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Tasmanian Basemetals Project

E.L. 32/94

Gowrie Park

Vol 1 of 1

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HELD BY: Renison Limited

MANAGER & OPERATOR: RGC Exploration Pty Ltd

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**AUTHOR(s):
R. Stockwell**

10 December, 1997

PROSPECTS: Gowrie Park

| | | |
|--------------------------|--------------------------------|--------------------------------|
| MAP SHEETS: | 1:25,000: Cethana | 1:100,000: Forth |
| GEOGRAPHIC COORDS | Min East: 432 000 mE | Max East: 438 000 mE |
| | Min North: 5 406 000 mN | Max North: 5 409 000 mN |

COMMODITY(s): Cu, Pb, Zn, Ag, Au

KEY WORDS: Massive sulphide, geochemistry, geophysics, alteration

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ANNUAL REPORT - EL 32/94
RGC EXPL. GOWRIE PARK

SUMMARY

Exploration this period has focussed on defining a drill hole (GPDD1) which best tests those geophysical/geochemical anomalies defined by previous exploration, and the rhyolite-andesite contact (inferred sea-floor position) at a depth of greater than 200m. In-fill soil sampling was only possible over the VLF-EM anomaly in the northeast of the grid, due to a thick scree cover, and did not repeat encouraging results from previous surveys.

Future exploration will be dependant on the results of GPDD1. If encouraging results are obtained from drilling, exploration should continue to focus on mineralisation hidden from previous exploration techniques, such as the rhyolite - andesite contact at depth and the base of, and beneath Mt Claude.

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1. INTRODUCTION

EL 32/94 - Gowrie Park is held by Renison Limited and is explored by RGC Exploration, both wholly owned subsidiaries of RGC Limited. The licence is located in northern Tasmania approximately 34 km south of Devonport, and is situated on the northwestern slopes of Mount Roland, Mount Van Dyke and Mount Claude (Fig.1). It was granted on January 13, 1995 and covers an area of 18 sq km.

Gowrie Park is a disused Hydro Electric Commission village and is approximately 10 km southwest of Sheffield, the nearest major township. The major access to the EL from Sheffield is via Claude Road.

The topography of the EL is dominated by the steep northern and western slopes of Mount Claude, Van Dyke and Roland. These form an elevated ridge that defines the southern and eastern boundaries to the EL. The Dasher River flows in a broad valley in the northwest corner of the lease and O'Neills Creek forms a major drainage in the gully between Mount Roland and Mount Claude.

The vegetation consists predominantly of medium eucalypt forest and light tea tree scrub with some patches of rainforest. The more fertile soils of the Dasher Valley have been cleared and support farming activities.

The area was acquired for its potential to host Rosebery style Cu-Pb-Zn-Ag mineralisation.

2. TENURE

The EL comprises:

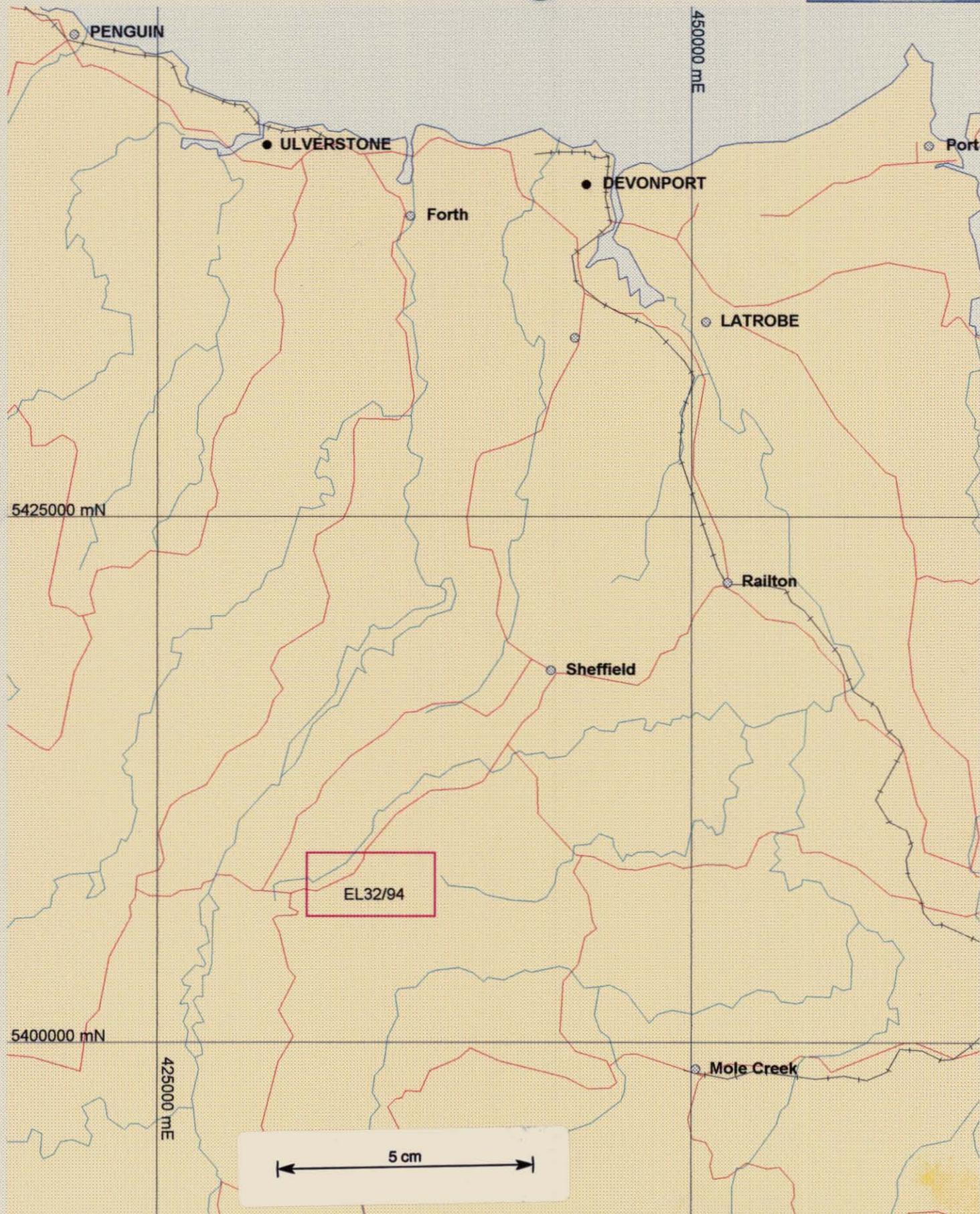
- Private Property
- Crown Land (Deferred Forest Land)
- Crown Land
- Land Vested in HEC.

3. PREVIOUS EXPLORATION

3.1 INTRODUCTION

Gowrie Park consists of a sequence of Cambrian andesitic - rhyolitic lavas and volcanoclastics of the Mt Read Volcanic Belt. It has been explored since 1973 for VHMS mineralisation. Within Gowrie park a sericite alteration zone with a coincident poorly defined Pb - Zn soil anomaly within rhyolitic volcanics has been the focus for much of the previous exploration. The previous exploration history of EL 32/94 prior to 1995 has been summarised by Halley & Vicary (1996).

Gowrie Park Location Plan Figure 1



3.2 RGC EXPLORATION JANUARY 1995 - JANUARY 1996

Exploration during this period consisted of a literature review of previous exploration, relogging of old diamond drill holes, the establishment of a 37 line km grid which was subsequently mapped, rock chip sampled, soil sampled and covered with ground magnetics. Previous exploration by CRA defined a zone of galena - sphalerite stockworking near the western EL boundary. RGC Exploration aims were to locate the seafloor position during the time sub-seafloor lead - zinc stockwork mineralisation was occurring.

Mapping divided the geology into three parts, a rhyolitic dominated sequence at the western end, an andesitic dominated sequence at the north-eastern end, and a mixed volcanoclastic sequence in between (plan 1).

The rhyolites are massive quartz - feldspar phyric lavas and associated breccias with minor volcanoclastics, they are weakly sericitised and the feldspar phenocrysts have been replaced by pale green sericite. An unusual coarse grained equigranular quartz - feldspar rhyolite intrusive occurs within the sequence.

In the middle and eastern parts of the grid fine grained, aphyric chloritic rocks occur within the rhyolite sequence and are interpreted to represent dacitic volcanoclastics.

At the eastern end of the grid the rhyolites are overlain by a package of fine to coarse grained quartz phyric volcanic sandstones. The coarser facies was noted to include rounded Precambrian quartzite clasts which have a strong resemblance to the upper Tyndall Group.

Two zones of quartzite were recognised at the eastern end of the grid, which resemble Precambrian quartzites.

The andesites occur at the eastern end of the grid and lithologies recognised include conglomerates, sandy volcanoclastics and coherent lavas. The andesites are very similar to the Beulah Formation in terms of chemistry and facies associations.

In the south-eastern part of the grid is a mixed zone of andesitic and rhyolitic lithologies.

Due to the lack of bedded sediments in the grid area, the facing direction is not known.

Moderate to strong sericite alteration occurs over the tested CRA anomalies at the western end of the grid. Weaker patchy sericite alteration is more or less continuous around the southern edge of the andesites at the base of Mount Van Dyke.

3.

Although not a zone of strong alteration, an asymmetrical alteration occurs at the andesite - rhyolite contact. This distinct change in volcanism may represent the sea floor position at the time the base metal stockwork mineralisation was occurring.

Soil samples and rock chips were collected over the grid in areas that were not covered by scree. The samples were analysed for base metals and a suite of other elements by Neutron Activation Analysis (NAA). The NAA results are useful as a guide to mapping lithological boundaries in areas of poor outcrop. Base metal analysis better defined the known CRA soil anomaly to the north and also picked up a small anomaly on the south end of line 435600mE. A program of infill soil sampling over this area failed to detect any anomalism.

A ground magnetic survey illustrated the pronounced change in geology from east to west, from andesitic to rhyolitic volcanics. The boundary between the two magnetic units is marked by a major north-west trending trough which may be the northern extension of the fault that offsets the Roland Conglomerate between Mt Van Dyke and Mt Roland. The magnetic survey also illustrated the extent of the andesite to the west at the northern end of the grid under cover.

Based on the lead and zinc geochemistry and corresponding moderate sericite alteration within the rhyolites near the andesite contact, it was recommended that a series of short drill holes be drilled to test the along strike potential of a possible sea-floor position, and that each hole should be subsequently tested by DHEM.

3.3 RGC EXPLORATION JANUARY 1996 - JANUARY 1997

Exploration during this period consisted of two soil sampling programmes, the drilling of 6 short RC drillholes and the cutting of 3.4 kilometres of infill gridlines to facilitate a comprehensive geophysical coverage. Geophysical surveys included an infill Self Potential survey, a Gradient Array IP survey and a VLF-EM survey.

The northern extensions of 2600E, 3000E and 3400E were sampled to test the strike extension of anomalous soil assay results from grid line 3200E. Only 22 samples were taken by auger, during the first soil programme, due to thick unconsolidated scree. Base metals concentrations were low, and inspection of the anomalous zone on 3200E revealed a sub-surface drainage channel, which may indicate the anomalous assay results were from transported material.

The second soil sampling programme was designed to test mobile metal ions in the scree covered areas of 2600E, 3000E and 3400E.

4.

The results suggested that no anomalous base metals mineralisation exists beneath the scree (Gregory, 1996). However, this method has not been proven in temperate climates, having been developed for intensely weathered soil profiles in arid regions. The results therefore were considered inconclusive.

The position and dip of the inferred sea-floor position, the contact between footwall rhyolite and hangingwall andesite, was located below scree between 3000E and 3300E by RC drilling. The contact between the unaltered andesite and the apparently underlying, variably sericite-pyrite-altered rhyolite, dips 80 degrees to the north. Base metals from drill-chip assays in both the andesite and rhyolite were low, however base metals concentration and alteration intensity appear to increase from west to east.

Results from the Self Potential survey were inconclusive. All anomalous zones were coincident with scree slopes and/or topographic highs. It was considered likely that the SP response may have been caused by the elevation differential between topographic highs and base stations, and the resultant streaming potential of water between the two locations, or electrochemical reactions between scree and underlying bedrock, or both.

Gradient Array IP and VLF-EM surveys delineated two possible zones of mineralisation (Dauth, 1996). The first occurs over the rhyolite-andesite contact, proximal to the inferred Cambrian growth fault at the base of Mt Claude, on lines 5000E-5400E, 6600E-6900E, and is coincident with a large zone of de-magnetisation. The second occurs in the northeast of the grid in an andesite dominated area on lines 5600E-6400E, 8090N-7920N.

A low apparent resistivity anomaly was defined by the Gradient Array IP survey, coincident with the CRAE UTEM anomaly drilled in 1984 (drill hole number RD84GPC4). The anomaly extends from line 3600E to 4600E on a northing of 7750N, it overlies the inferred sea-floor position, and lies in the direction (vector) of increased mineralisation potential determined by RC drilling to the west. This zone represents the most conductive area encountered during the IP survey, however a coincident apparent chargeability response was absent.

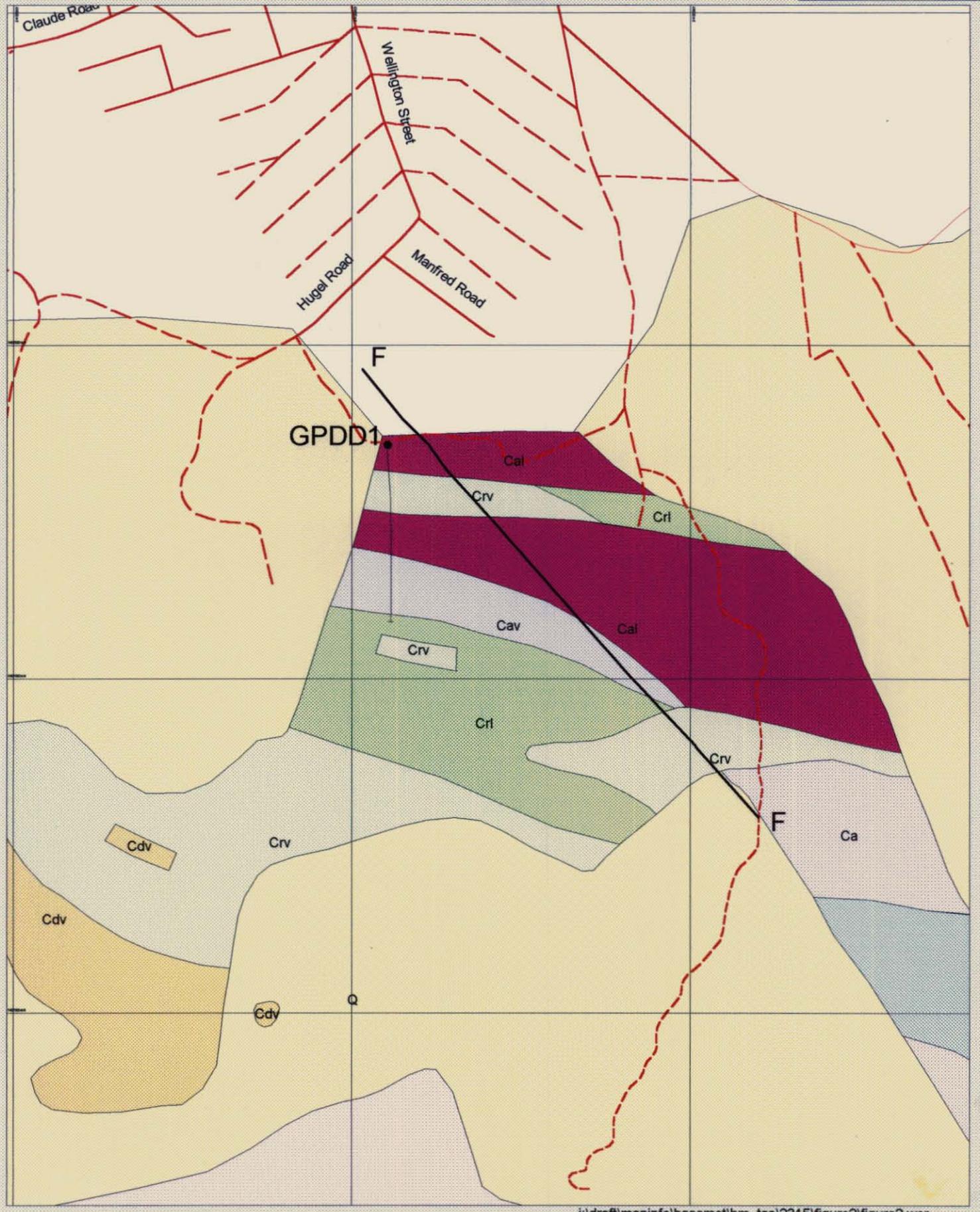
Recommended for follow-up were both VLF-EM anomalies, the andesite-rhyolite contact at depth (>200m) and the low apparent resistivity anomaly coincident with the CRAE UTEM anomaly.

4. WORK COMPLETED

In the period January 1997 - January 1998 the work completed in EL 32/94 - Gowrie Park includes the following:-

- 1) A review of past and future exploration, by Wally Herrmann (Appendix 2)

Gowrie Park Geology Figure 2



5.

- 2) 43 soil samples
- 3) One diamond drill hole - GPDD1 was planned.

5. RESULTS

5.1 SOIL GEOCHEMISTRY

In-fill sampling of lines; 5800E, 7600N-8150N and 6200E, 7600N-8200N was conducted to further test the Gradient Array IP/VLF-EM anomaly in the northeast region of the Gowrie Park grid. Sampling of in-situ soils was conducted at 25 metre intervals. Assay results were low (<23ppm Cu, <44ppm Pb & <128ppm Zn) and did not repeat encouraging results in the same area from lines 6000E and 6400E (<1264ppm Cu, <169ppm Pb & <1392ppm Zn).

Sampling of the IP/VLF-EM anomaly at the base of Mt Claude (5000E-5400E, 6600E-6900E) could not be performed due to a thick coverage of Owen Conglomerate scree.

5.2 DIAMOND DRILLING

One diamond drill hole (GPDD1) is planned for December, collared at 5407850mN, 434525mE, and drilled at 65/177 degrees. Core orientation is recommended for GPDD1 to constrain the structural character of the area. The target is the andesite-rhyolite contact (inferred sea-floor position) at a depth of 300m. This target has been chosen for its proximity to the inferred Cambrian growth fault (~300m), its position within an apparent resistivity low which is coincident with a CRAE UTEM anomaly, and the presence of a large de-magnetised zone, apparent on ground magnetic TMI. DHEM will be performed on completion of this diamond drill hole.

6. RECOMMENDATIONS

Geological features at Gowrie Park common to VHMS systems are;

- 1) a bimodal volcanic package
- 2) characteristic sericite-pyrite footwall? alteration
- 3) jasper - quartz - carbonate veining in the footwall? rhyolites
- 4) occurrence of albite alteration in the hangingwall? andesites
- 5) an asymmetrical alteration halo
- 6) known base metal stringers of Cambrian age in the footwall? rhyolites.

6.

Consideration of the above features and the extent of the alteration from Gowrie Park to Cethana suggests that a substantial hydrothermal system was operative.

Previous exploration efforts have focused along strike from the CRA known mineralisation and along strike of the andesite - rhyolite contact and have failed to locate significant mineralisation close to the surface.

It is proposed that exploration focus on those areas hidden from previous exploration techniques, such as at depth and underneath the Roland Conglomerate of Mt Claude.

Exploration this period focused on the definition of a conceptual drill target to best test the andesite - rhyolite contact, which is interpreted to be the seafloor position at the time the rhyolite-hosted stringer mineralisation was depositing. Geophysical techniques, RC drilling and soil geochemistry have failed to outline any obvious near surface targets, forcing exploration to look below the well-tested upper 200m.

Further work at Gowrie is dependant on the results from GPDD1 and the subsequent DHEM survey. Two exploration programs are proposed for Gowrie Park based on good results from GPDD1.

The first and priority target will be continued drilling of the andesite - rhyolite contact at depth in order to test the continuity of mineralisation, or vectors to mineralisation from the alteration assemblage.

Of equal importance, if alteration is coincident with the demagnetised zone in GPDD1, would be the drilling of the VLF-EM anomaly at the base of Mt Claude, on lines 5000E-5400E, 6600E-6900E. This site is also proximal to the fault, and lies within a much larger/more intense zone of demagnetisation.

A mass-balance study of alteration at Gowrie would be appropriate (based on the recommendations of Herrmann, 1997: Appendix 2). The volume-comparison of major element species between least- and most-altered rock types (especially the footwall? rhyolite) at Gowrie will define the intensity of alteration, and hence, the likelihood of mineralisation.

Multiple rock samples of the unaltered protoliths should be collected to be averaged for comparison with altered rocks. GPDD1 will provide un-weathered rocks necessary for the study, however other rocks from surface outcrop may have to be collected to ensure all pertinent rocktypes at Gowrie are represented (care should be taken to avoid weathered material). The method of Gresens (1967) is recommended, in conjunction with the isocon (linear relationship) method of Grant (1986) or Baumgartner and Olson (1995).

7. REFERENCES

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APPENDIX 1

GOWRIE PARK SOIL SAMPLES

| Sample No. | Grid N | Grid E | Au (ppm) | Au(R) | Cu (ppm) | Pb (ppm) | Zn (ppm) | Ag (ppm) |
|------------|--------|--------|----------|-------|----------|----------|----------|----------|
| W263235 | 7600 | 6200 | X | - | 4 | 6 | 23 | 1 |
| W263236 | 7650 | 6200 | X | - | 16 | 4 | 57 | 2 |
| W263237 | 7675 | 6200 | X | - | 14 | 9 | 54 | 2 |
| W263238 | 7700 | 6200 | X | - | 14 | 14 | 79 | 1 |
| W263239 | 7725 | 6200 | X | - | 19 | 12 | 100 | 2 |
| W263240 | 7750 | 6200 | X | - | 23 | 14 | 102 | 2 |
| W263241 | 7775 | 6200 | X | - | 13 | 13 | 128 | 3 |
| W263242 | 7800 | 6200 | X | - | 14 | 15 | 63 | 2 |
| W263243 | 7825 | 6200 | X | X | 15 | 5 | 66 | 2 |
| W263244 | 7850 | 6200 | X | - | 13 | 12 | 55 | 3 |
| W263245 | 7875 | 6200 | X | - | 15 | 5 | 83 | 3 |
| W263246 | 7900 | 6200 | X | - | 17 | 4 | 53 | 3 |
| W263247 | 7925 | 6200 | X | - | 15 | 6 | 67 | 3 |
| W263248 | 7950 | 6200 | 0.03 | - | 14 | X | 57 | 3 |
| W263249 | 7975 | 6200 | 0.01 | - | 16 | 7 | 59 | 3 |
| W263250 | 8000 | 6200 | 0.01 | - | 21 | 5 | 65 | 2 |
| W263251 | 8025 | 6200 | 0.01 | 0.01 | 18 | X | 52 | 2 |
| W263252 | 8050 | 6200 | X | - | 20 | X | 51 | 3 |
| W263253 | 8075 | 6200 | 0.01 | - | 15 | 22 | 39 | 2 |
| W263254 | 8100 | 6200 | 0.01 | - | 21 | 19 | 46 | 3 |
| W263255 | STD | | 0.32 | 0.38 | >10000 | 19 | 40 | 6 |
| W263256 | 8125 | 6200 | X | - | 21 | 15 | 51 | 3 |
| W263257 | 8150 | 6200 | X | - | 18 | 11 | 43 | 2 |
| W263258 | 8175 | 6200 | X | - | 16 | 44 | 64 | 1 |
| W263259 | 7700 | 5800 | X | - | 6 | 12 | 34 | 1 |
| W263260 | 7725 | 5800 | X | - | 5 | 24 | 36 | X |
| W263261 | 7750 | 5800 | X | - | 5 | 8 | 39 | X |
| W263262 | 7775 | 5800 | X | - | 7 | X | 25 | X |
| W263263 | 7800 | 5800 | X | - | 7 | X | 14 | 1 |
| W263264 | 7825 | 5800 | X | - | 8 | 6 | 14 | 1 |
| W263265 | 7850 | 5800 | X | X | 8 | X | 12 | X |
| W263266 | 7875 | 5800 | 0.01 | X | 5 | 3 | 108 | 1 |
| W263267 | 7900 | 5800 | X | - | 7 | 4 | 65 | 1 |
| W263268 | 7925 | 5800 | X | X | 7 | 3 | 23 | X |
| W263269 | 8000 | 5800 | X | - | 6 | X | 26 | 1 |
| W263270 | 8025 | 5800 | X | - | 13 | 11 | 27 | X |
| W263271 | 8050 | 5800 | X | - | 14 | 9 | 40 | X |
| W263272 | 8075 | 5800 | X | - | 7 | 19 | 23 | X |
| W263273 | 8100 | 5800 | X | - | 7 | 12 | 14 | 1 |
| W263274 | 8125 | 5800 | X | - | 6 | X | 7 | X |
| W263275 | STD | | 0.4 | 0.38 | 3551 | >10000 | 2786 | 14 |
| W263276 | 8150 | 5800 | X | - | 9 | 9 | 18 | 1 |
| W263277 | 8175 | 5800 | X | - | 11 | 15 | 25 | 1 |
| METHOD | | | F650 | F650 | A101 | A101 | A101 | A101 |
| DET.LIMIT | | | 0.01 | 0.01 | 2 | 3 | 2 | 1 |
| ANALABS | | | | | | | | |

| Sample No | Grid N | Grid E | Cu (%) | Pb (%) |
|-----------|--------|--------|--------|--------|
| W263235 | 7600 | 6200 | N.A. | N.A. |
| W263236 | 7650 | 6200 | N.A. | N.A. |
| W263237 | 7675 | 6200 | N.A. | N.A. |
| W263238 | 7700 | 6200 | N.A. | N.A. |
| W263239 | 7725 | 6200 | N.A. | N.A. |
| W263240 | 7750 | 6200 | N.A. | N.A. |
| W263241 | 7775 | 6200 | N.A. | N.A. |
| W263242 | 7800 | 6200 | N.A. | N.A. |
| W263243 | 7825 | 6200 | N.A. | N.A. |
| W263244 | 7850 | 6200 | N.A. | N.A. |
| W263245 | 7875 | 6200 | N.A. | N.A. |
| W263246 | 7900 | 6200 | N.A. | N.A. |
| W263247 | 7925 | 6200 | N.A. | N.A. |
| W263248 | 7950 | 6200 | N.A. | N.A. |
| W263249 | 7975 | 6200 | N.A. | N.A. |
| W263250 | 8000 | 6200 | N.A. | N.A. |
| W263251 | 8025 | 6200 | N.A. | N.A. |
| W263252 | 8050 | 6200 | N.A. | N.A. |
| W263253 | 8075 | 6200 | N.A. | N.A. |
| W263254 | 8100 | 6200 | N.A. | N.A. |
| W263255 | STD | | 1.24 | N.A. |
| W263256 | 8125 | 6200 | N.A. | N.A. |
| W263257 | 8150 | 6200 | N.A. | N.A. |
| W263258 | 8175 | 6200 | N.A. | N.A. |
| W263259 | 7700 | 5800 | N.A. | N.A. |
| W263260 | 7725 | 5800 | N.A. | N.A. |
| W263261 | 7750 | 5800 | N.A. | N.A. |
| W263262 | 7775 | 5800 | N.A. | N.A. |
| W263263 | 7800 | 5800 | N.A. | N.A. |
| W263264 | 7825 | 5800 | N.A. | N.A. |
| W263265 | 7850 | 5800 | N.A. | N.A. |
| W263266 | 7875 | 5800 | N.A. | N.A. |
| W263267 | 7900 | 5800 | N.A. | N.A. |
| W263268 | 7925 | 5800 | N.A. | N.A. |
| W263269 | 8000 | 5800 | N.A. | N.A. |
| W263270 | 8025 | 5800 | N.A. | N.A. |
| W263271 | 8050 | 5800 | N.A. | N.A. |
| W263272 | 8075 | 5800 | N.A. | N.A. |
| W263273 | 8100 | 5800 | N.A. | N.A. |
| W263274 | 8125 | 5800 | N.A. | N.A. |
| W263275 | STD | | N.A. | 1.24 |
| W263276 | 8150 | 5800 | N.A. | N.A. |
| W263277 | 8175 | 5800 | N.A. | N.A. |
| METHOD | | | A103 | A103 |
| DET.LIMIT | | | 0.01 | 0.01 |
| ANALABS | | | | |

APPENDIX 2

GEOLOGICAL NOTE - WALLY HERRMANN

Geological Note: Wally Hermann**Gowrie Park EL 32/94**

RGC's mapping indicates that there is an E-W trending contact near Claude Road between altered quartz phyric felsic volcanics (rhyolite and minor dacite) to the south, and andesites to the north. The andesites are magnetic and hornblende phyric, suggesting a possible analogy with the Suite II CVC andesites of the Anthony Road area north of Queenstown, (Crawford et al. 1992). The stratigraphic facing is unknown but younging to north is the favourite guess; the magnetic character of the andesites enables interpretation of the contact from aeromagnetic data. If the Suite II correlation is correct, and these represent a particular magmatic-time unit, it would put the Gowrie Park contact at about the same volcano-stratigraphic level as the Newton Creek spillway favourable horizon, which is locally famous for its massive sulphide clasts. Alternatively, a southward facing could place it at an equivalent to the Henty favourable horizon.

There are patchy Pb & Zn soil geochemical anomalies in the felsic rocks but generally no IP anomalies, close to the contact, which could be construed to reflect footwall alteration zones below a VHMS favourable horizon. CRAE drill tested the strongest IP anomalies several hundred metres south of the contact but intersected only minor mineralisation.

RGC carried out an SP survey which returned booming anomalies (to - 200mV) apparently related to N-S trending topography and spur like talus slopes; interpreted as possibly due to groundwater streaming or reaction between reduced bedrock ground water and oxidised talus derived from Roland Conglomerate. These SP anomalies remain untested because they do not appear to be related to volcanic stratigraphy etc.; I think that is fair enough given the absence of IP and EM responses.

RGC drilled a series of six short RC percussion holes, mainly to pin down the andesite contact and establish its relationship to magnetics. The holes intersected the contact in two places about 200m apart and enabled the interpretation that it dips steeply to the north. "Moderate" alteration including quartz-hematite veins at the contact, exists in the felsic rocks and appears to increase in intensity to the east where pink albite? Alteration occurs in the postulated hangingwall andesites.

This appears to be an alteration "vector" perhaps indicating greater hydrothermal flux and a more proximal zone to the east. The question is: should more holes be drilled into this contact to the east and down dip to map out alteration zonation? The surface geochemical indications, IP and etc. are not very favourable so the shallow potential is reduced and drilling >200m is indicated.

There does not seem to be much enthusiasm in the RGC camp for this track, and I suspect that is due to the rather shaky geological concept and the doubts about it representing a significant favourable horizon. If it were a recognised favourable horizon close to a known deposit (cf: White Spur) there would probably be little reluctance to chase up an apparent alteration vector.

However, unrecognised favourable horizons may still exist and alteration indicates hydrothermal activity. So my inclination would be to follow through and drill a few more moderate depth holes to map out alteration and hopefully refine the vectors and geological setting. My main reservations concern the style of alteration. The apparent increase in intensity of alteration at Gowrie Park appears to be based on visual estimation with no wholerock geochemical data to indicate the degrees (say) of sodium depletion, mass changes and isotopic characteristics. This part of the greater problem of the lack of convincing interpretation of the geological setting and the controls and styles of alteration in the Cethana area generally. From the little core I have seen from Cethana East, it is evident that hematite alteration is widespread and it appears to predate pervasive quartz-sericite-pyrite alteration. This implies that the hematite is related to some very early (diagenetic-oxidation?) phase of alteration, perhaps but not necessarily, associated with the same hydrothermal system which produced ~synvolcanic quartz-sericite-pyrite or that the latter is not synvolcanic.

I would be tempted to have a closer look at the whole rock geochemistry of altered and least altered rocks to determine if alteration at Gowrie Park is as strong as it seems, and that it is the right type to be expected in VHMS systems. It may not be conclusive, in which case you would have to take a punt and drill anyway, but if it turns out to be convincingly unlike VHMS related alteration then it would justify saving some drilling dollars.

