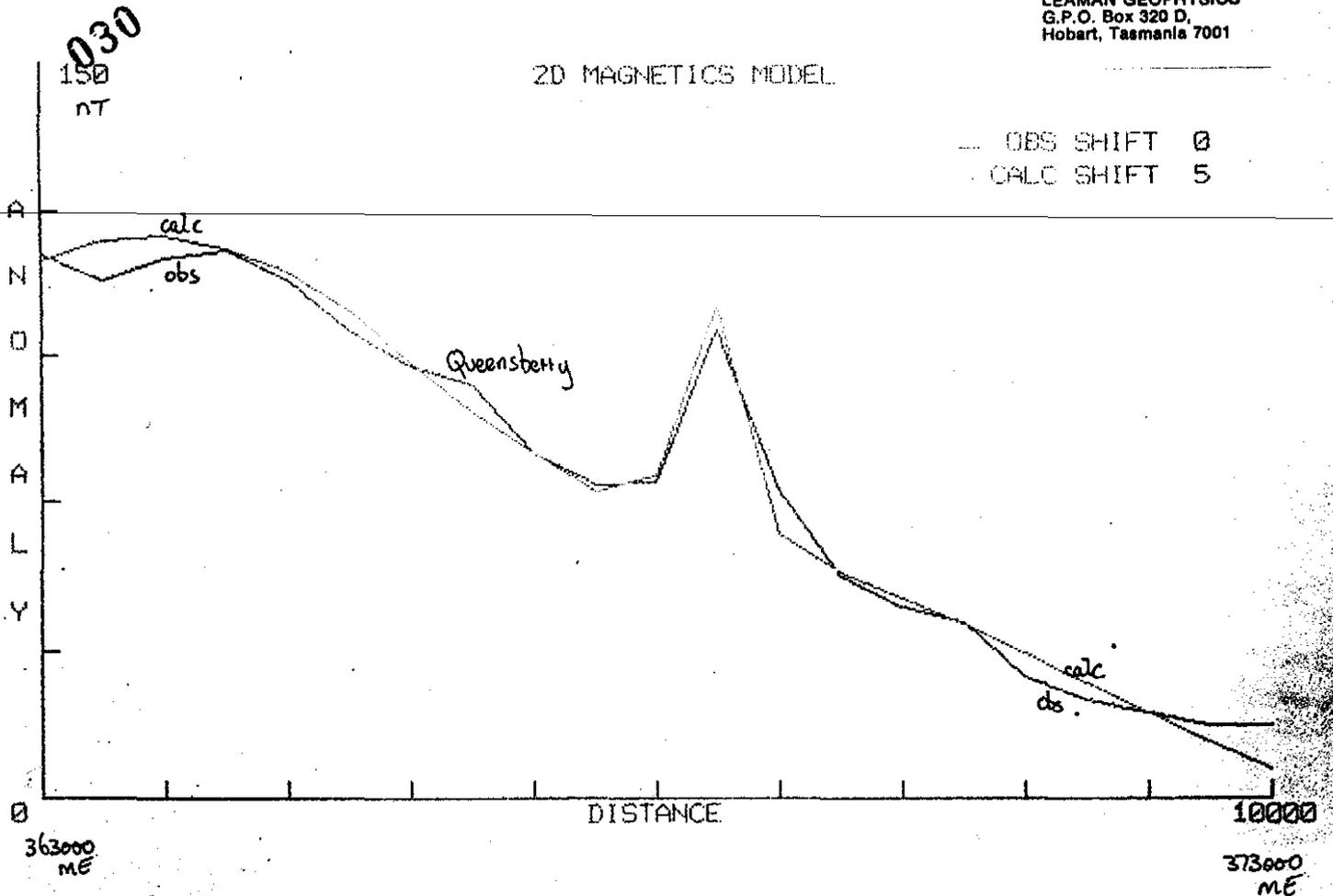
EL 10/85 OBSERVED MAGNETIC PROFILES  
 (1981 Mines Department aeromagnetic survey)

FIGURE 13

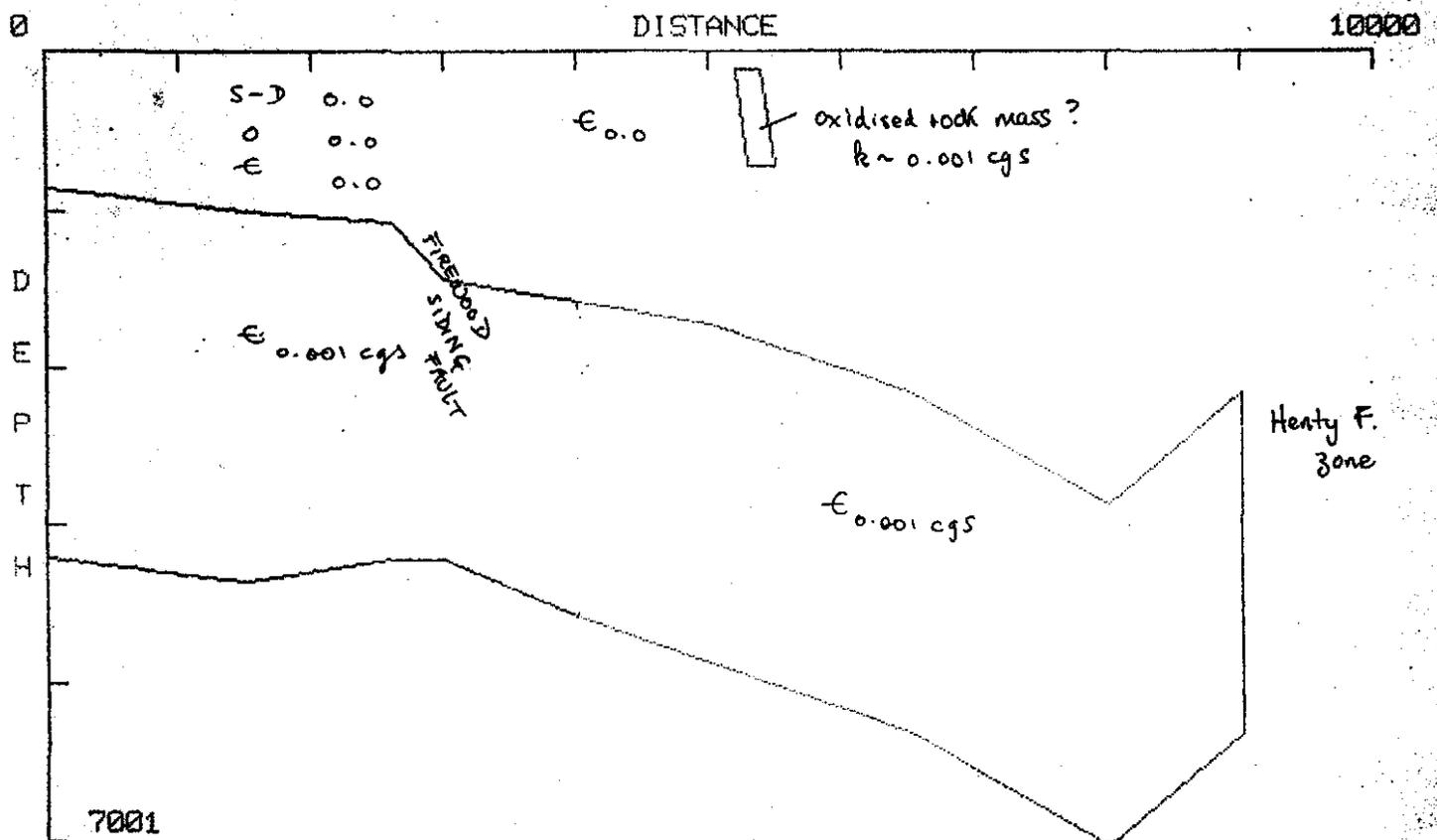


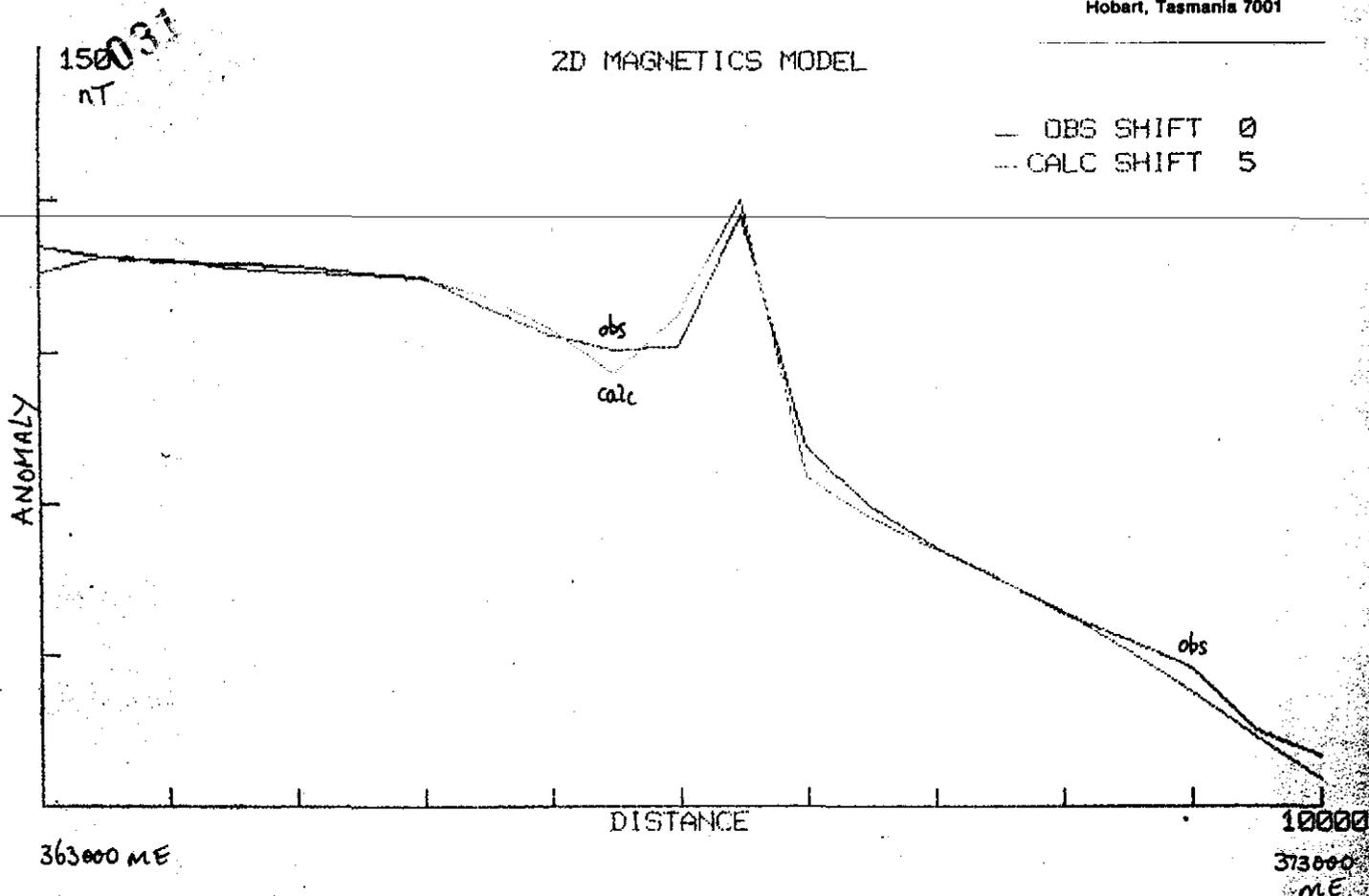
LINE PARAMETERS - ORIGIN, LIMIT, INCR : 0 10000 500

LEAMAN GEOPHYSICS  
G.P.O. Box 320 D,  
Hobart, Tasmania 7001

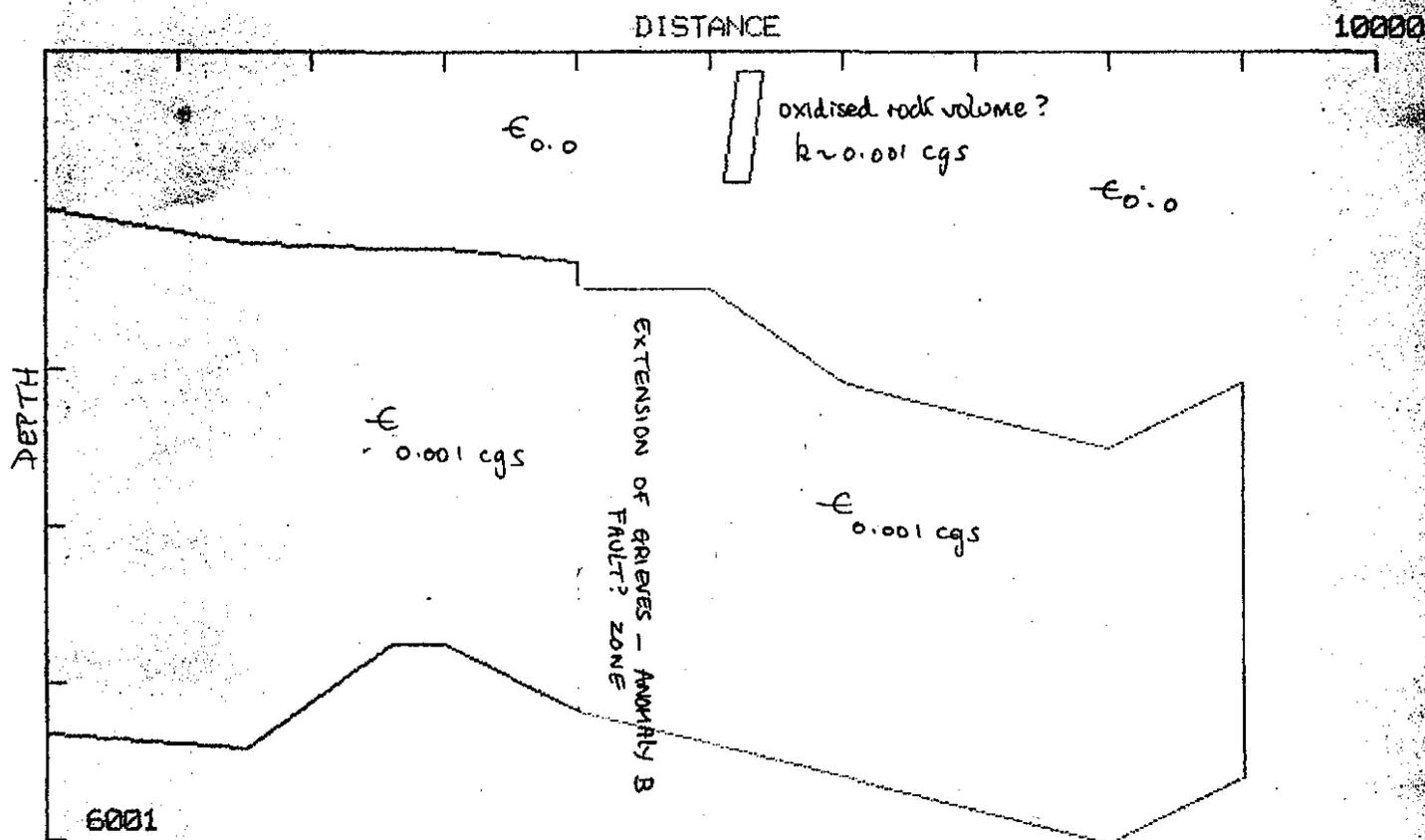


HENTY RIVER 5345N 363-373E  
ADJ K4 TO 00075





ENTY RIVER 5348N 363-373E  
D.J. 13 FAULT AREA



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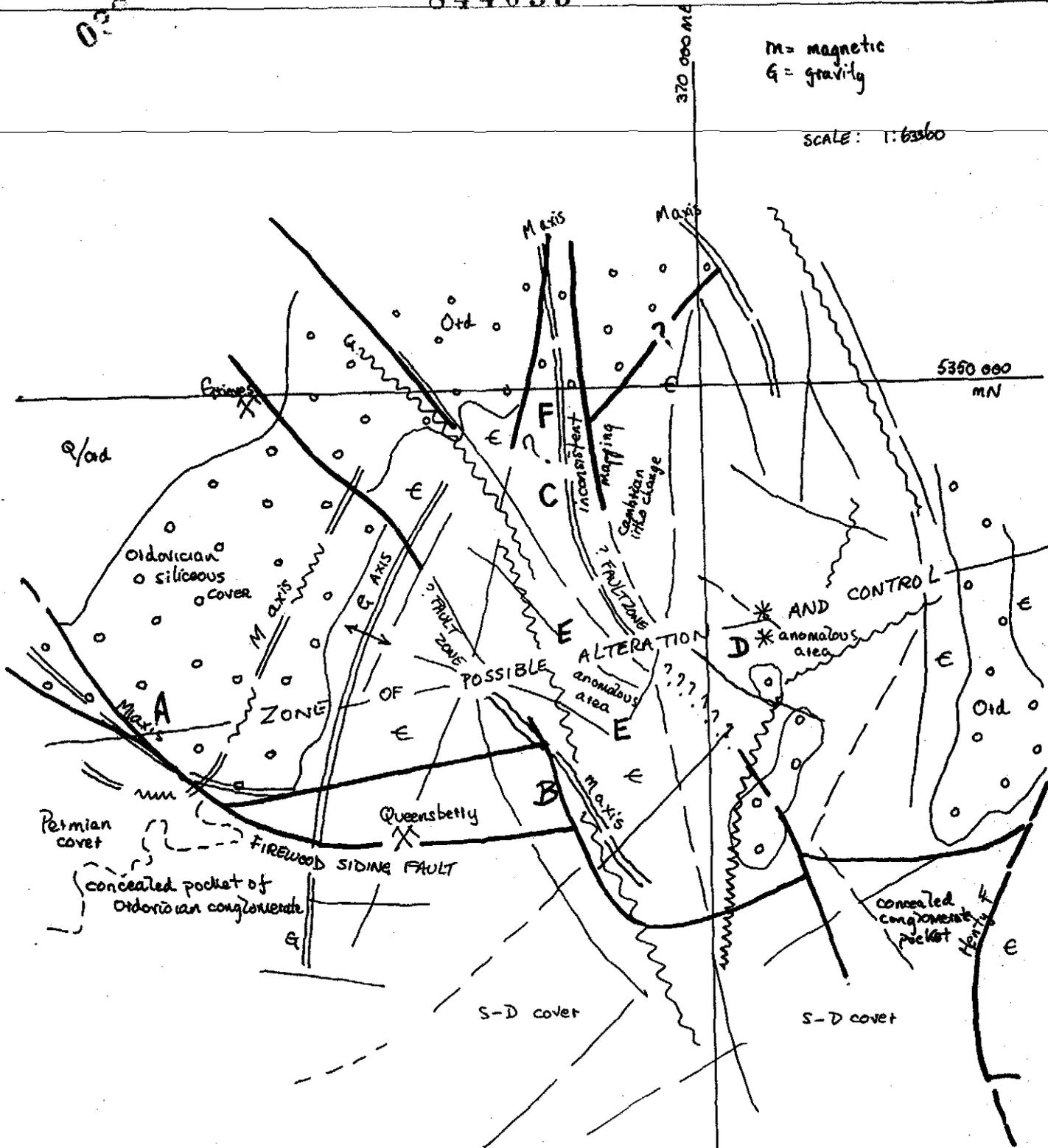
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844033

030

m = magnetic  
G = gravity

SCALE: 1:63560



## EL 10/85 INTERPRETATION SUMMARY

Mapping base from Blissett and Gulline (1962) and Baillie (1977)  
Trends inferred from gravity and magnetic data shown; symbol density suggests possible relative significance.

Anomalous areas A, B, C probably represent oxidation effects but may be mineralised. No explanation is offered for anomalous areas D, E and these sites should be reviewed. Mapping is required near F to resolve mismatch at sheet edges.

Further work should be concentrated along the near E-W control zone and features which abut it (A to E).